

FCM, UFSC

26 de Novembro de 2014

Aerosols, Clouds and Convection

Henrique Barbosa

Departamento de Física Aplicada

Instituto de Física da USP

hbarbosa@if.usp.br

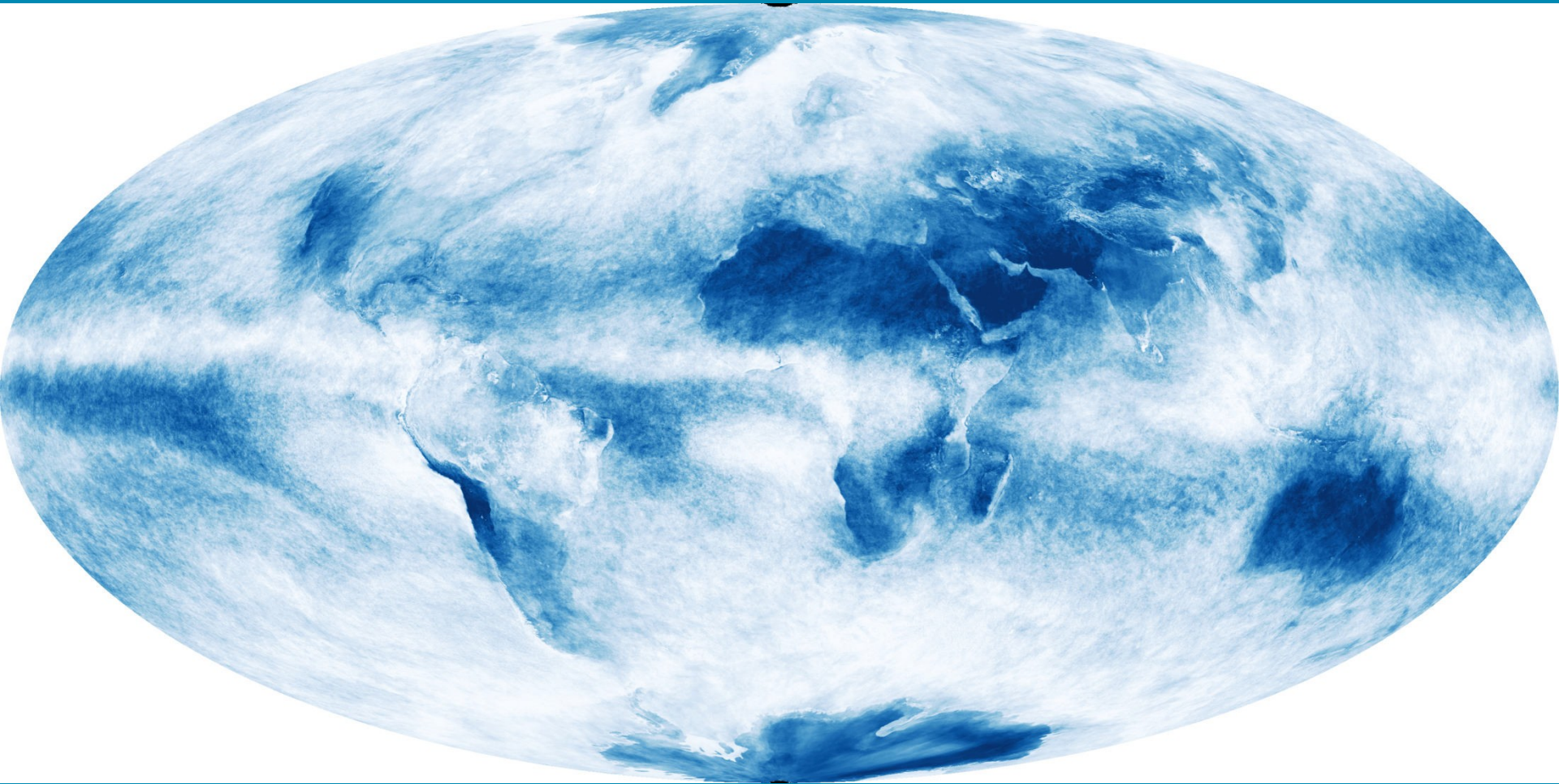
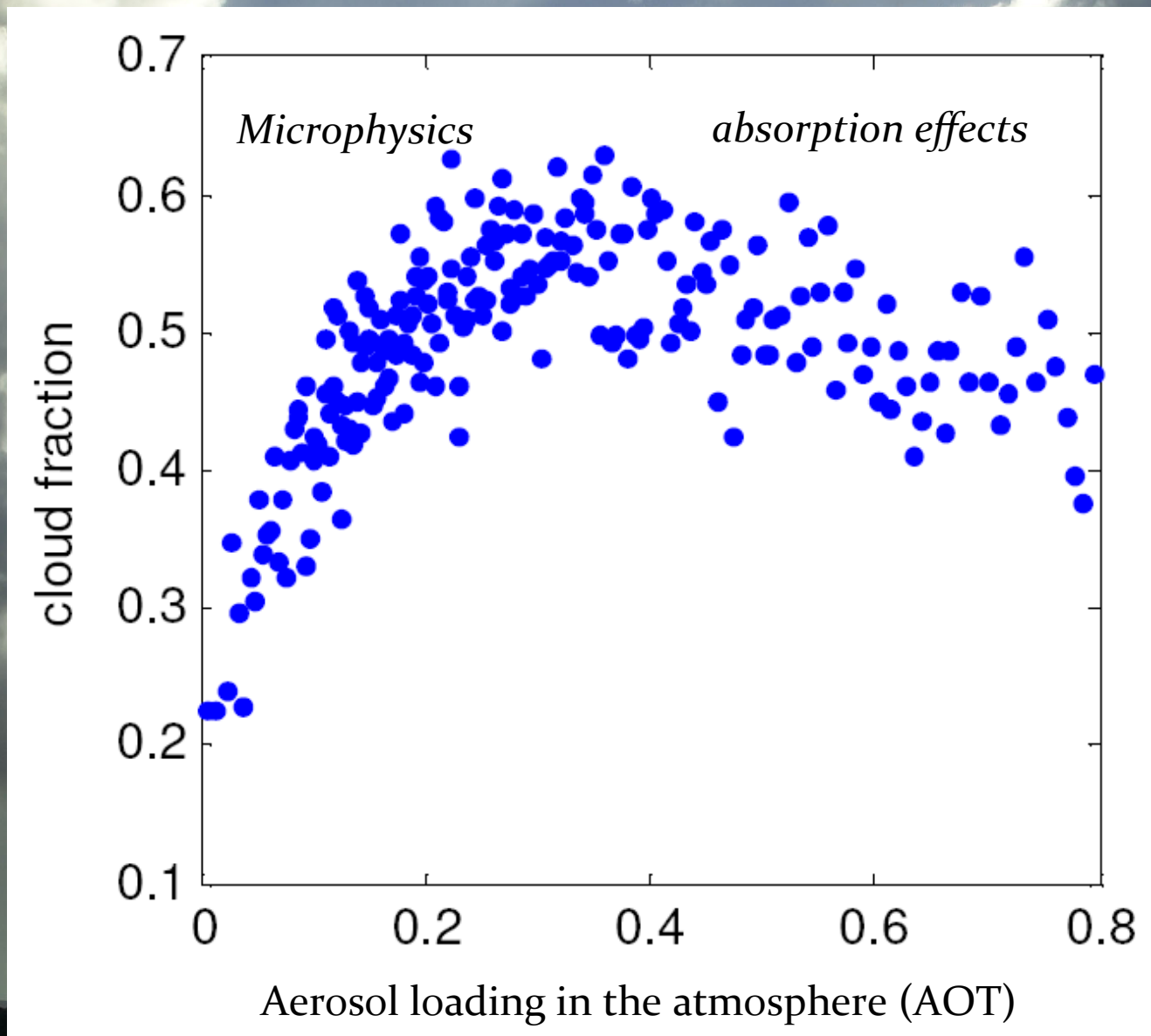
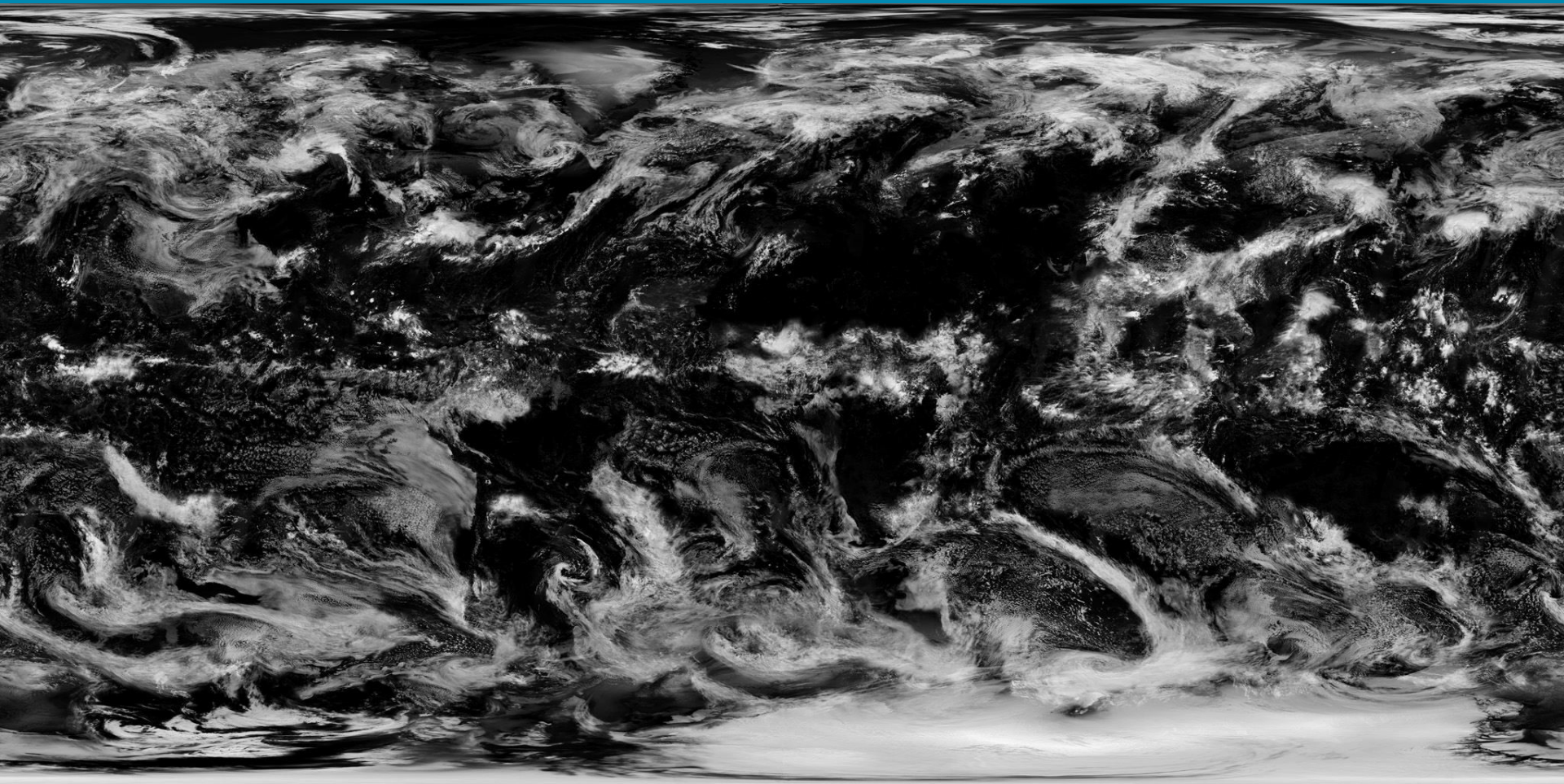




Photo: Ilan Koren

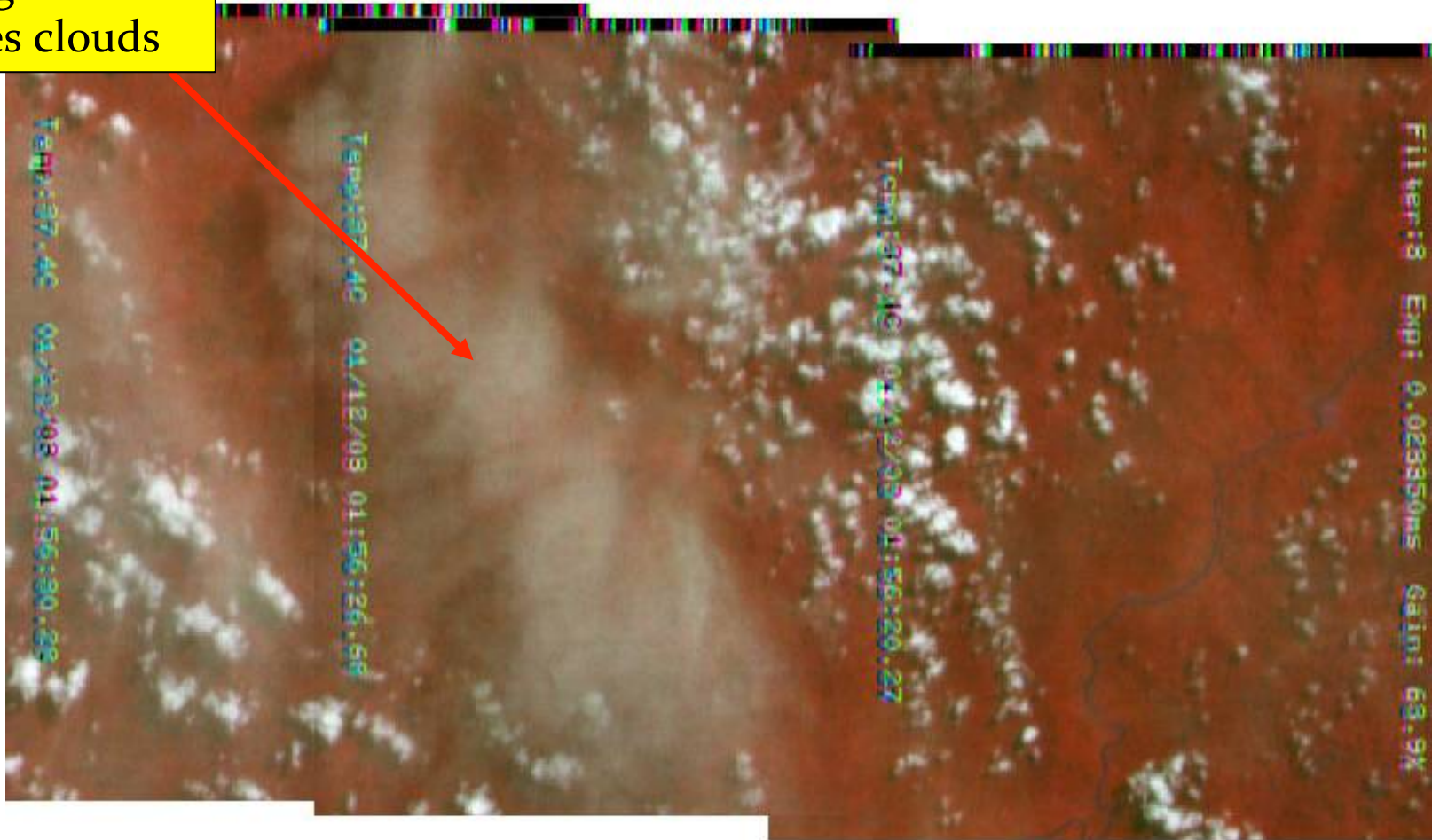
Relationships between cloud properties and aerosol loading in Amazonia





With too much aerosols: Cloud supression

Absorbing aerosol suppresses clouds



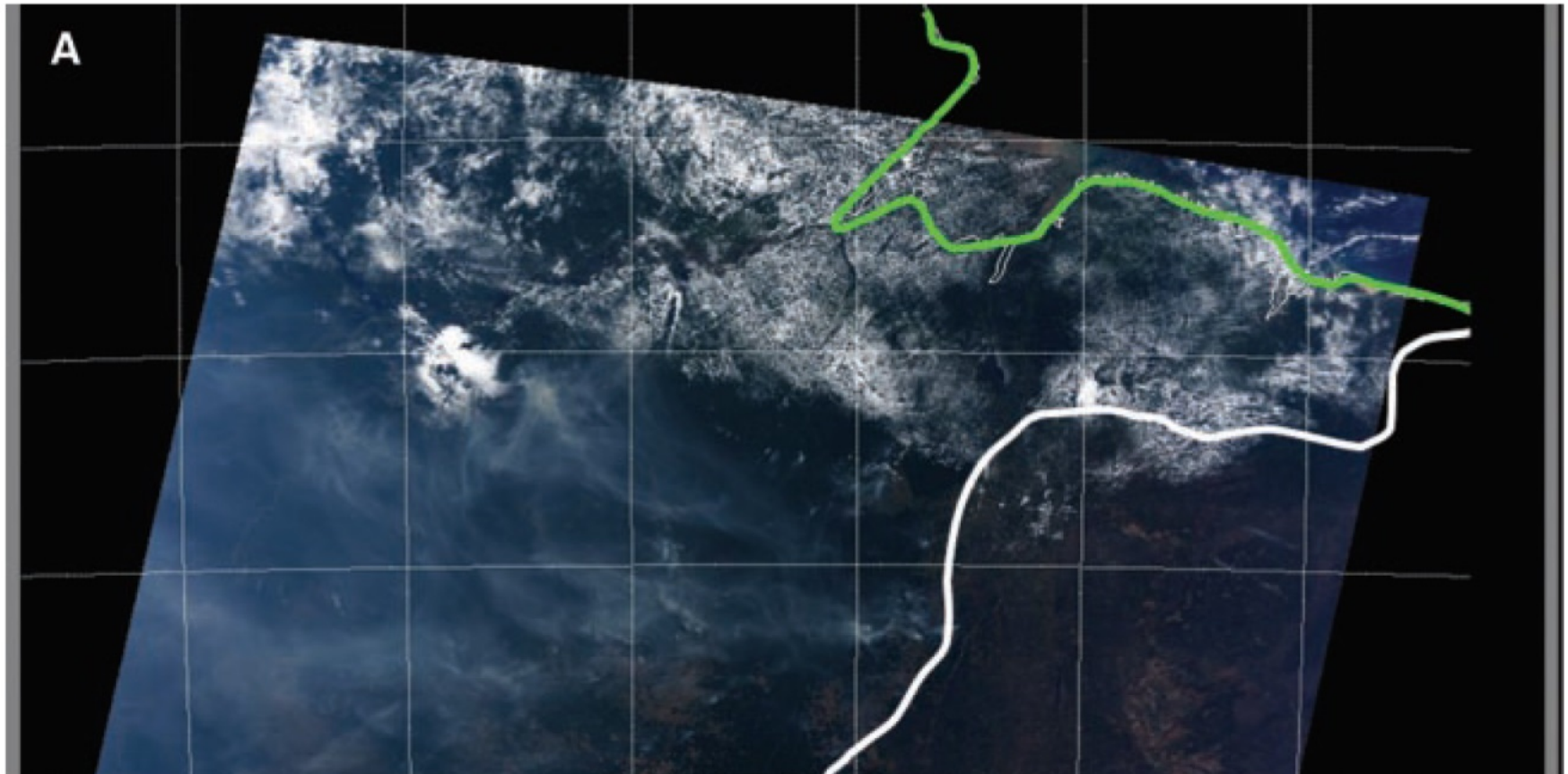
- Stabilization
- Suppression of surface fluxes
- Microphysical influences on droplets

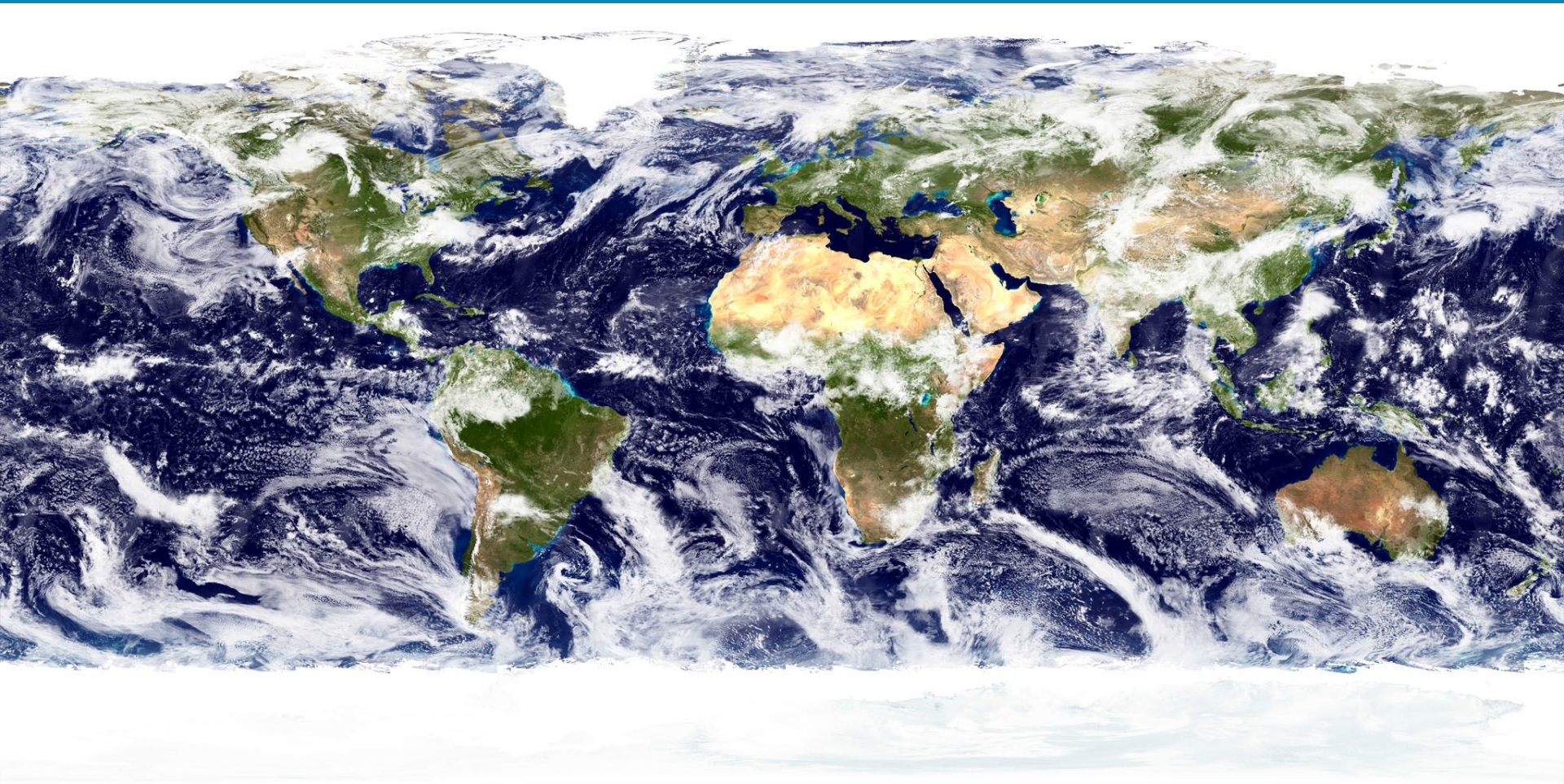
**Columbia
Shuttle
January 2003**



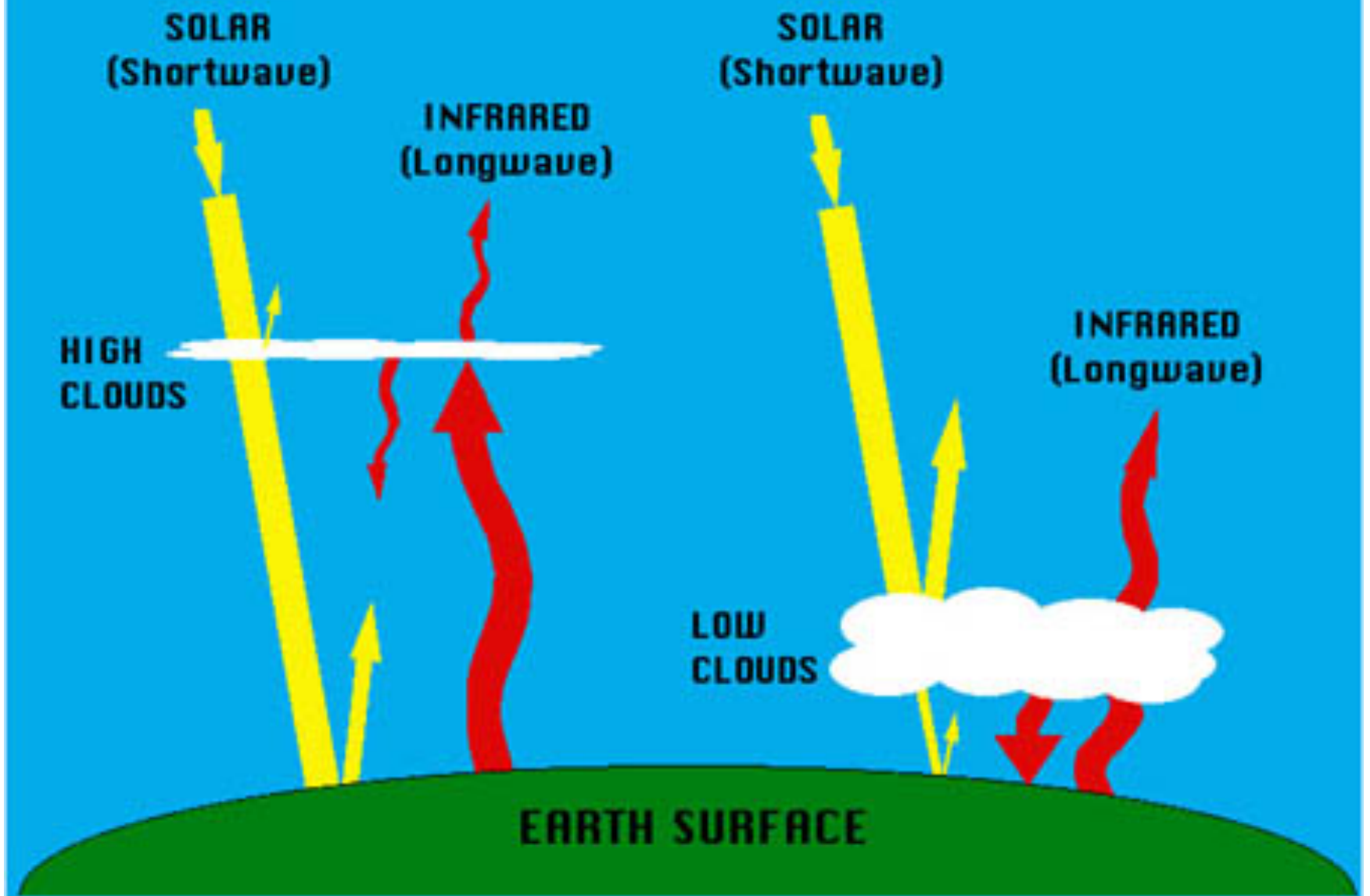
Terra and Aqua satellite images of the east Amazon basin, 11 August 2002. (*From Koren et al., 2004*)

Large scale low cloud suppression





CLOUD EFFECTS ON EARTH'S RADIATION

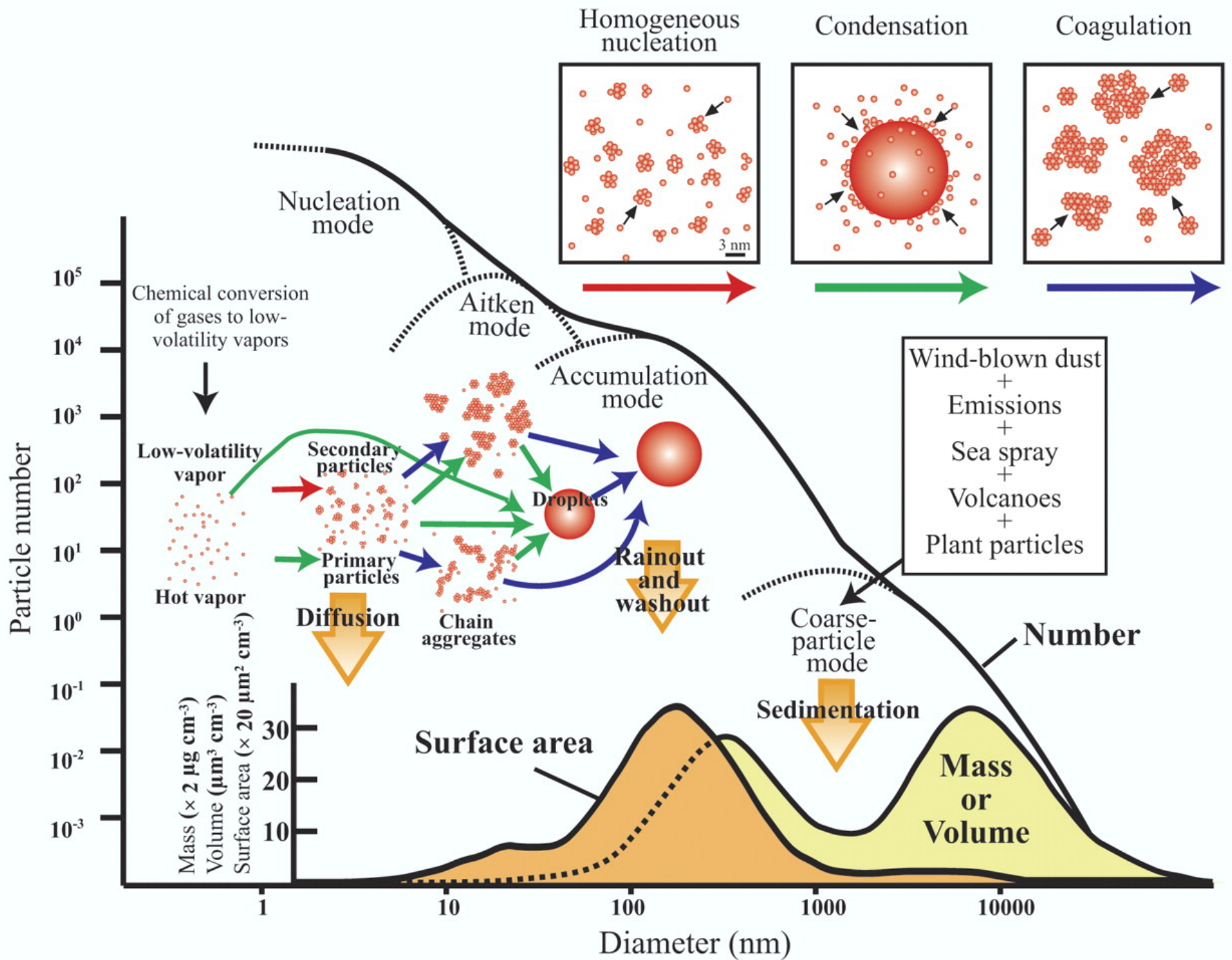


Key players of basic convective warm cloud (Cu)



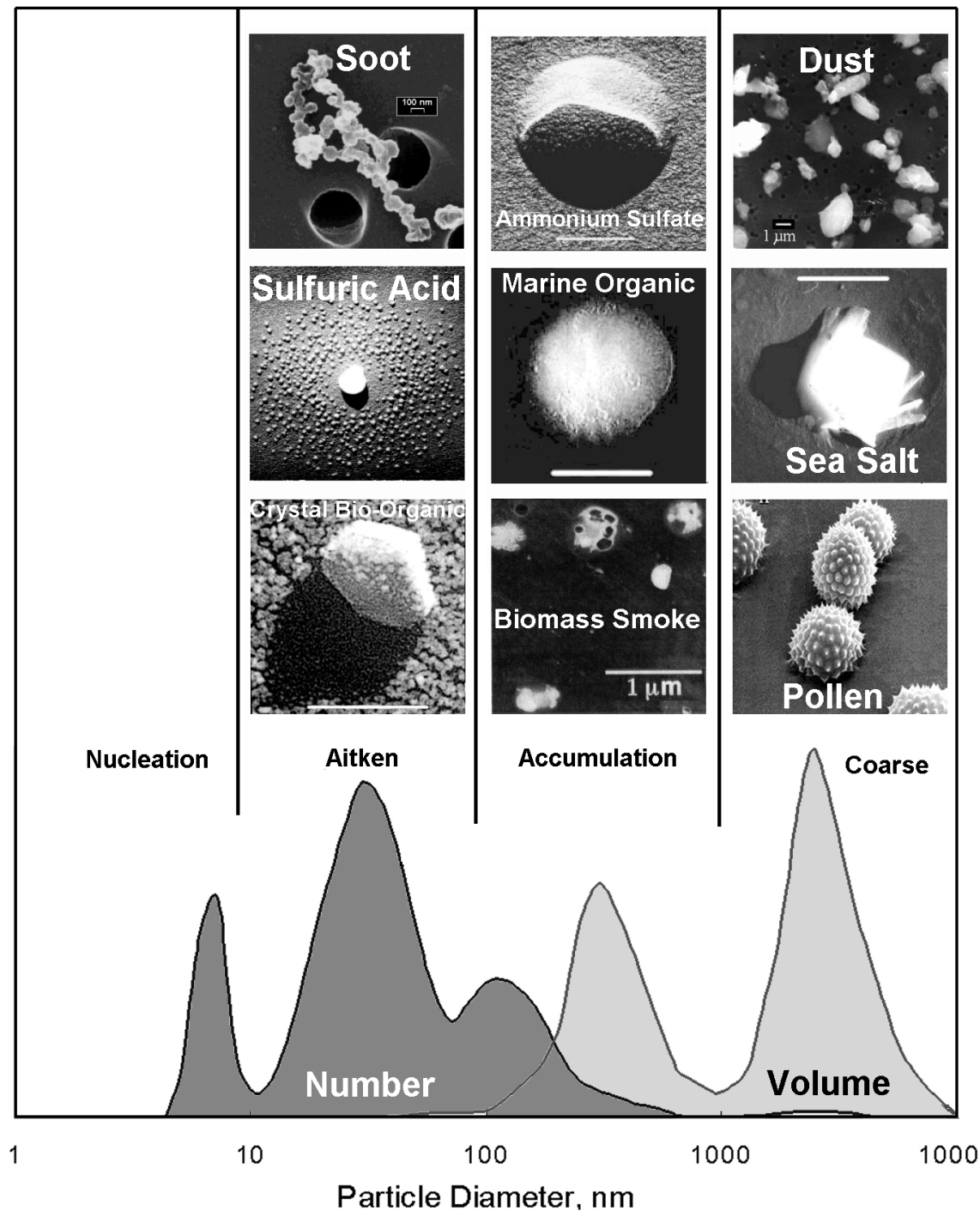
And all clouds need CCN to form





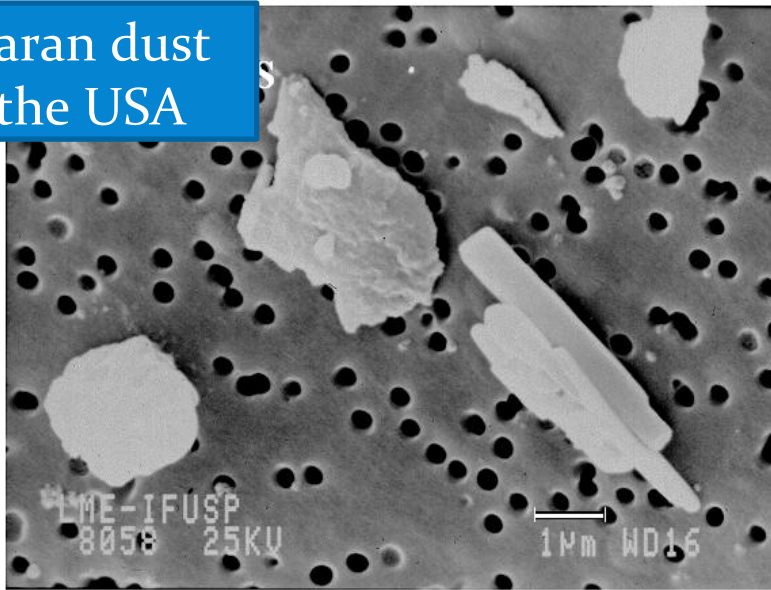
Aerosol...

- Interaction with the radiation and CCN property depends:
 - Size
 - Shape
 - Surface properties
 - Chemistry

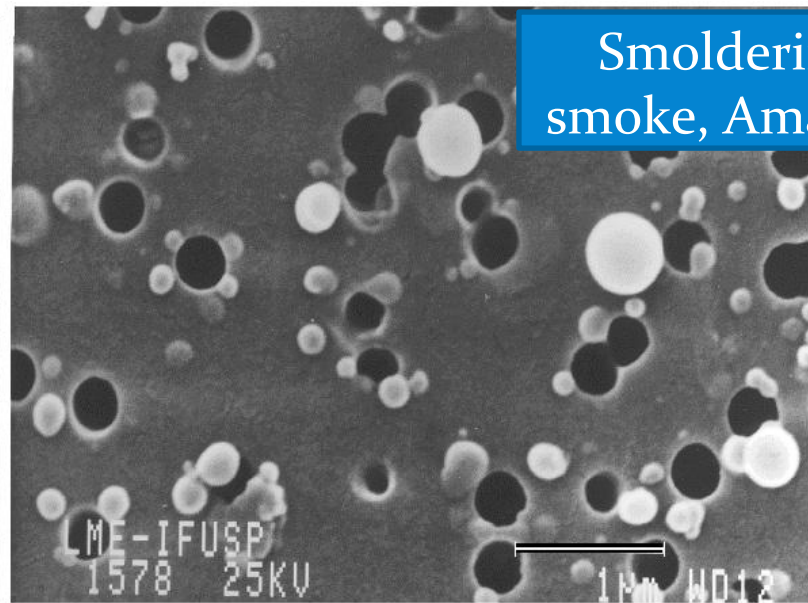


... Aerosols can be very different

Saharan dust
in the USA



Smoldering
smoke, Amazon

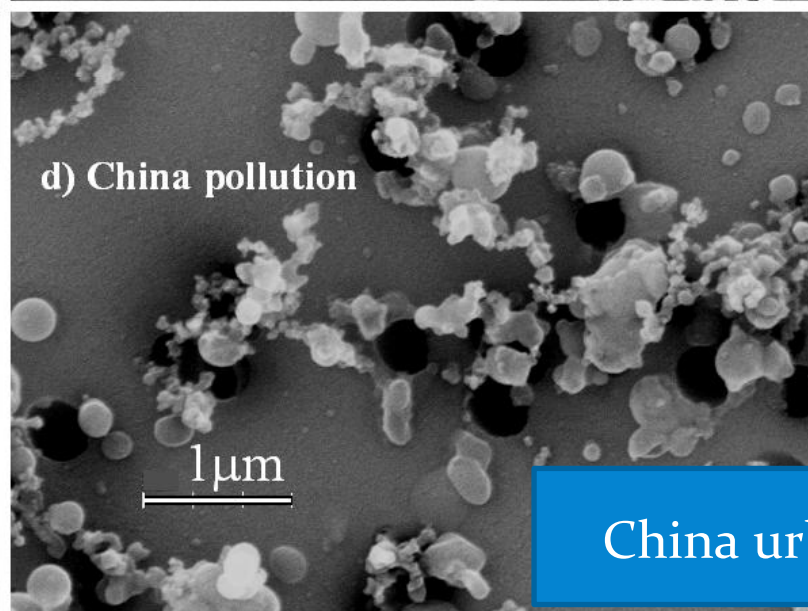


c) Smoke Cluster
from Amazon



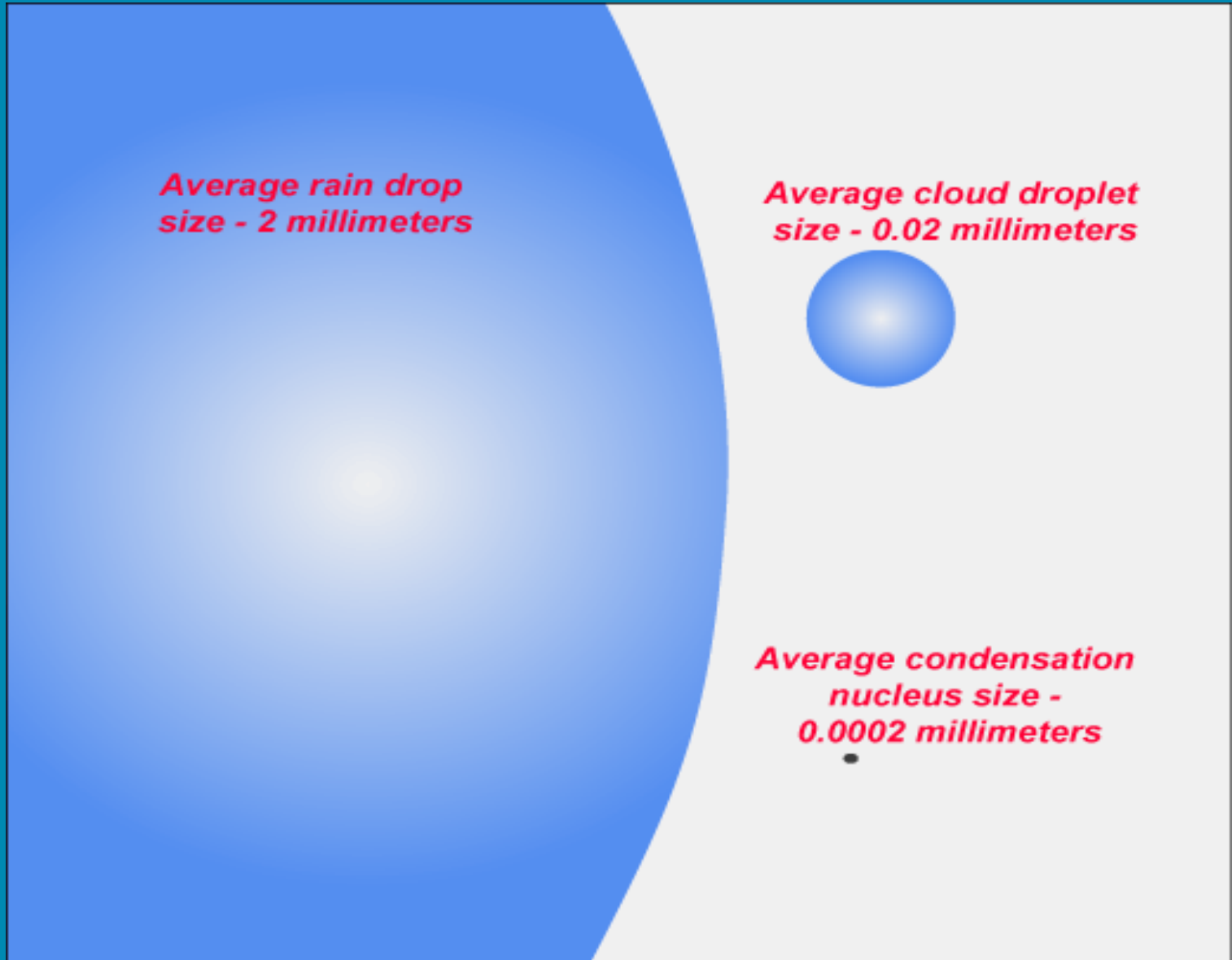
Flaming smoke,
Amazon

d) China pollution

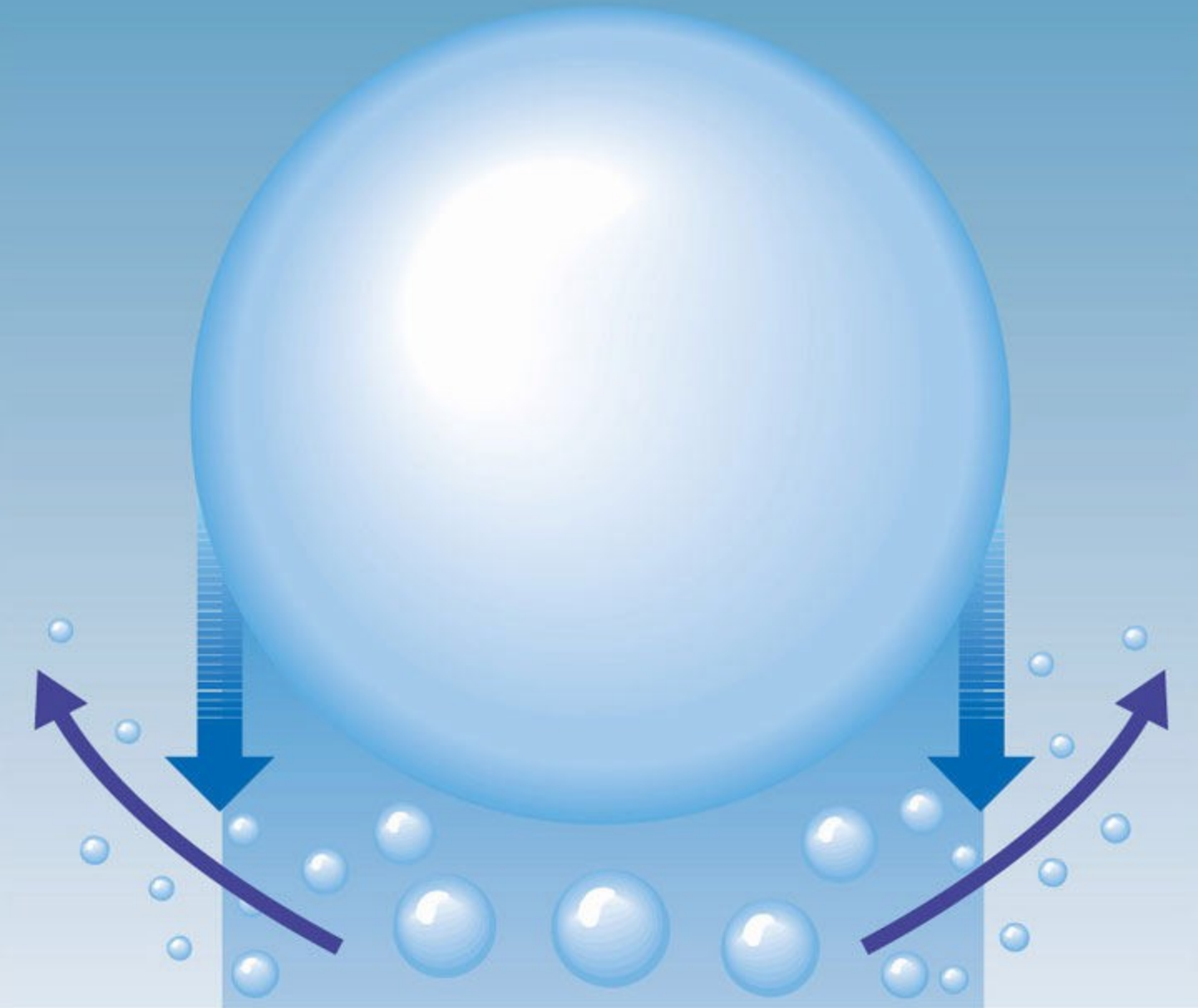


China urban

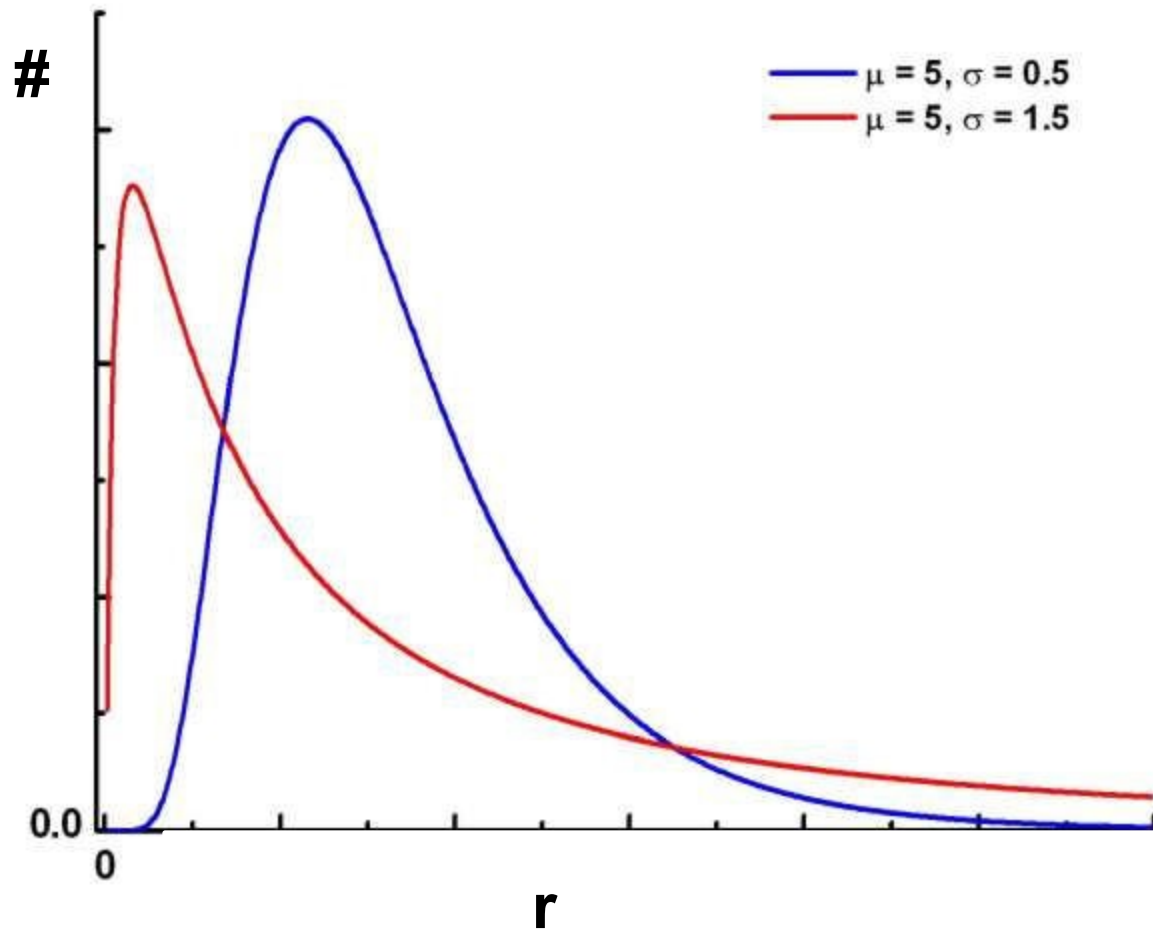
6 orders of magnitude in volume



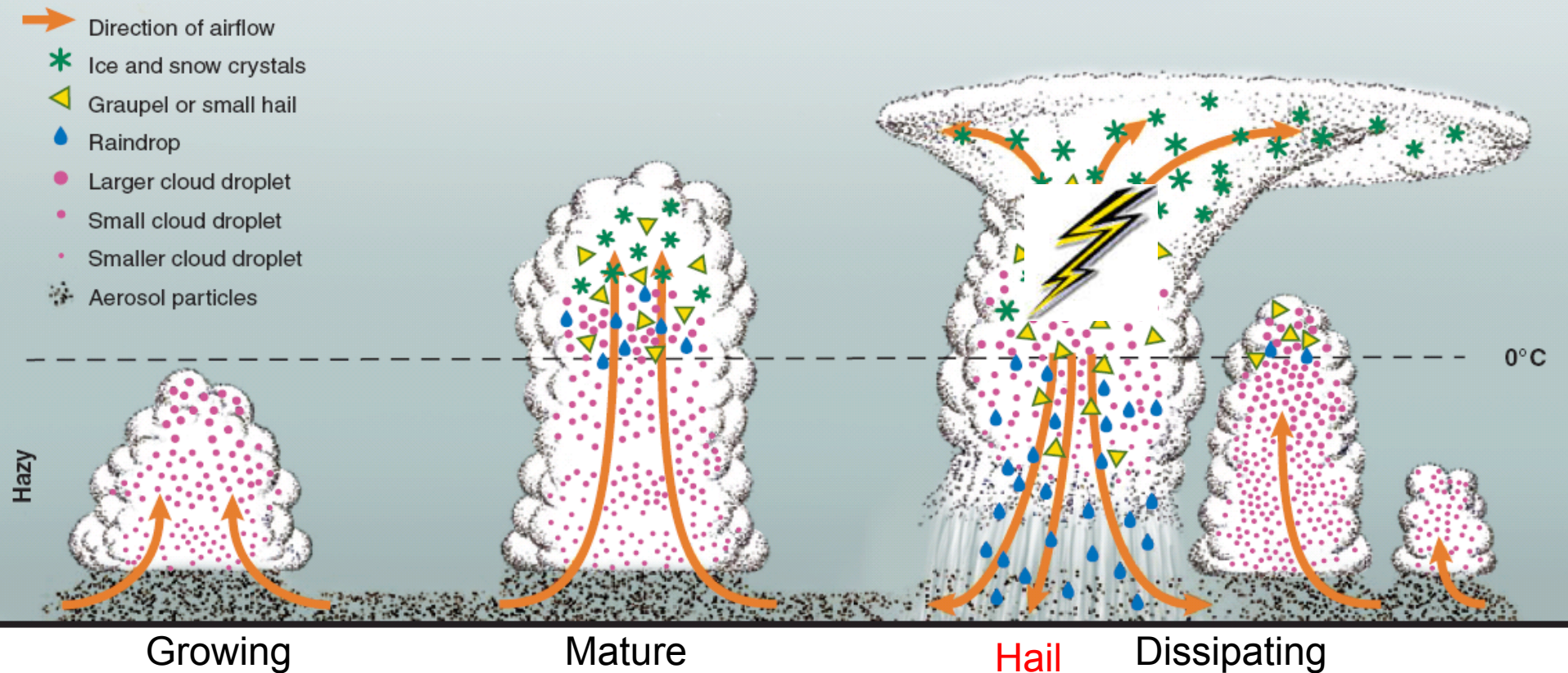
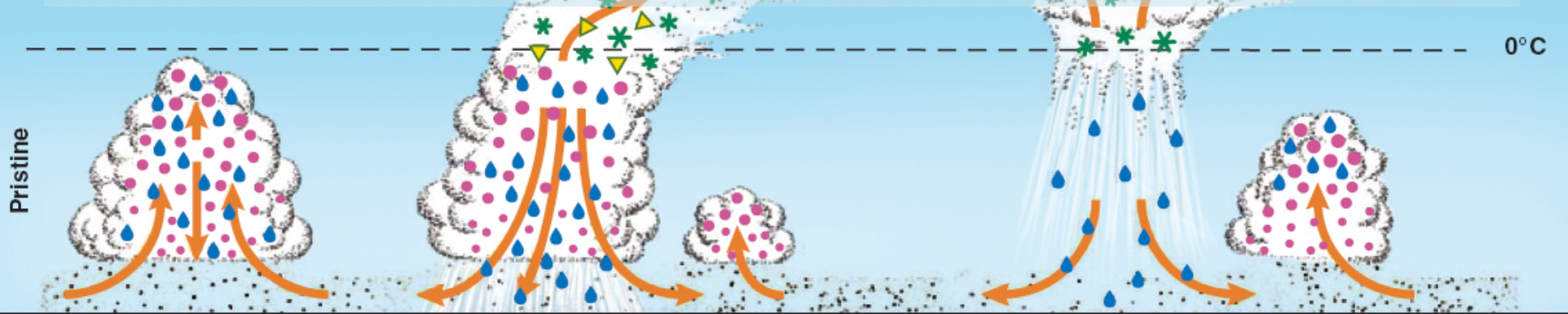
Interactions between different sizes is what makes rain



Polluted x pristine clouds



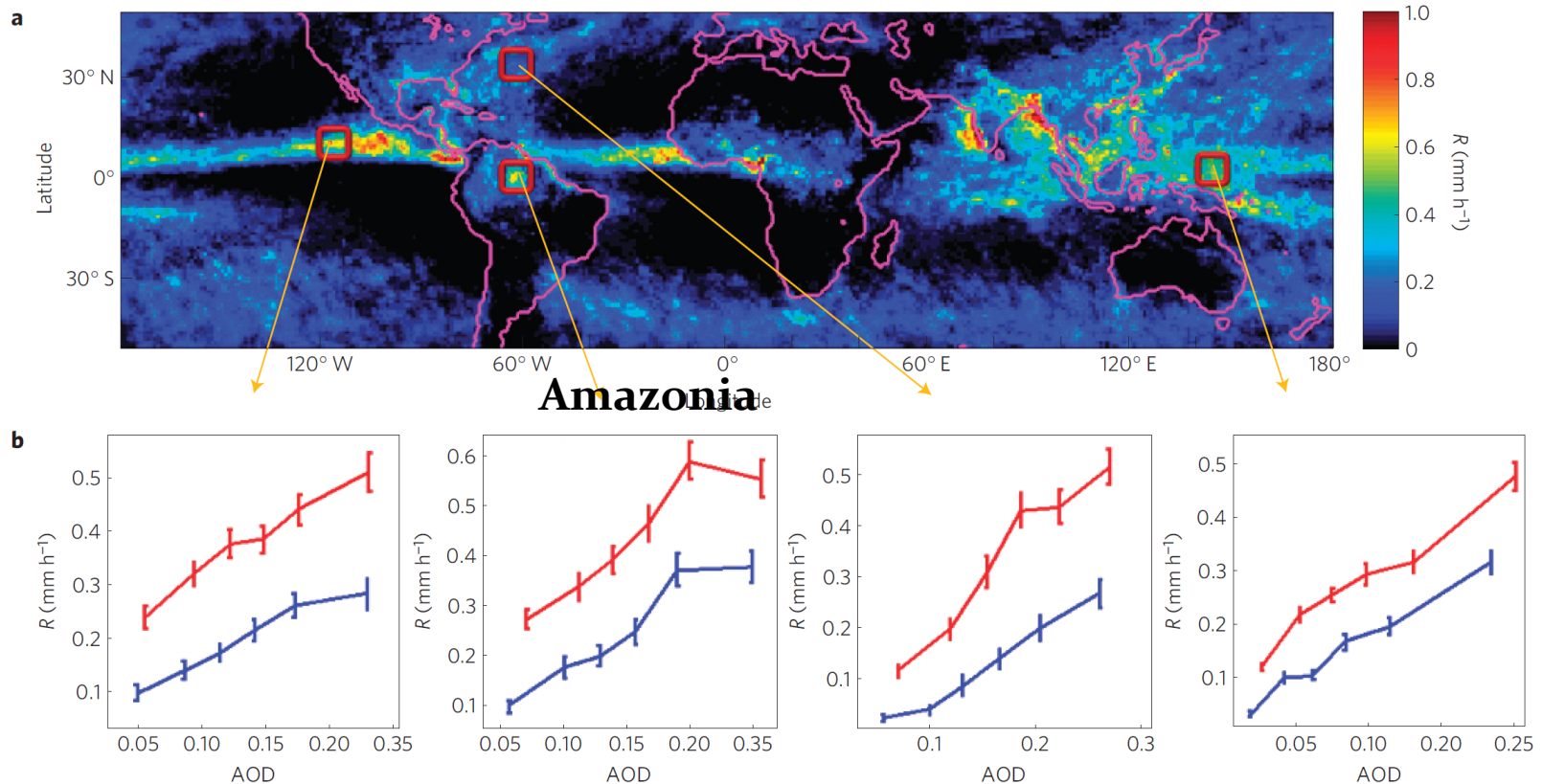
Rosenfeld D., U. Lohmann, G.B. Raga, C.D. O'Dowd, M. Kulmala, S. Fuzzi, A. Reissell, M.O. Andreae, 2008: Flood or Drought: How Do Aerosols Affect Precipitation? *Science*, 321, 1309-1313.



Rain rate (TRMM) versus Optical Depth (MODIS)

NATURE GEOSCIENCE DOI:10.1038/NCEO1364

LETTERS



13:30 local-time map of rain rate (R) and the observed trend with aerosol loading in four selected regions. Period: July and August 2007. **b**, The average R values are plotted for six aerosol-loading sets (blue, including zero R grid squares; red, without zero R grid squares). Note the R intensification as a function of AOD in all cases. (Koren et al., Nature 2012)

Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

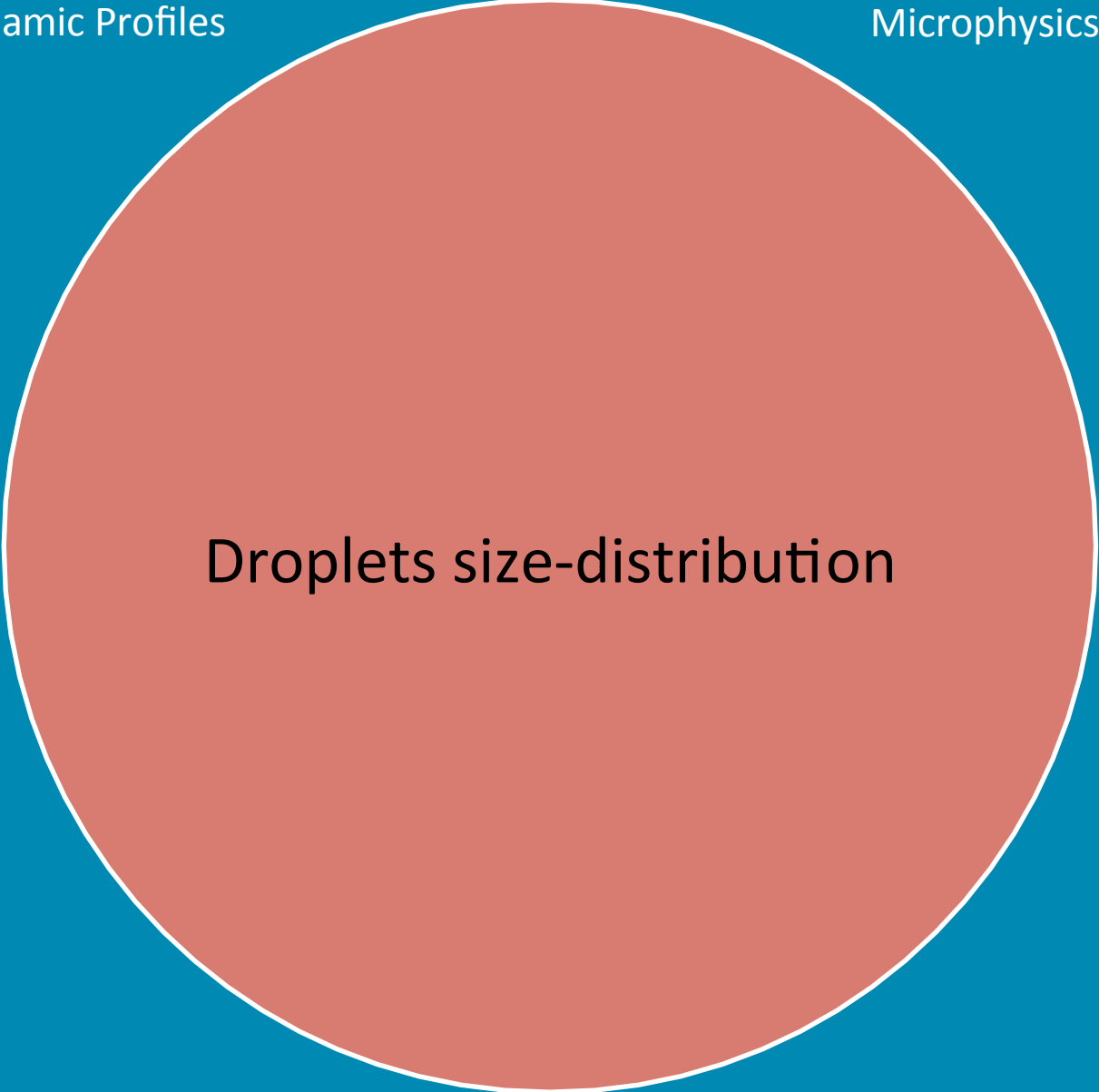
Microphysics - Aerosols



Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

Microphysics - Aerosols

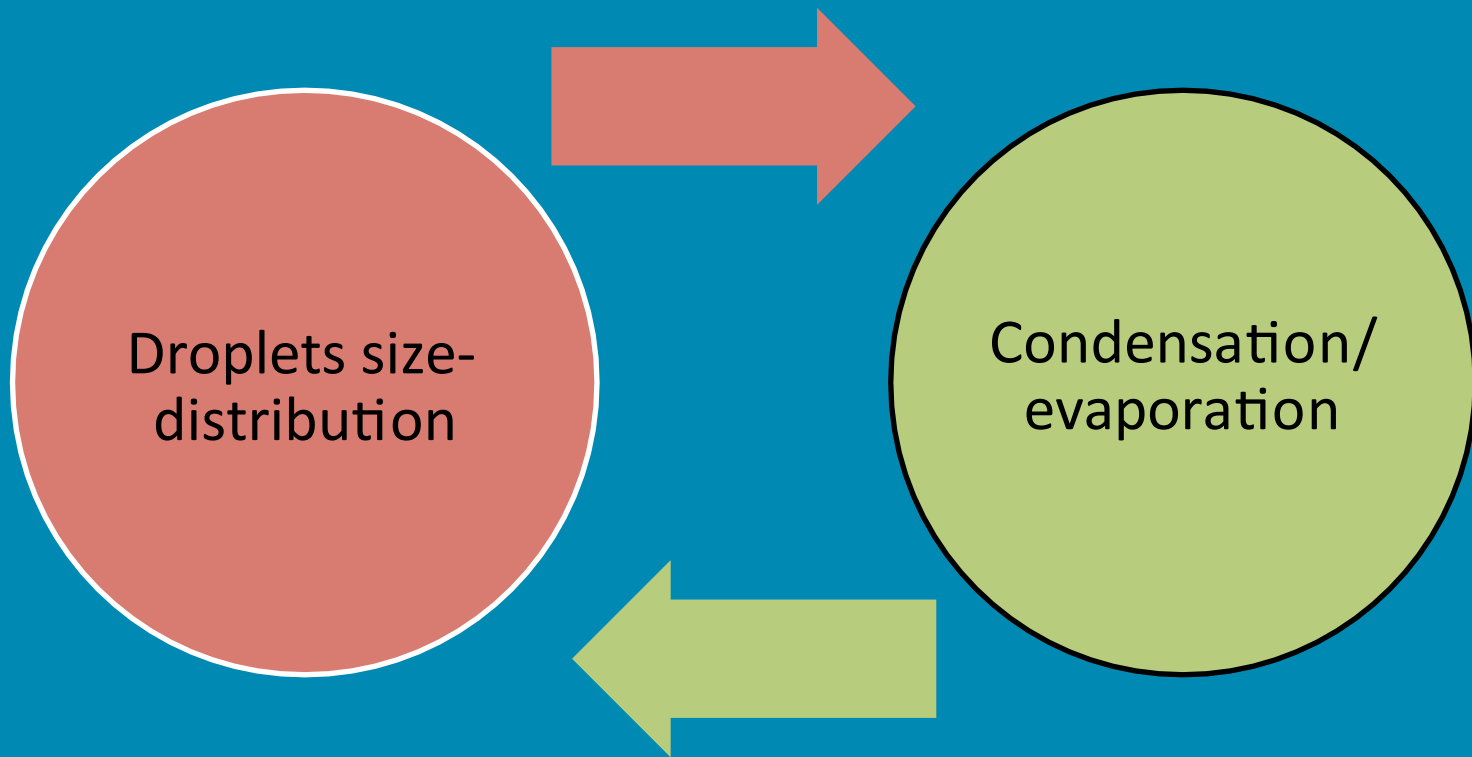


Droplets size-distribution

Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

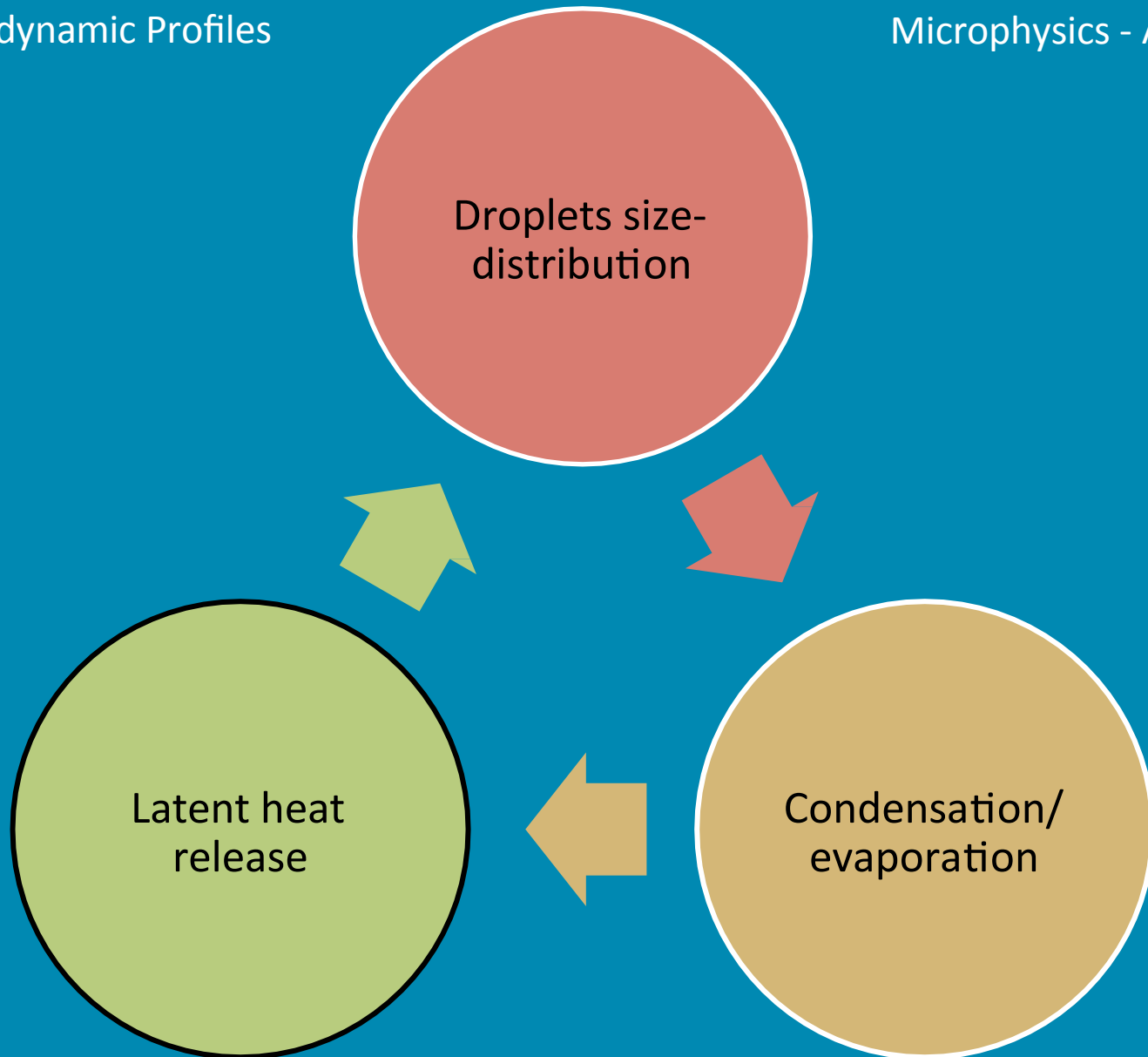
Microphysics - Aerosols



Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

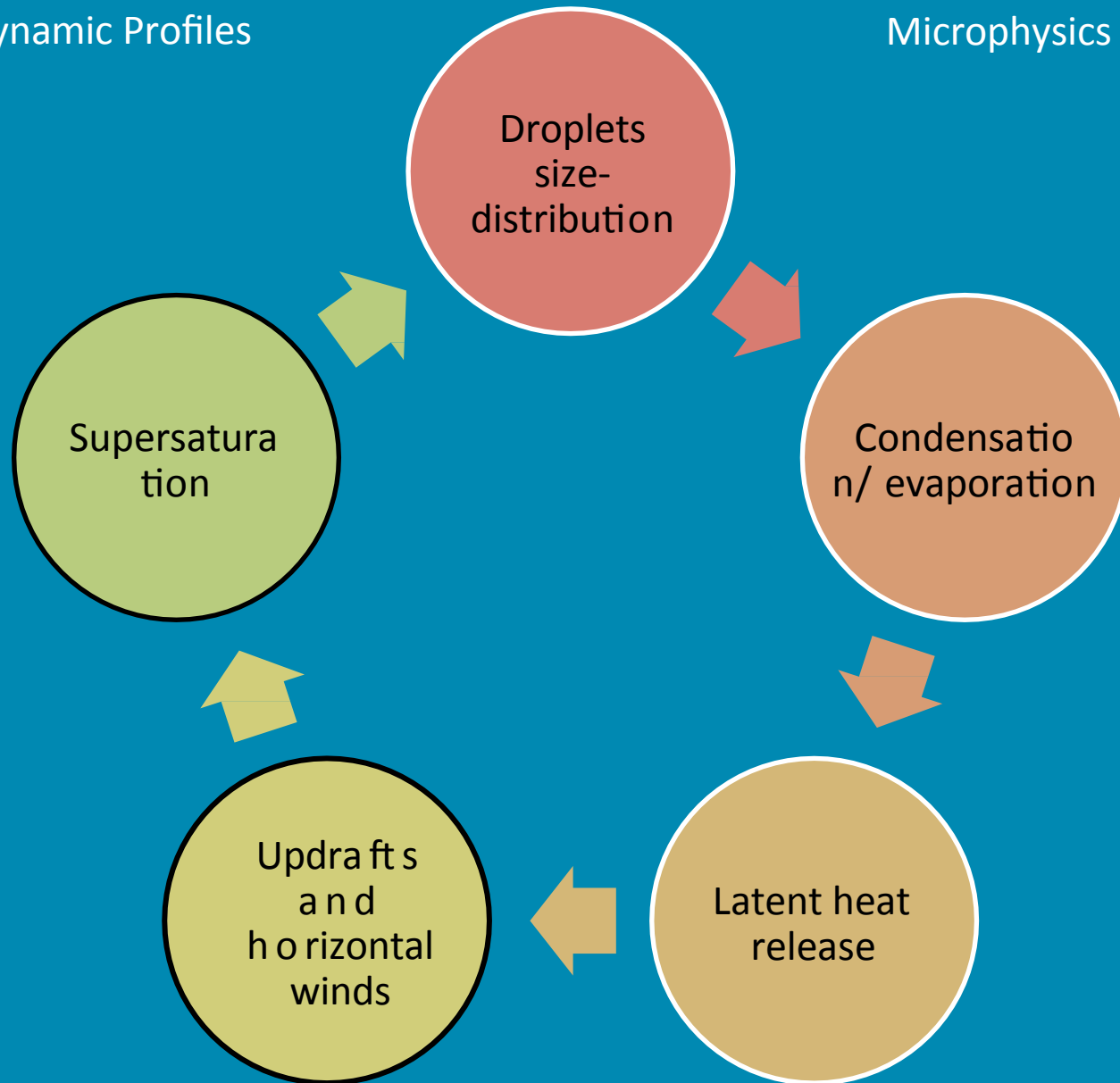
Microphysics - Aerosols



Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

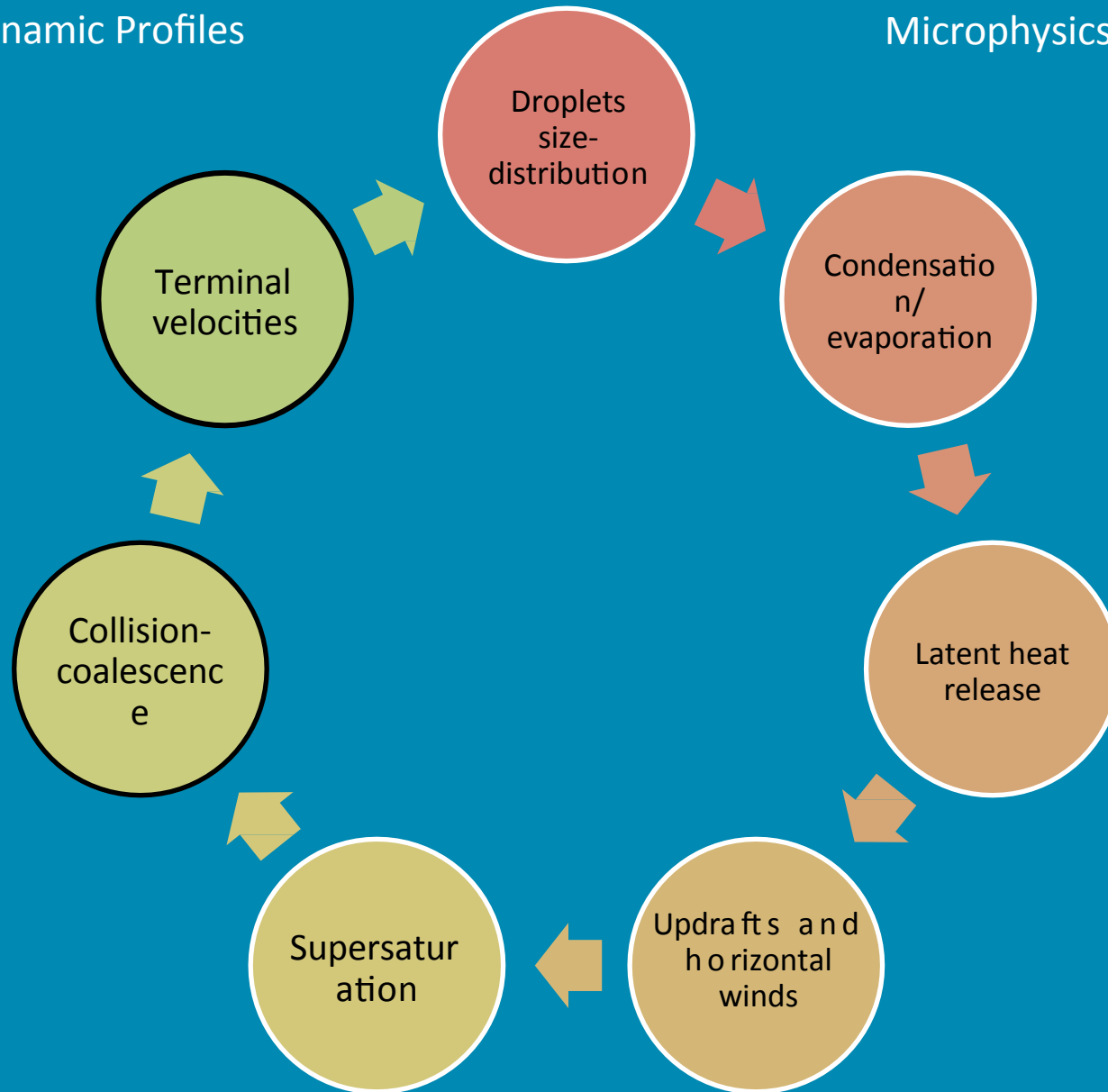
Microphysics - Aerosols



Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

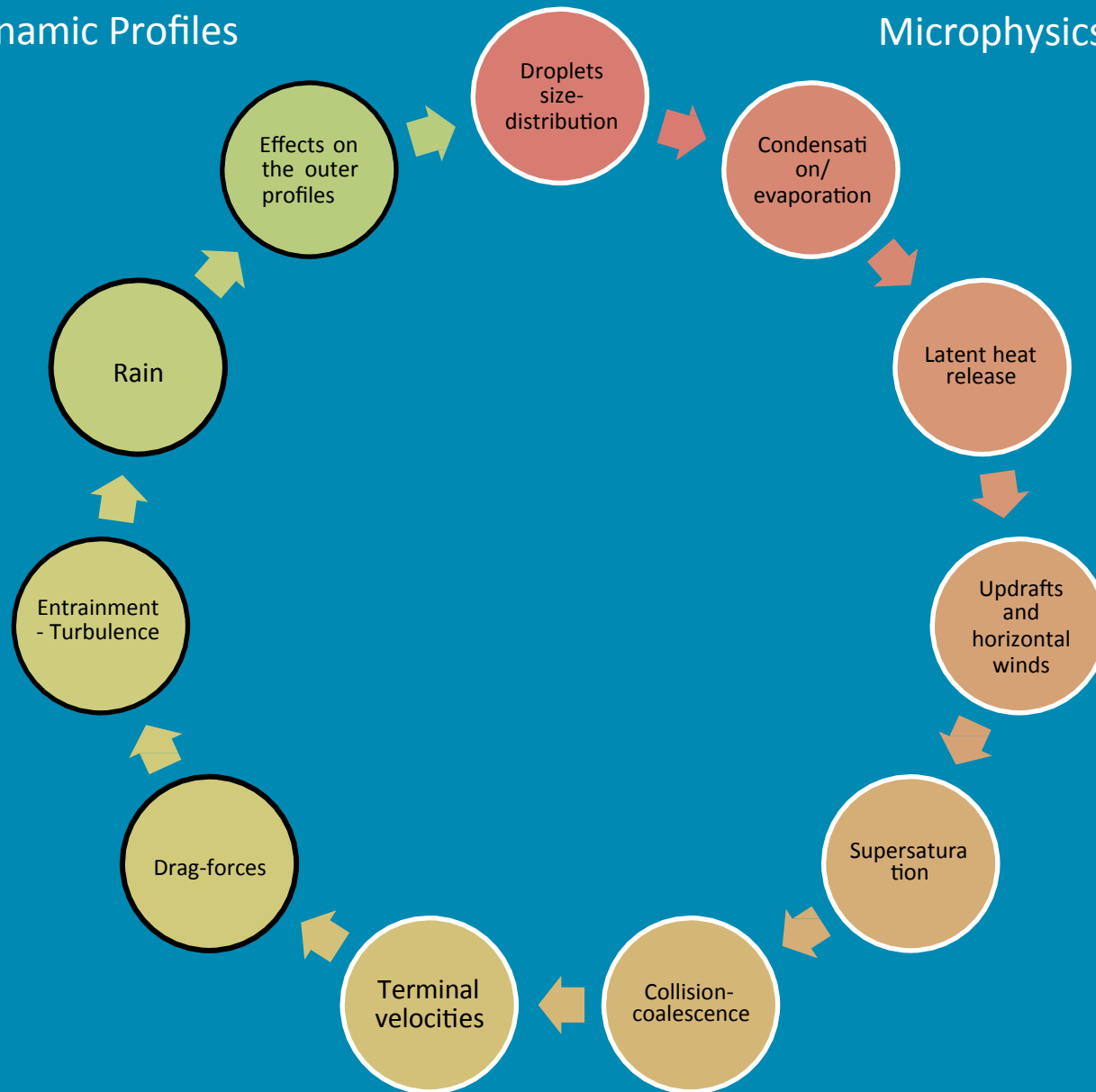
Microphysics - Aerosols



Key players of basic convective warm cloud (Cu)

Thermodynamic Profiles

Microphysics - Aerosols



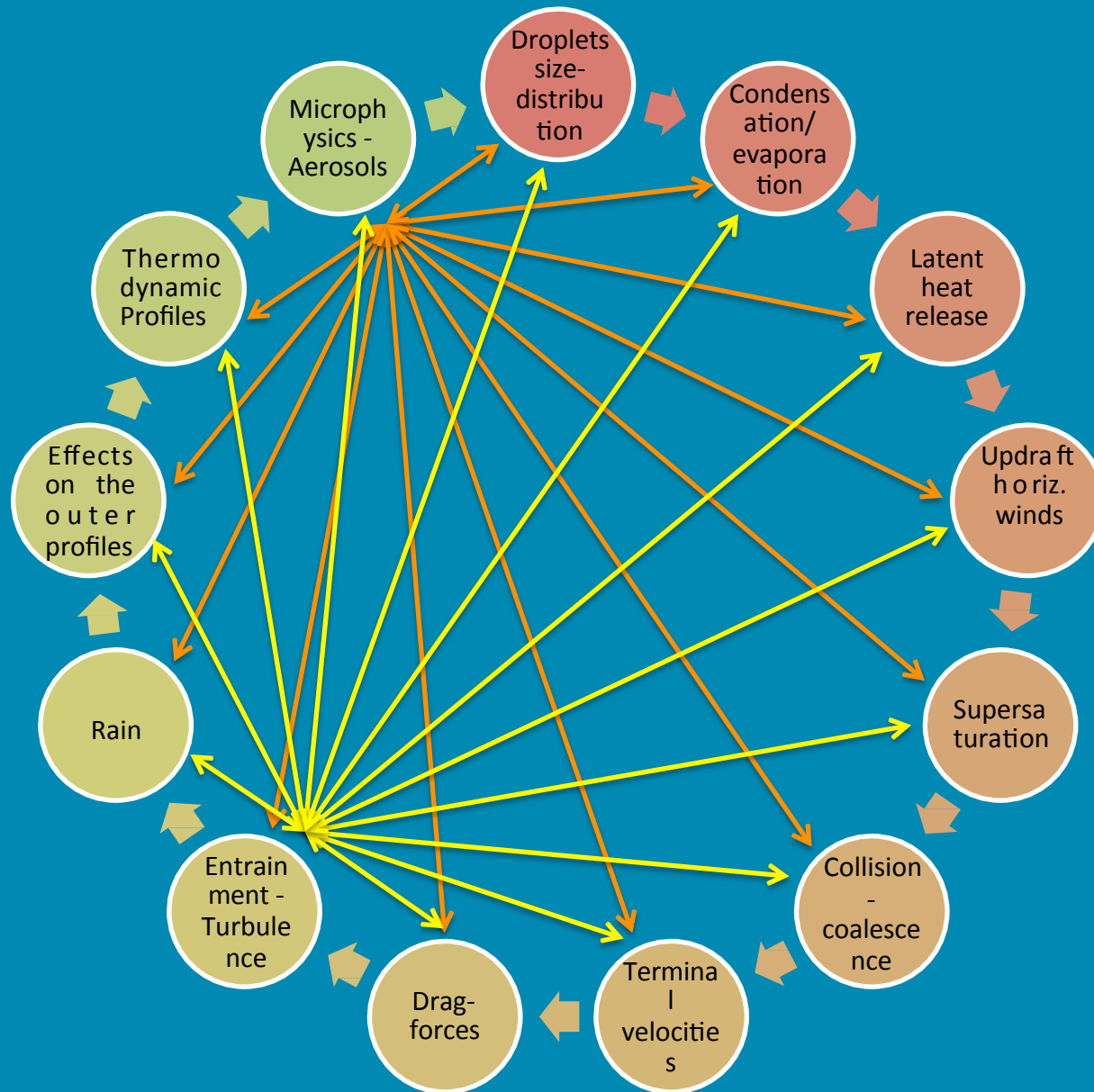
Key players of basic convective warm cloud (Cu)



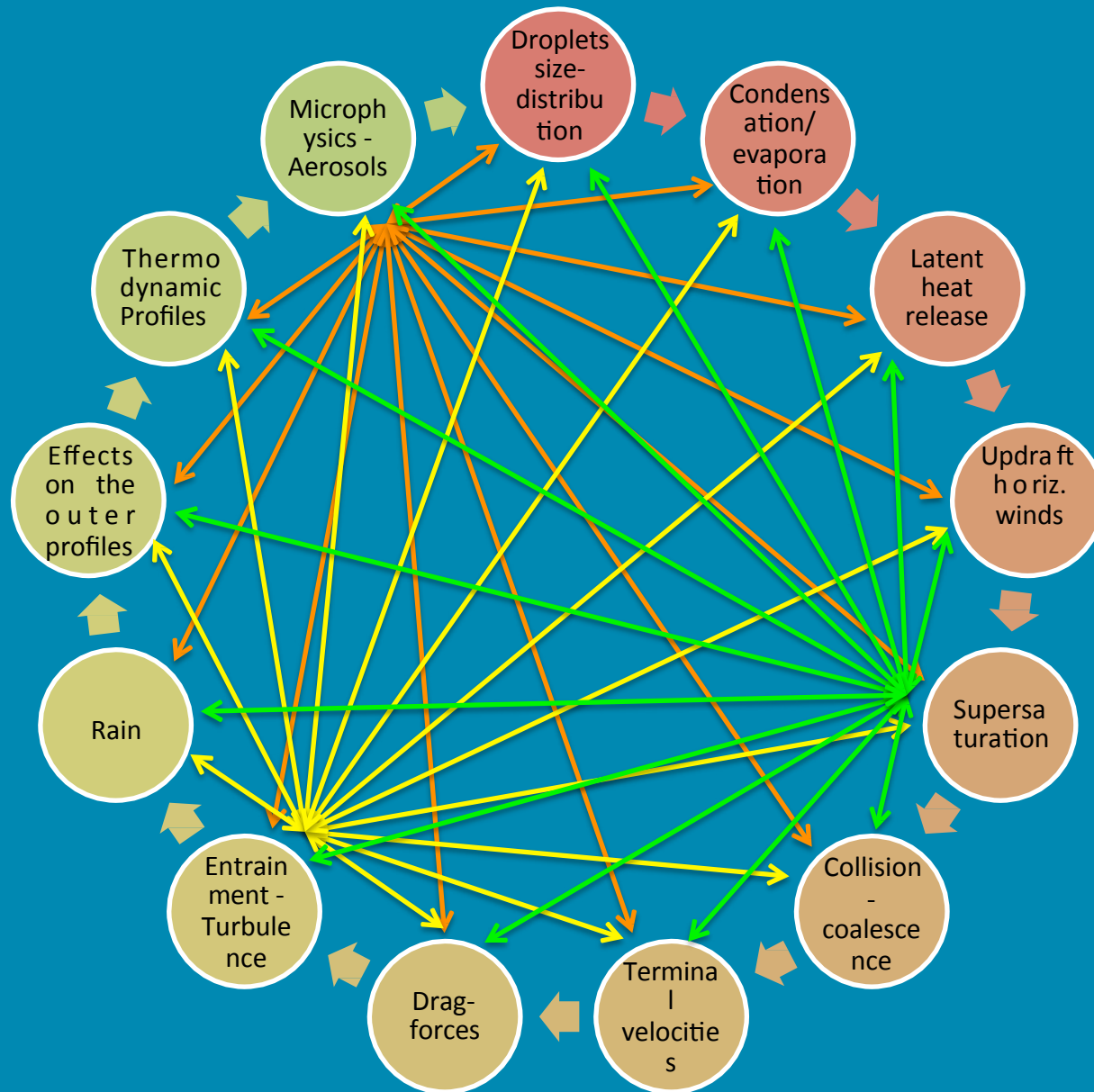
Key players of basic convective warm cloud (Cu)



Key players of basic convective warm cloud (Cu)



Key players of basic convective warm cloud (Cu)



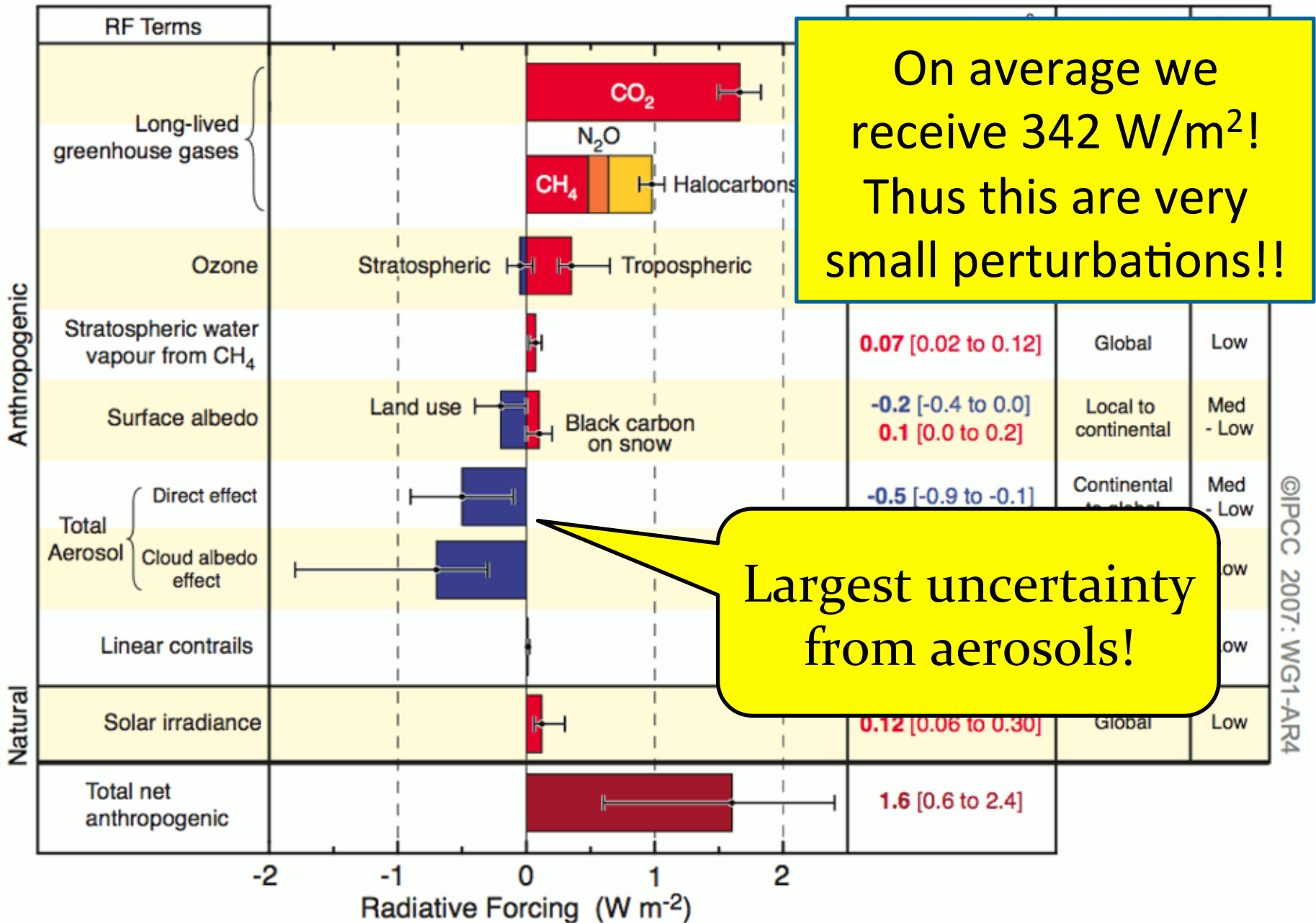
Clouds and Aerosols
after the
1st, 2nd, 3rd, ..., Nth
indirect effect



To further study the cloud (and cloud field) system

- 1) Not only microphysics – dynamics and microphysics are coupled (you can measure the coupling strength). Hence changes in aerosol properties affect cloud dynamics
- 2) Not only Marine Stratocumulus – strong aerosol signature in convective clouds
- 3) We do not understand well the warm phase and therefore how changes in aerosol properties affect it (many dynamical regimes)
- 4) Observations – models – theoretical analysis (toy models)

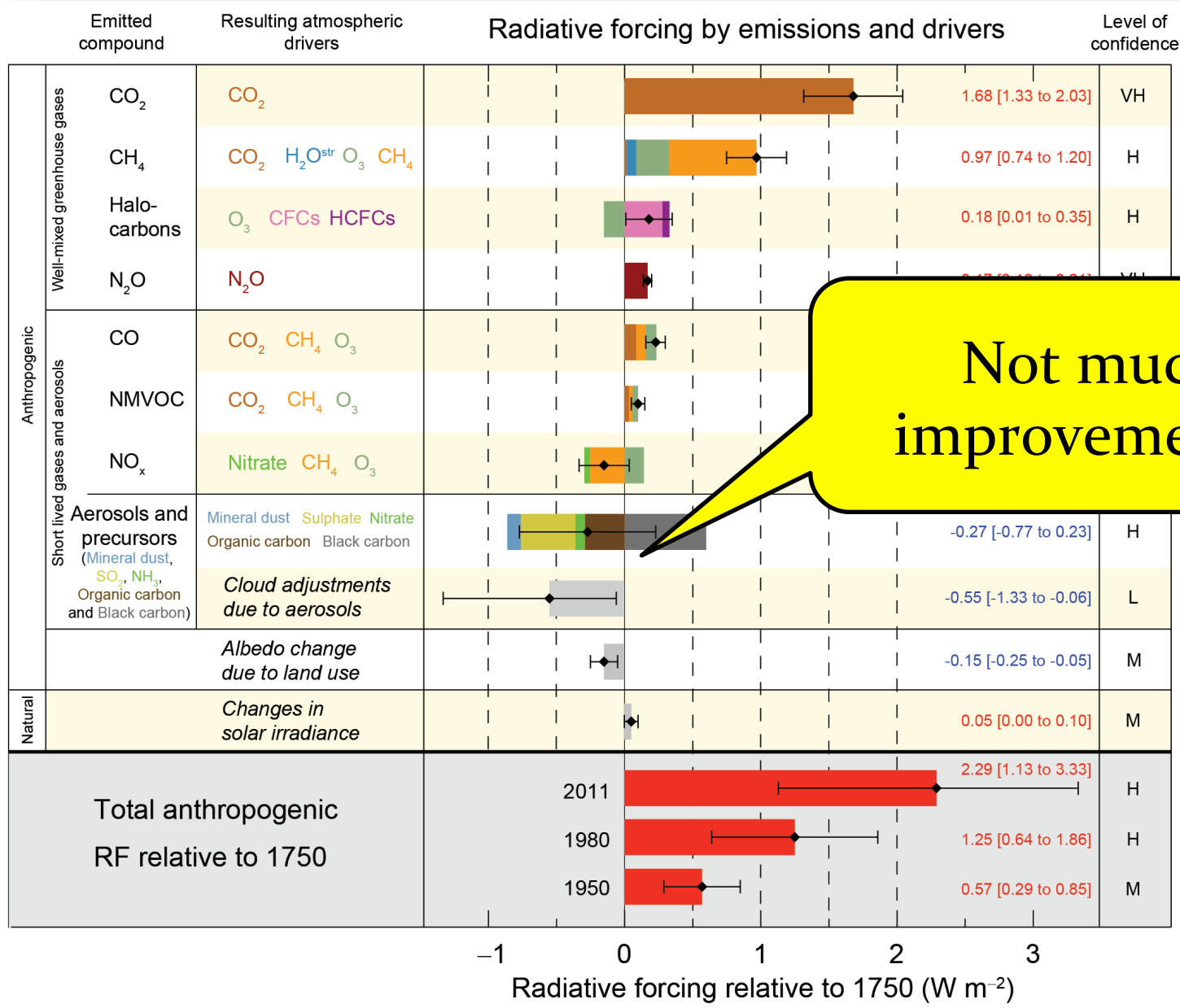
Radiative forcings of the global climate system IPCC 2007



On average we receive 342 W/m²! Thus this are very small perturbations!!

Largest uncertainty from aerosols!

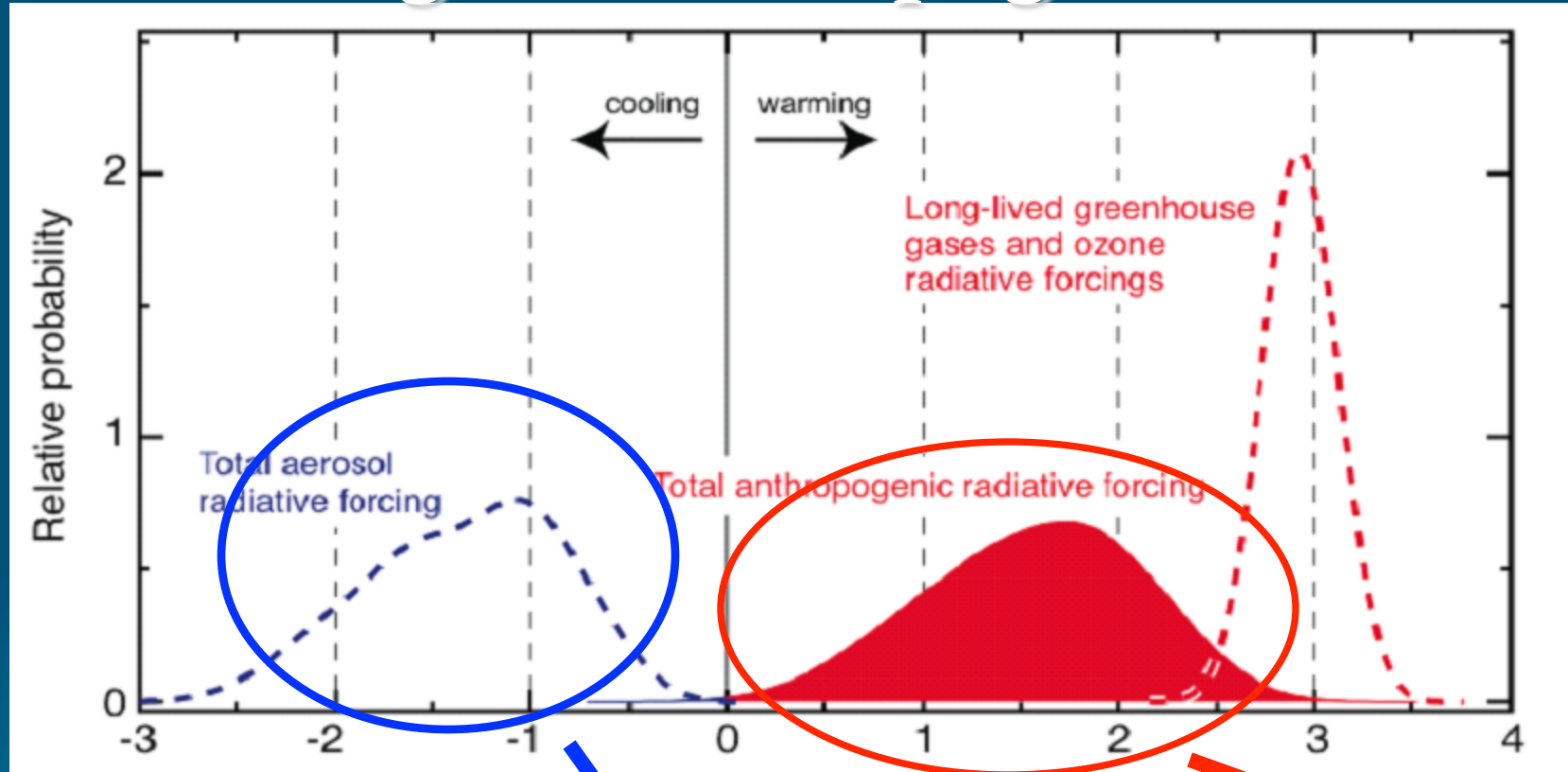
Forçante radiativa do sistema climático global (IPCC 2013)



Not much improvement...

Radiative forcing relative to 1750 (W m⁻²)

Combining all anthropogenic effects



What is being done to **this component** is critical to the **final forcing**

- **Combined anthropogenic forcing is not straight sum of individual terms.**
- **Tropospheric ozone, cloud-albedo, contrails → asymmetric range about the central estimate**
- **Uncertainties for the agents represented by normal distributions except: contrail (lognormal); discrete values → trop. ozone, direct aerosol, cloud albedo**
- **Monte Carlo calculations to derive probability density functions for the combined effect**

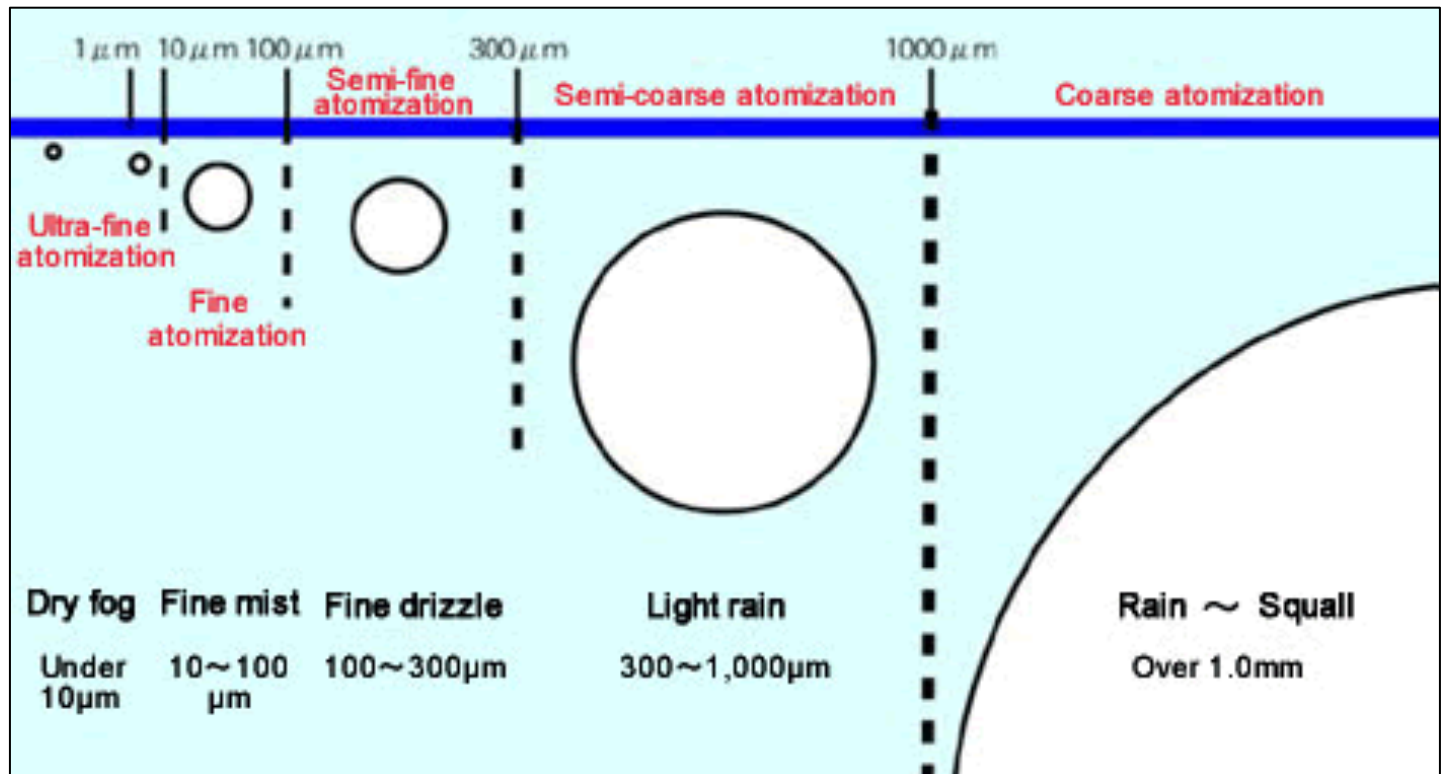
What about convection?

- 1 Sophisticated dynamics with parameterized convection (climate models)
- 2 Sophisticated dynamics with some microphysics (regional models): **bulk microphysics (N-moments)**
- 3 Simplified dynamics with sophisticated microphysics (LES/CRM): **bin microphysics**
- 4 No dynamics with explicit condensation equations (parcel models): **single particle microphysics**

4

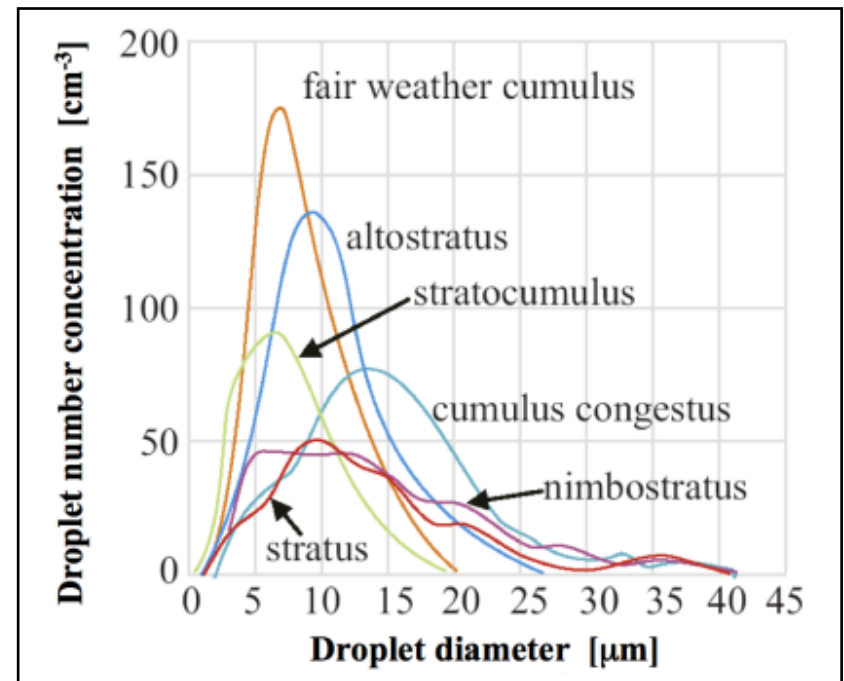
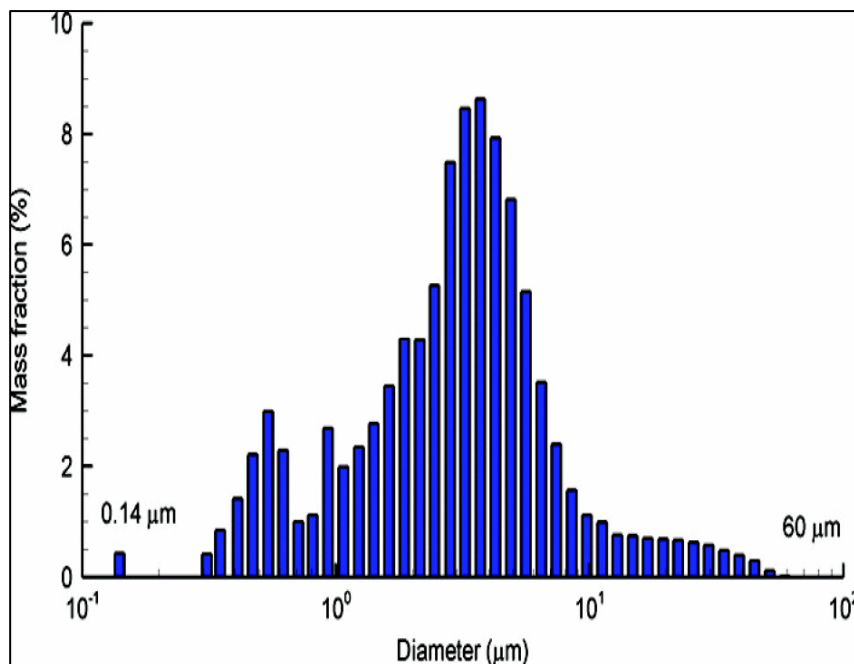
Single Particle Microphysics

- Parcel models with no dynamics, no mixing, that solve the diffusion equation explicitly for the growth of particles into droplets.



3 Bin Microphysics

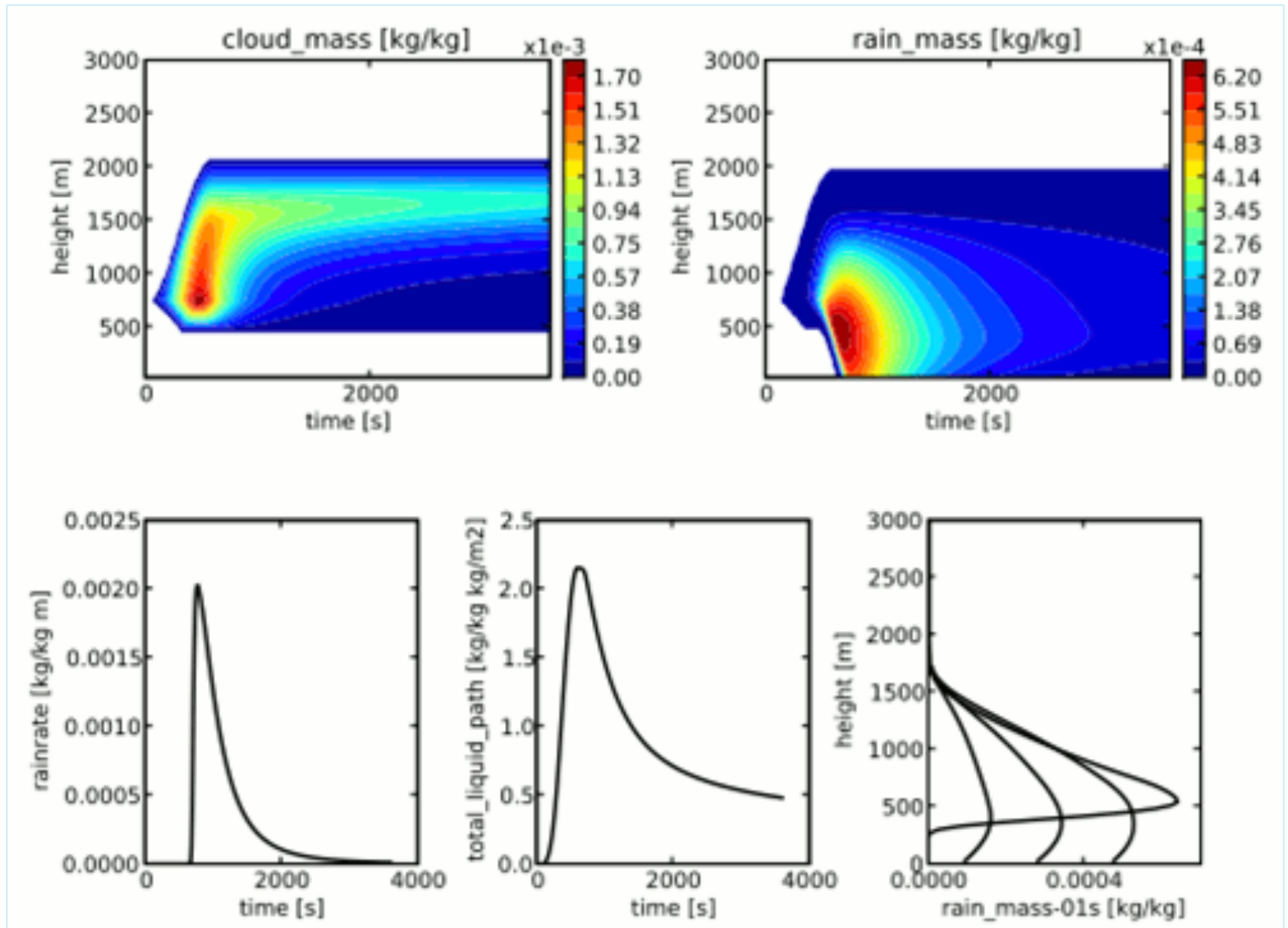
- Models that solve explicitly the size distribution of droplets as they grow, collide, and break.



- Generally used in models with simplified dynamics, typical 1D or 2D frameworks forced by large-scale tendency.

3

Ex.: Taubin – Feingold et al.



2

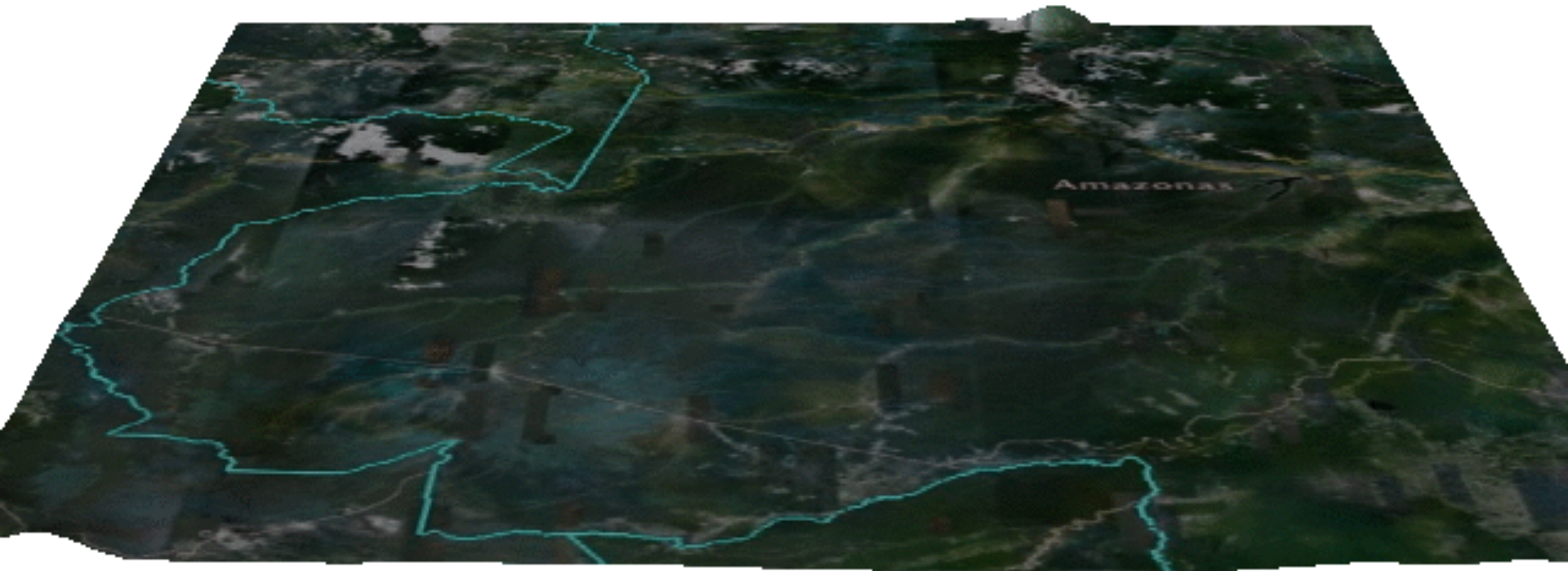
Regional Models w/ bulk Micphys

- Typically solve the budget equations for some of the moments of the size distribution. E.g., first and second moment (Kessler scheme):
 - Mass concentration (momentum 1)
 - Number concentration (momentum 0)
- Much faster than “bin” schemes, bulk parameterizations have to assume a shape for the size distribution and lose some reality.

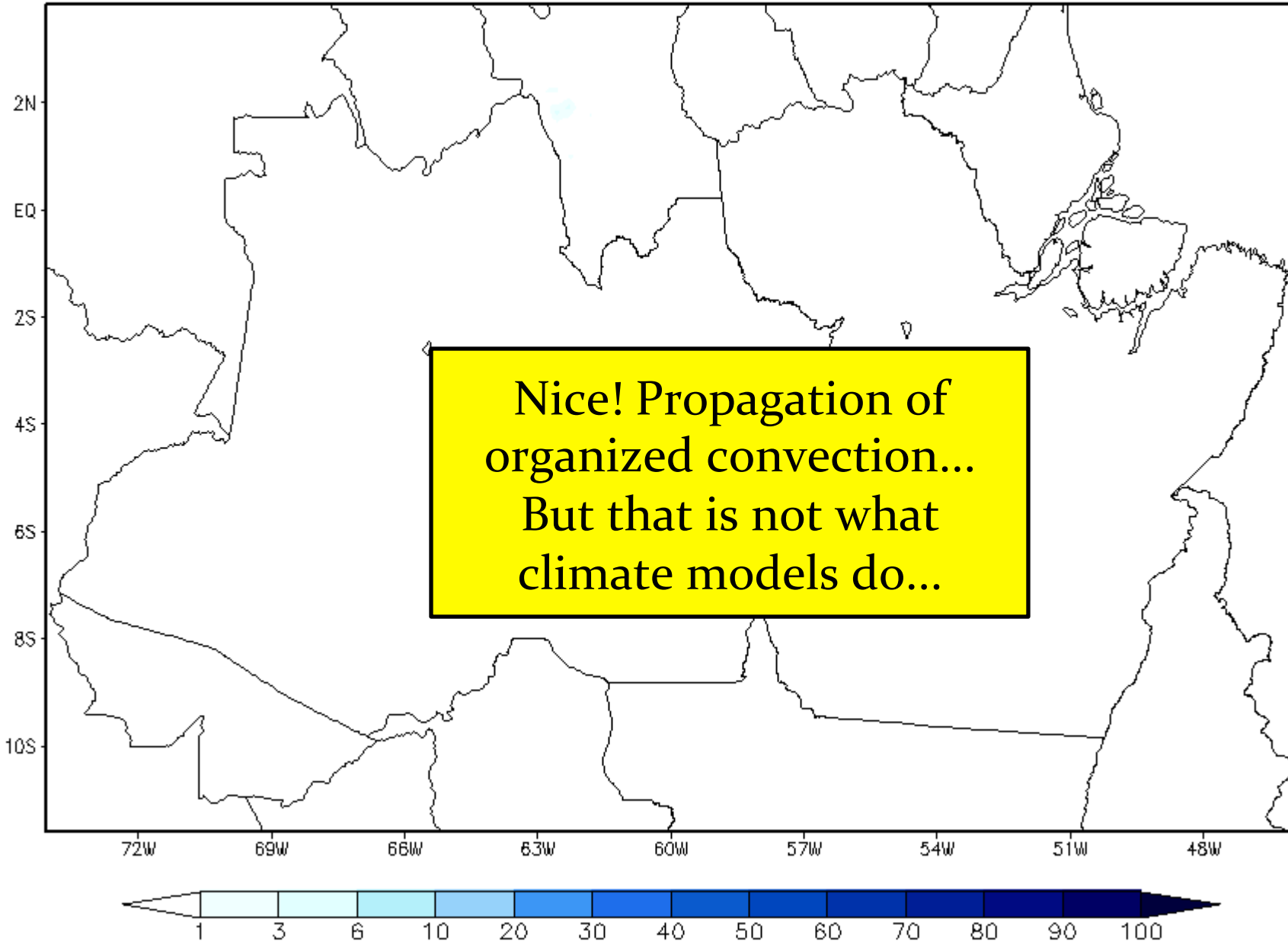
00:00:00
26 Apr 2007
1 of 16
Thursday

2

RAMS – 1 moment



Precipitacao AM PAR-OFF resolucao 3.5Km 01Z26APR2007



Parameterization:
NO organized
convection!

Precipitacao AM PAR-0N resolucao 17.5Km

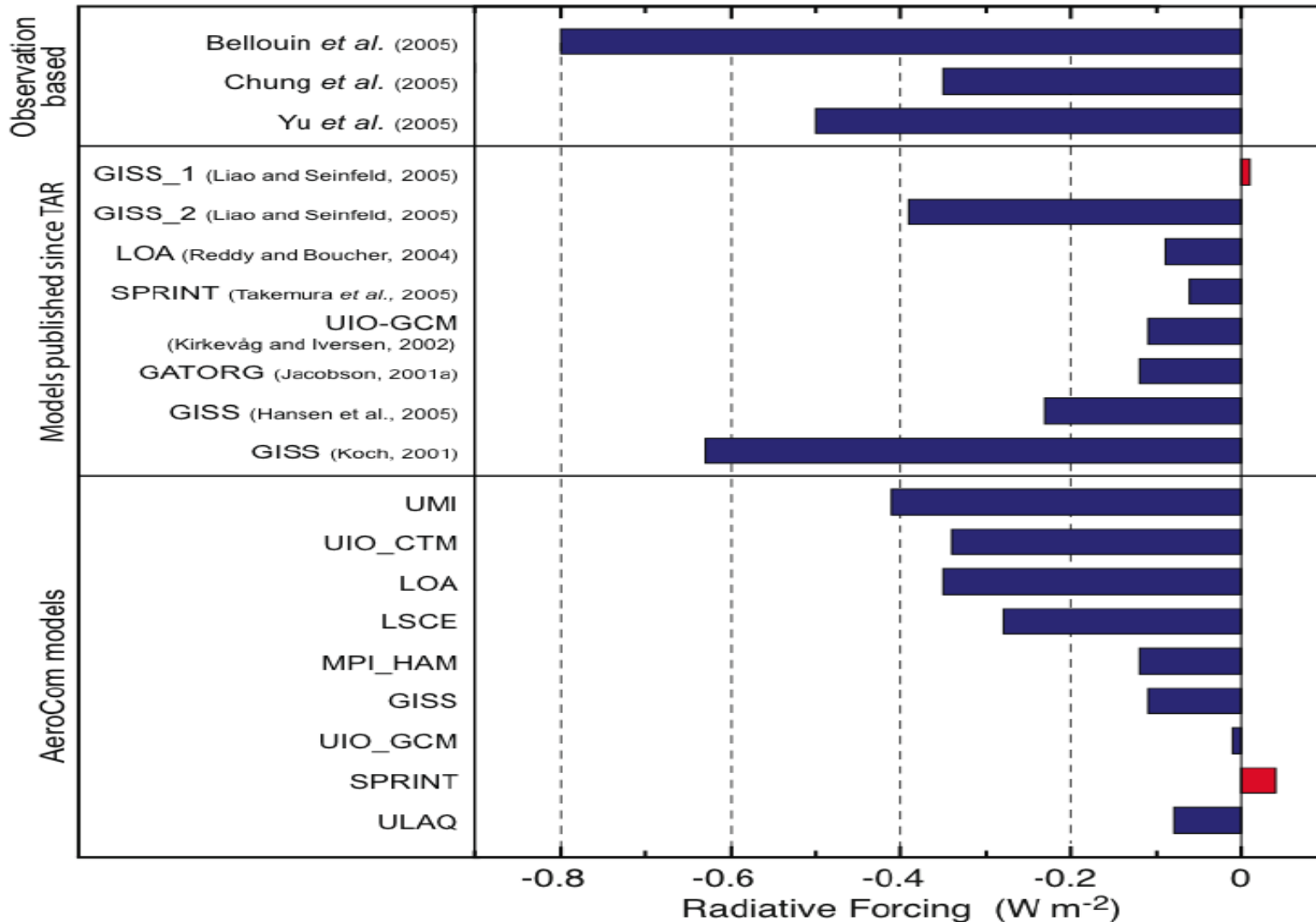


Not Nice! NO organized
convection @ 18km... NO
diurnal cycle... Rough
seasonal cycle...

But that is the best we can
do right now...

Wait, clouds are complicated but...

Disagreement between modeled DIRECT aerosol forcing!!!

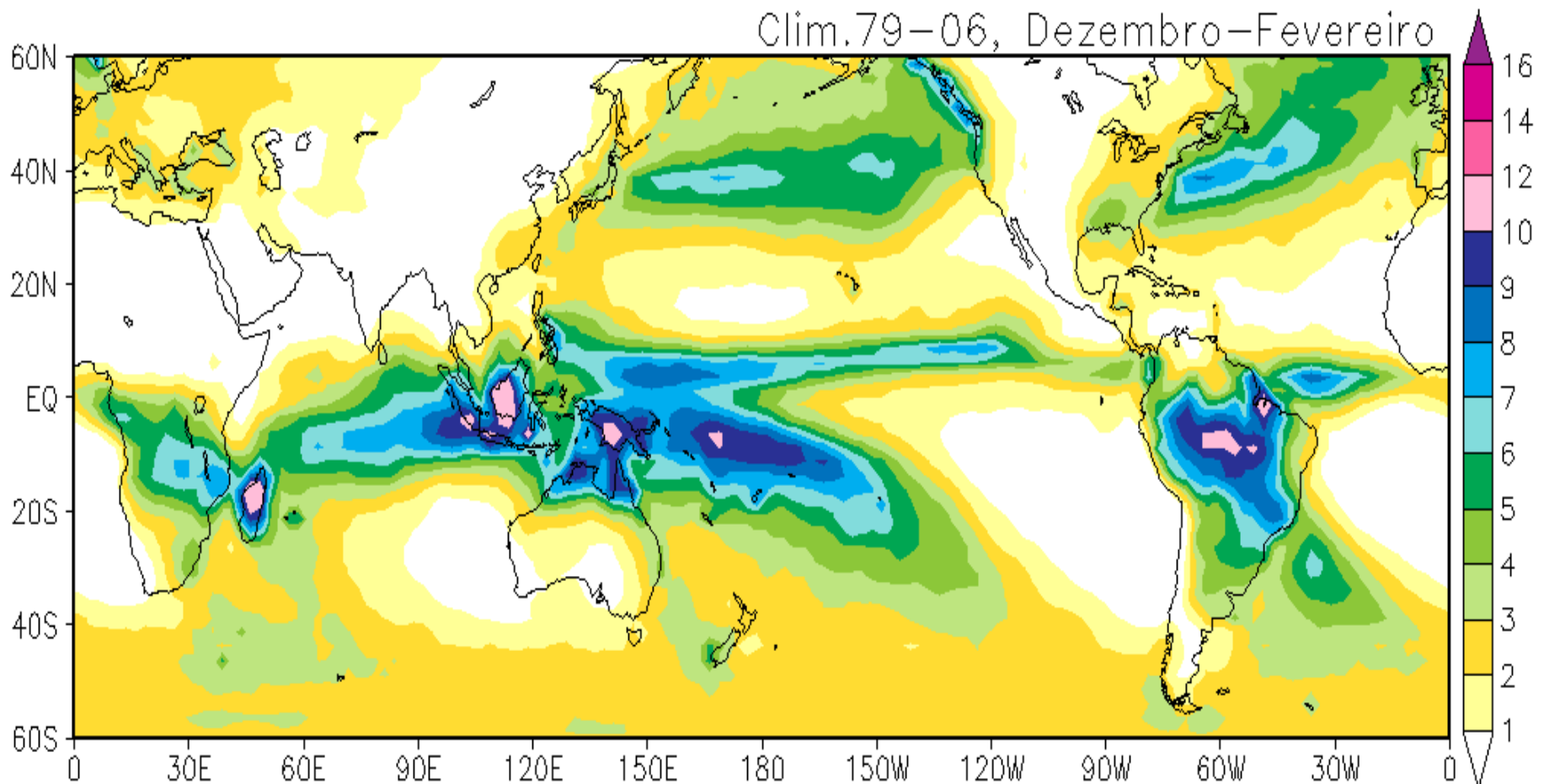




Manaus

Deforestation arch

Global Precipitation



Large Scale



Strong Diurnal
Cycle

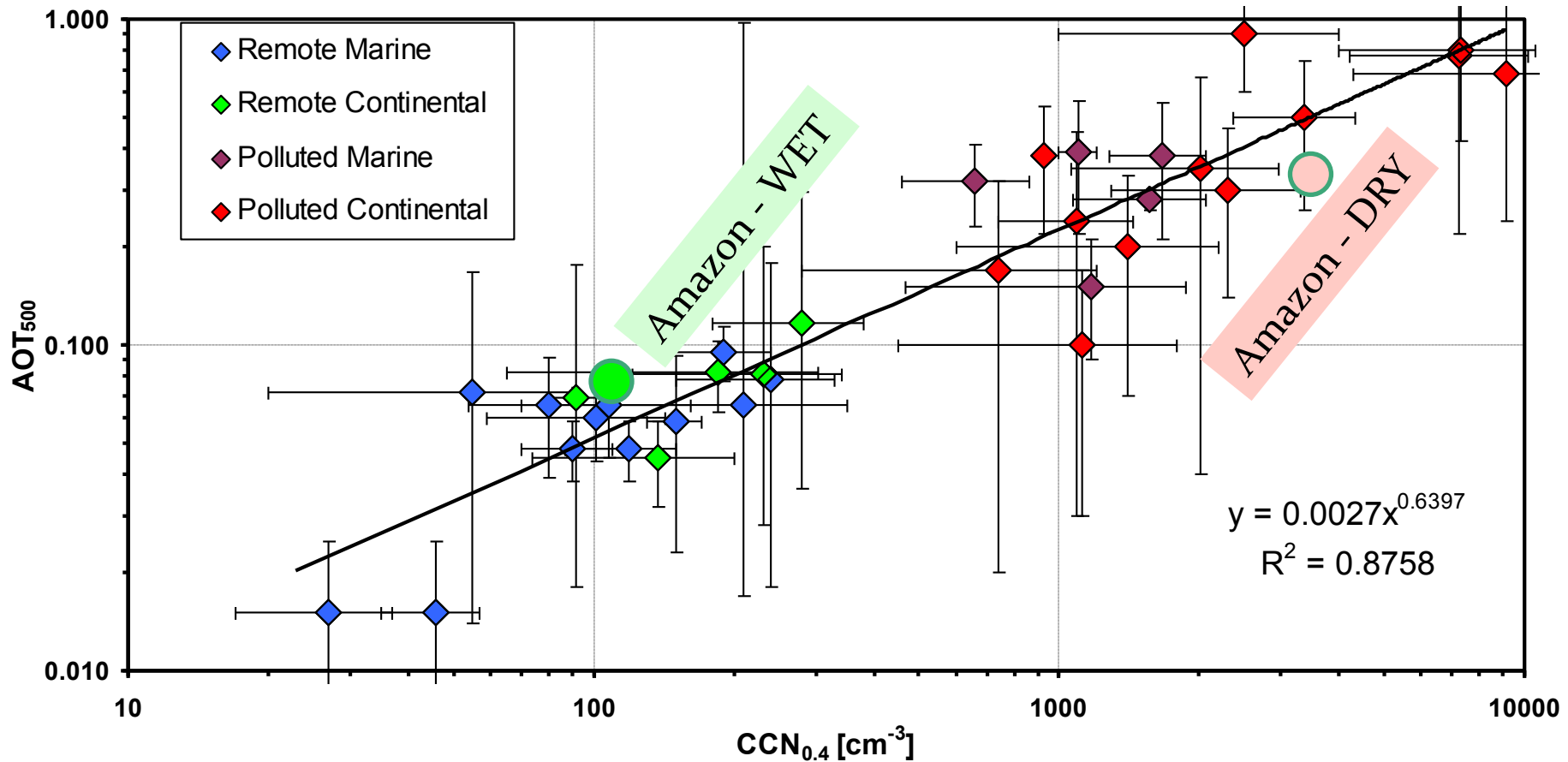
Squall Lines

ITCZ

SACZ

Fronts

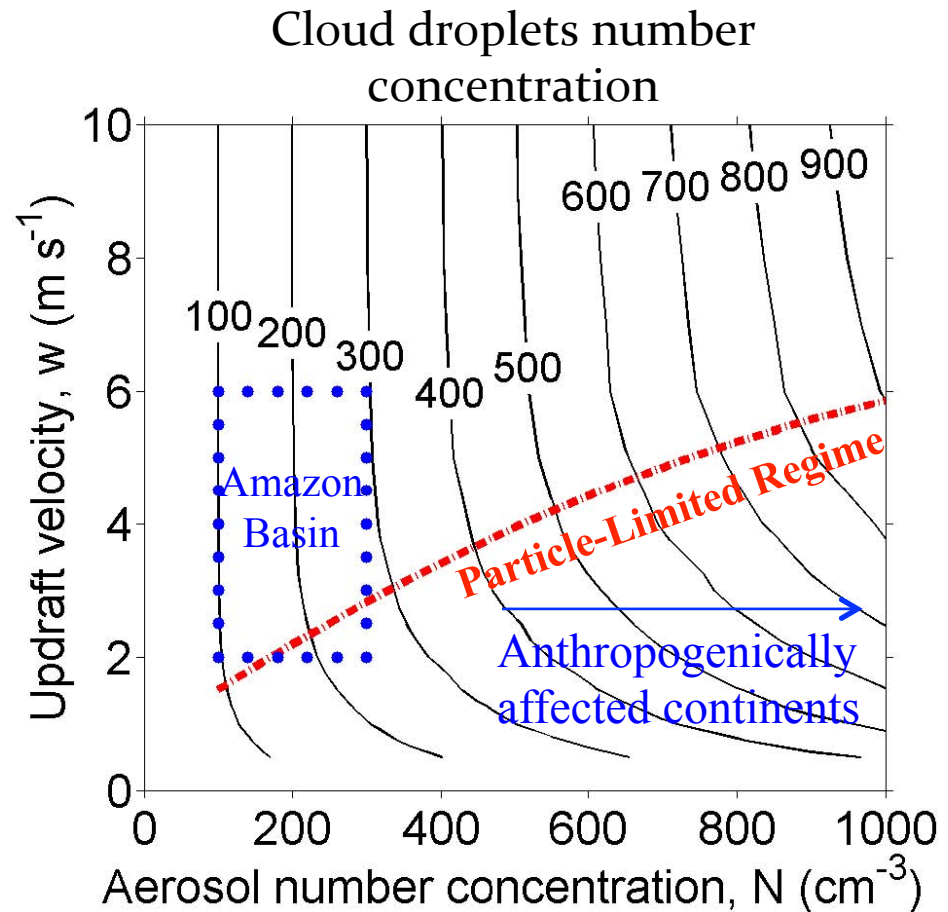
Observations of CCN and AOT



CCN concentrations and AOT over the cleanest continental sites are similar to the cleanest marine sites!

Possible Aerosol Effects

- The Amazon region is particularly susceptible to changes in CN because of the low background concentrations and high water vapor levels, indicating a regime of cloud properties that is highly sensitive to aerosol microphysics.





ATTO

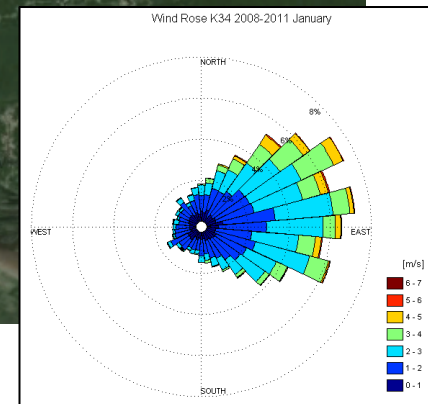
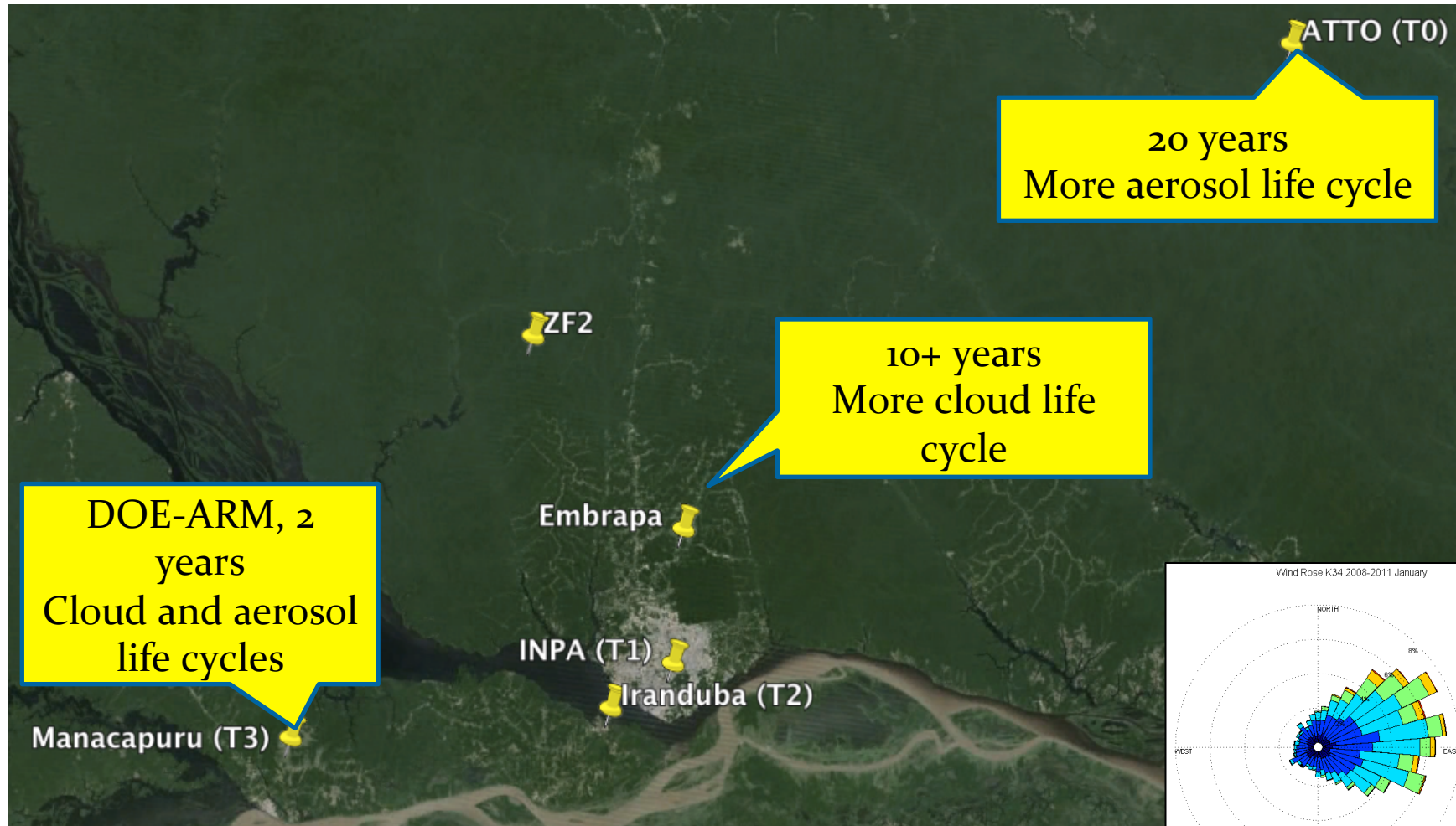
ZF2

Aconvex

Manaus

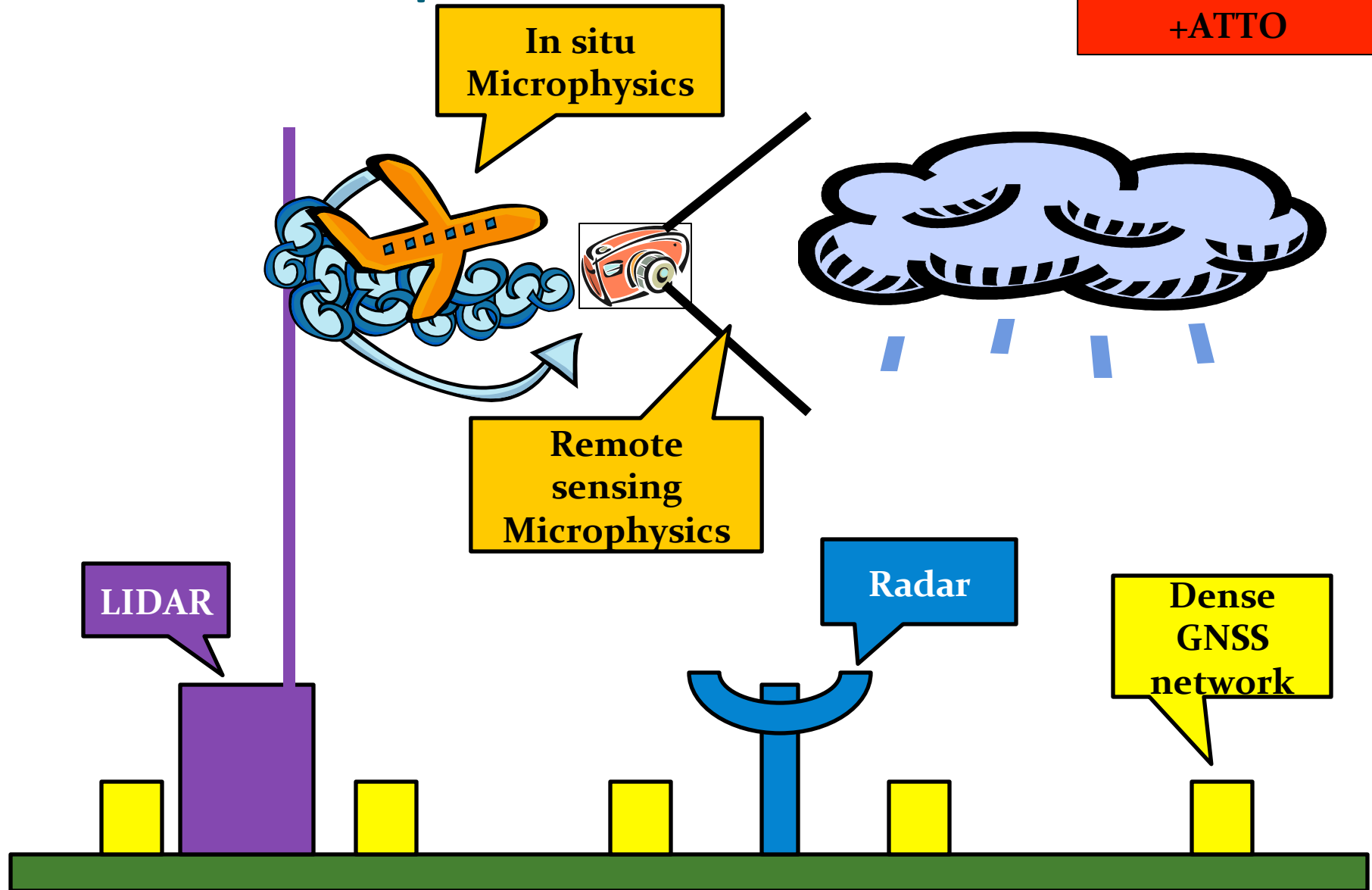
GoAmazon

GoAmazon + CHUVA

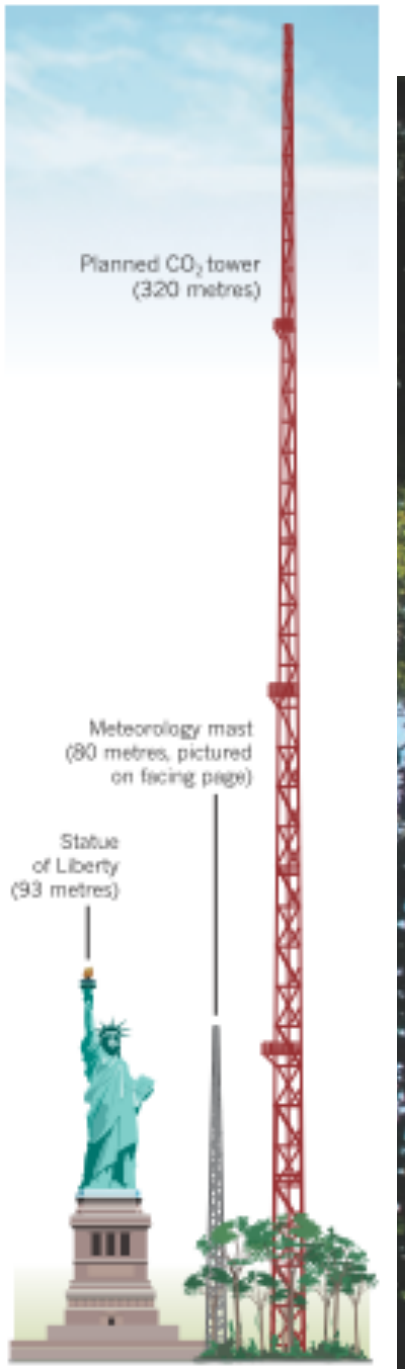


Futuros experimentos

+ CHUVA
+GoAmazon
+ATTO



ATTO – Torre pequena (80m)



ACONVEX – Aerosols, Clouds, Convection Experiment in Amazon



Raman Lidar



Rain Radar



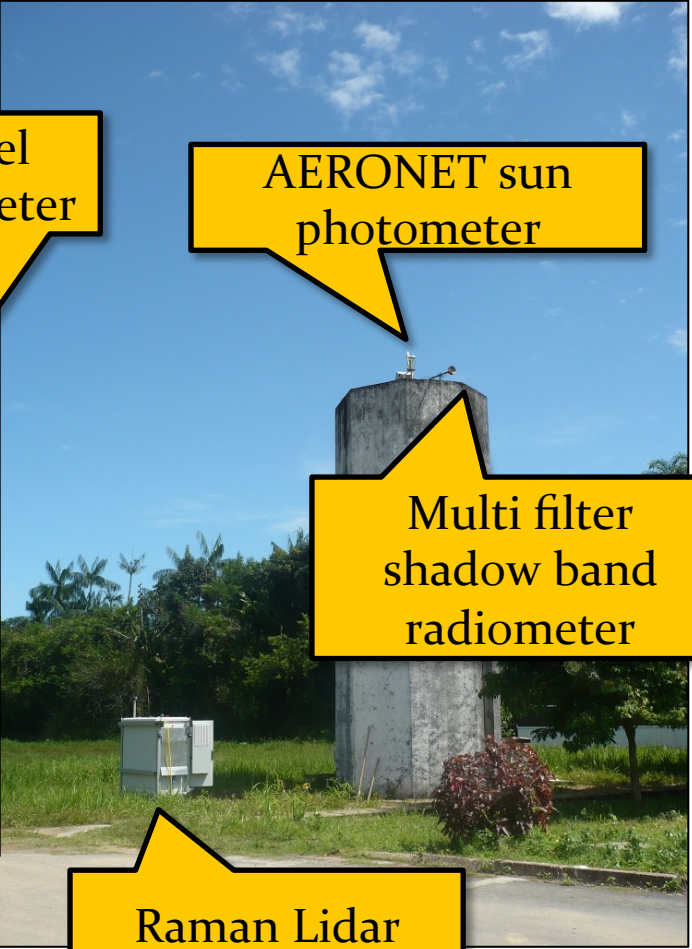
GNSS Trimble



Parsivel
disdrometer



AERONET sun
photometer



Multi filter
shadow band
radiometer



Raman Lidar

Manaus ZF₂ site

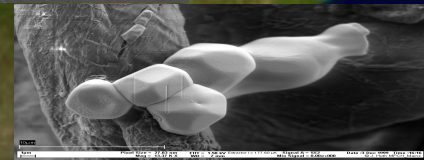
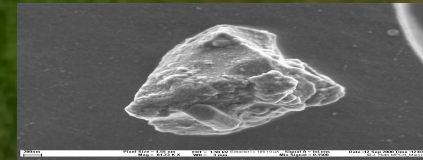
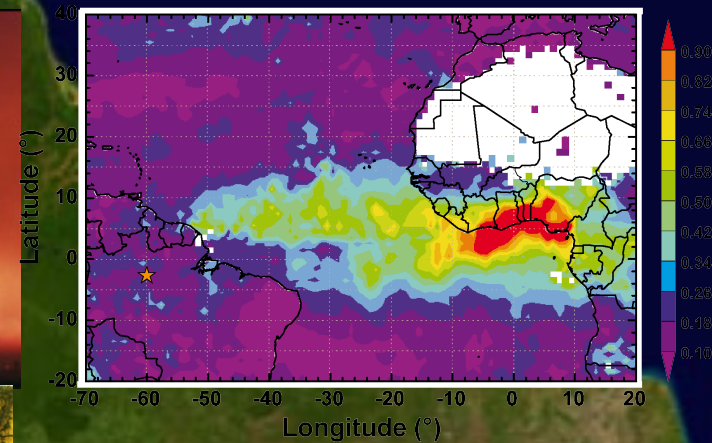
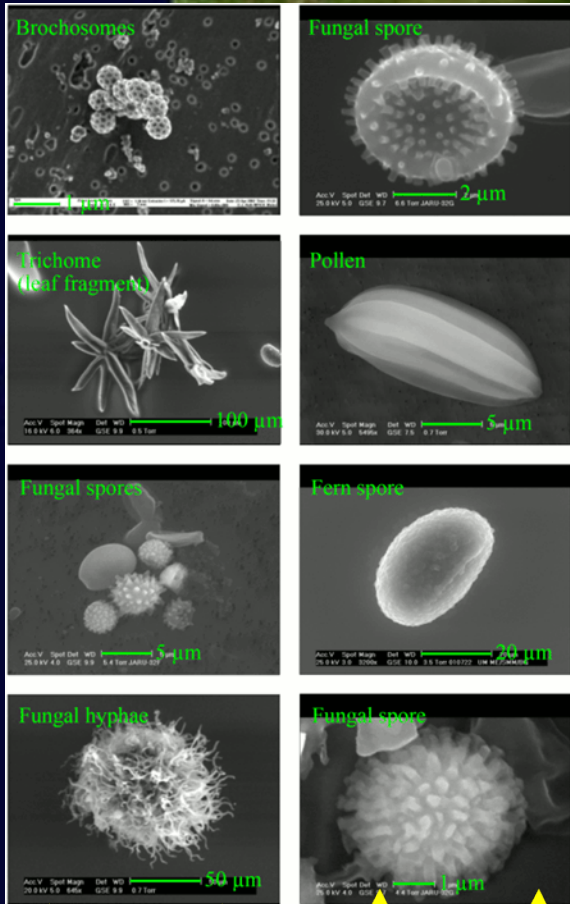


Amazonia: 3 different types of aerosols

Biogenic (primary and SOA)

Biomass Burning

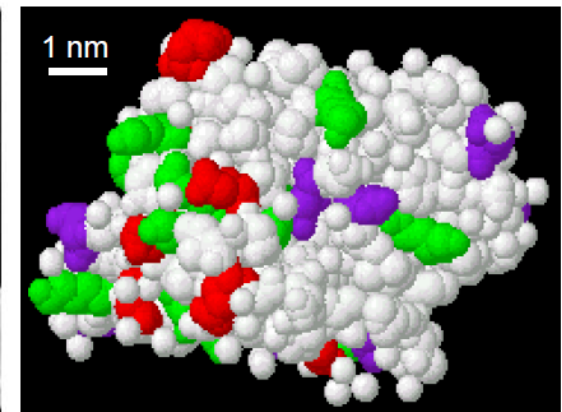
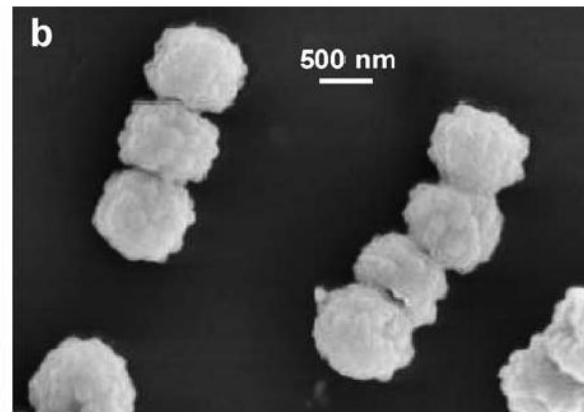
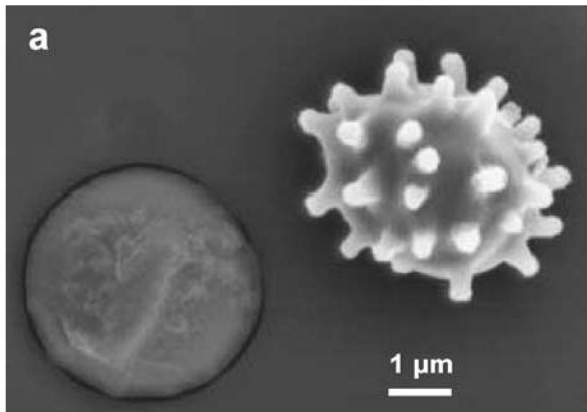
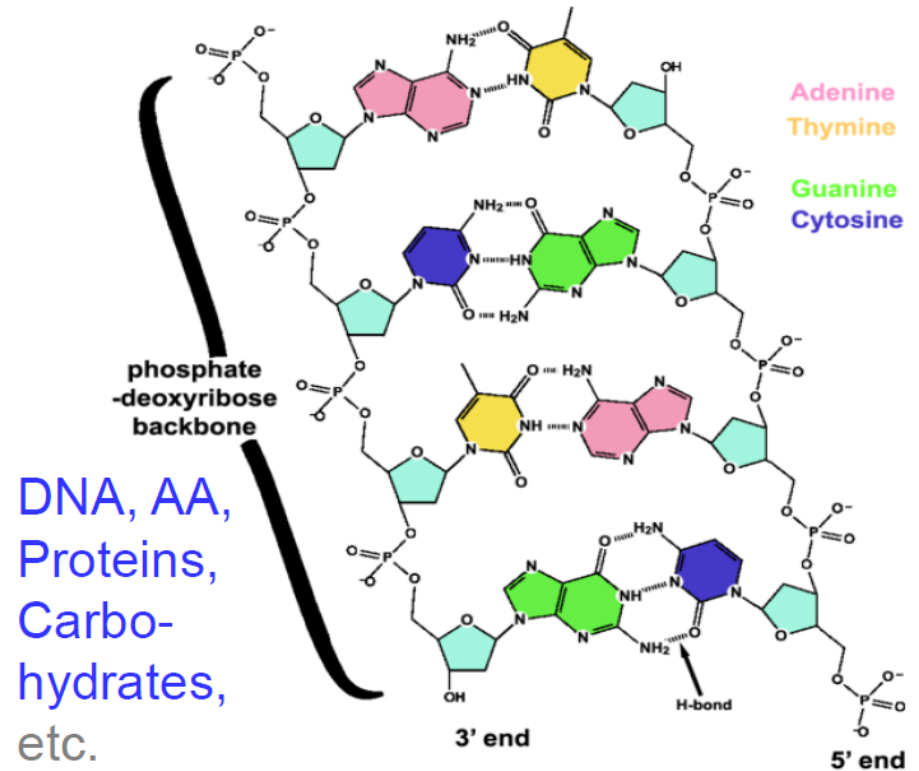
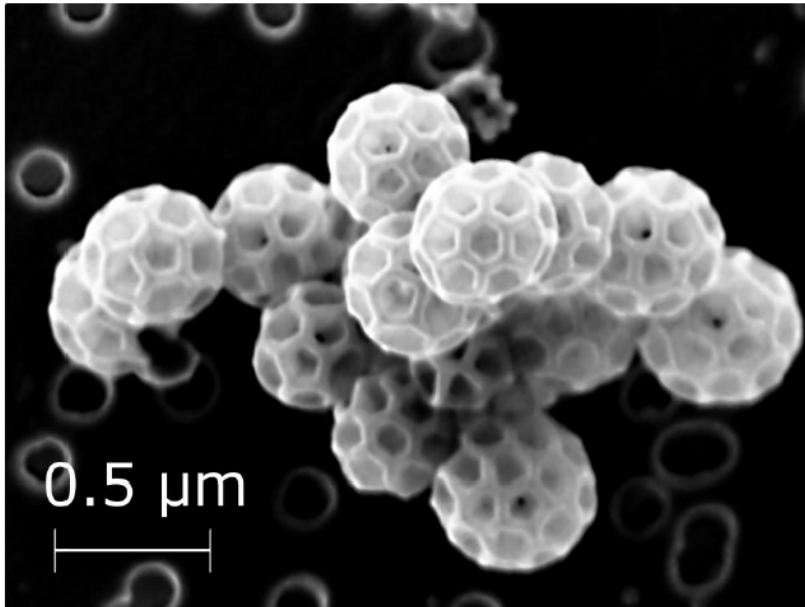
Dust from Sahara



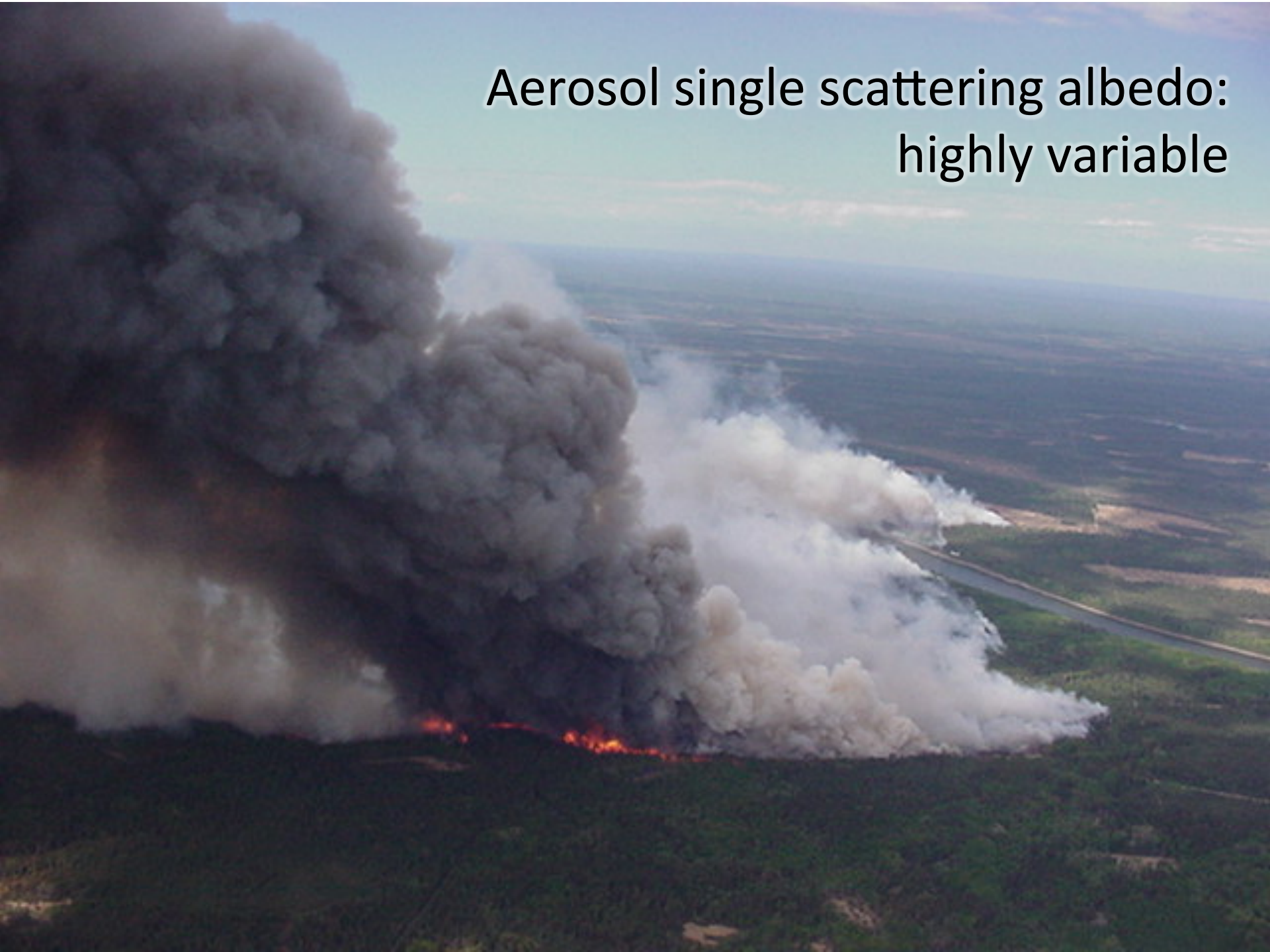
Each with VERY different properties and impacts

Biological Particles & Molecules

Bacteria, Brochosomes, Spores, Pollen, Plant Debris, etc.

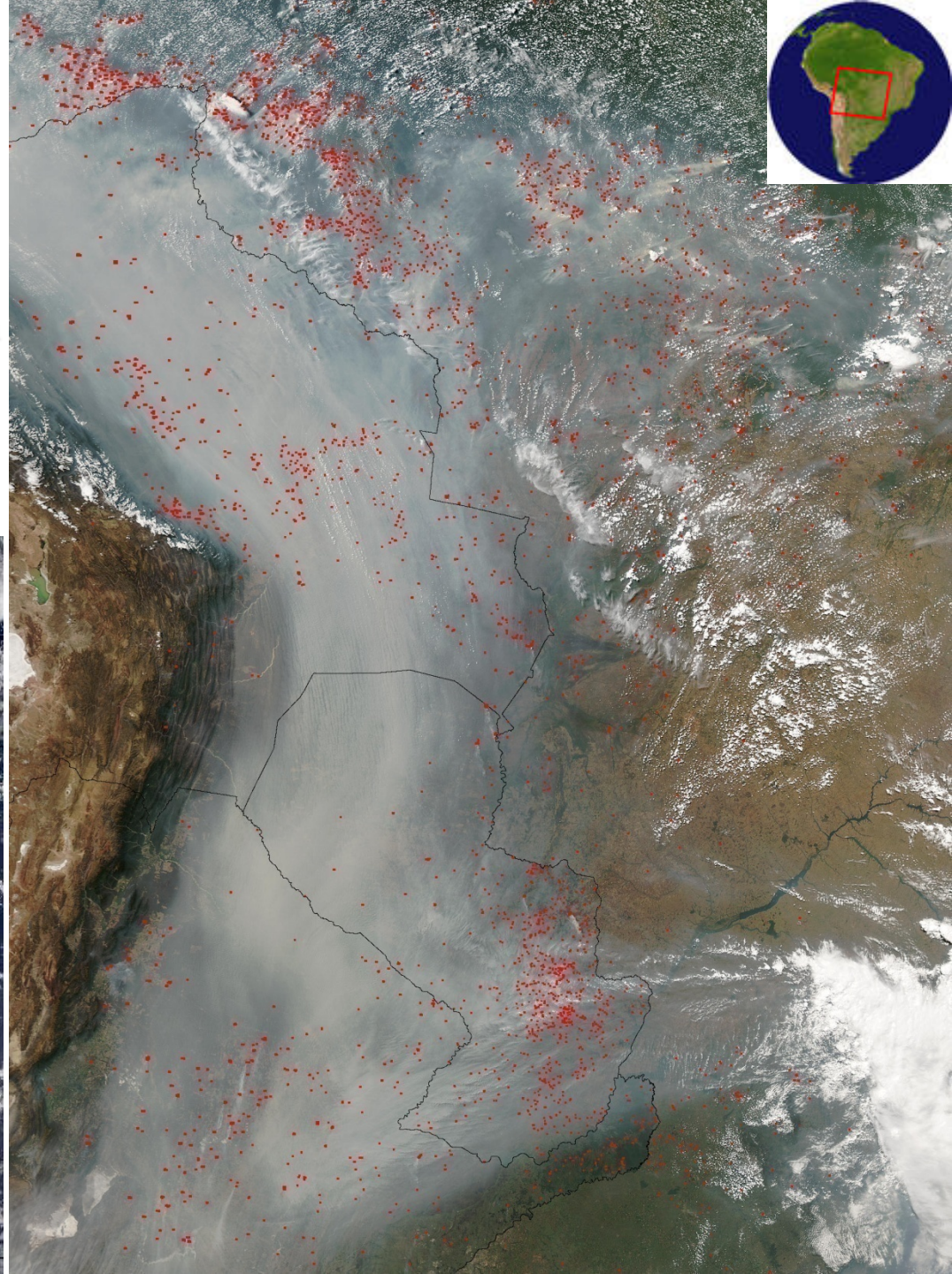
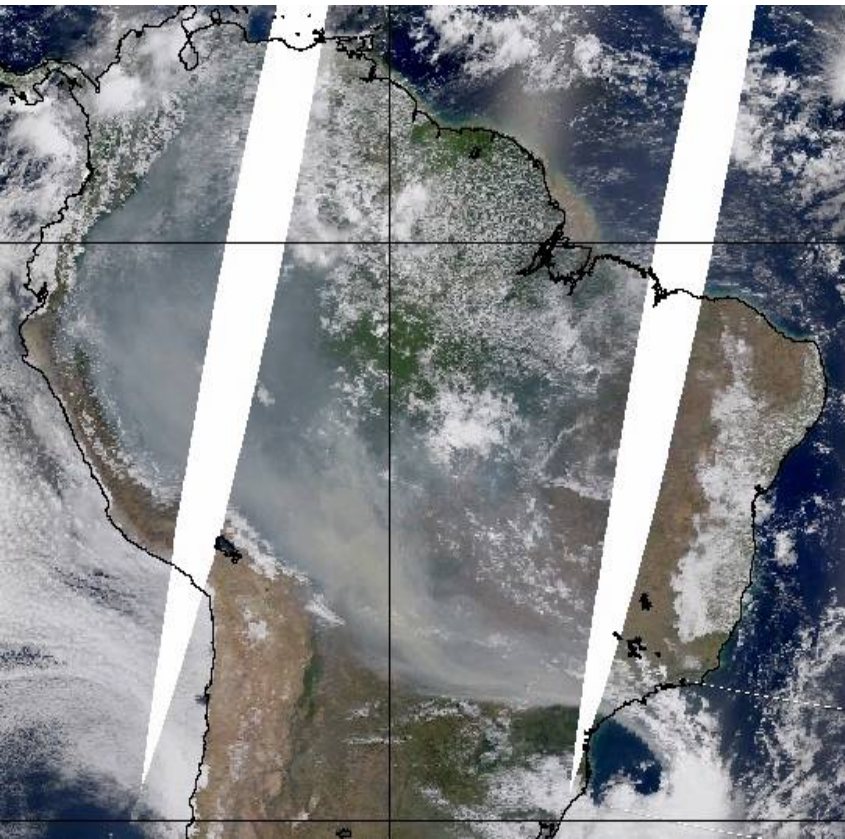


Aerosol single scattering albedo:
highly variable

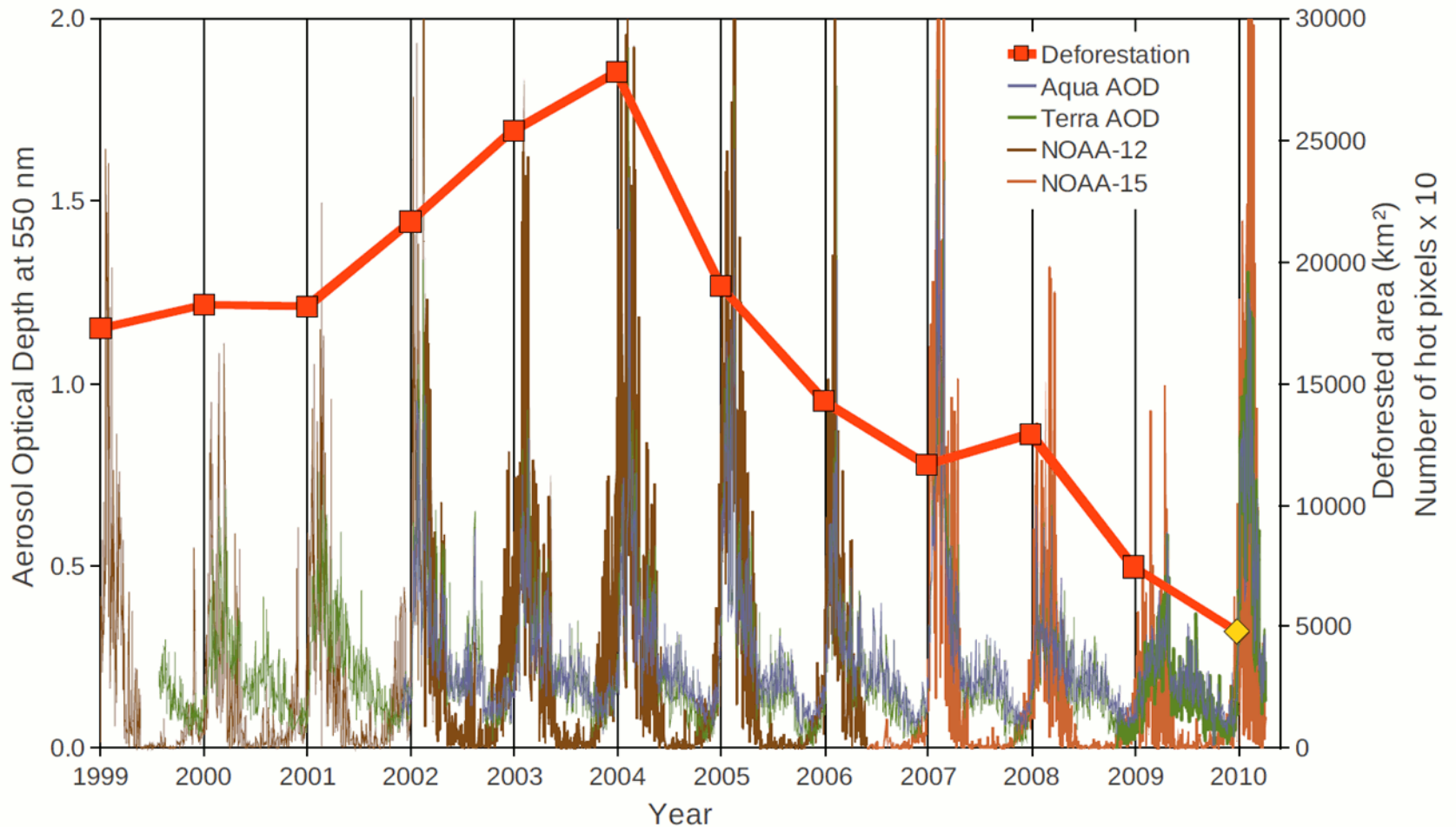


Large scale aerosol distribution in Amazonia

- Severe health effects on the Amazonian population (about 20 million people)
- Climatic effects, with strong effects on cloud physics and radiation balance.
- Changes in carbon uptake and ecosystem functioning



Yearly deforestation with MODIS AOD and hot pixels from NOAA



Yearly deforestation over the Brazilian Amazon region (INPE, 2010) compared to MODIS daily smoke optical depth and the daily number of hot pixels from NOAA-12 and NOAA-15.



Amazonia

Average aerosol forcing clear sky

Top: - 10 w/m²



Atmosphere: + 28 w/m²

Surface: - 38 w/m²

Conditions: surface: forest vegetation
AOT ($\tau=0.95$ at 500nm); 24 hour average
7 years (93-95, 99-02 dry season Aug-Oct)

INDOEX

average aerosol forcing clear sky

Top: - 7±1 w/m²



Atmosphere: + 16±2 w/m²

Surface: - 23±2 w/m²

Conditions: surface: ocean
AOT ($\tau=0.3$ at 630 nm); 24 hour average
Jan-Mar 99

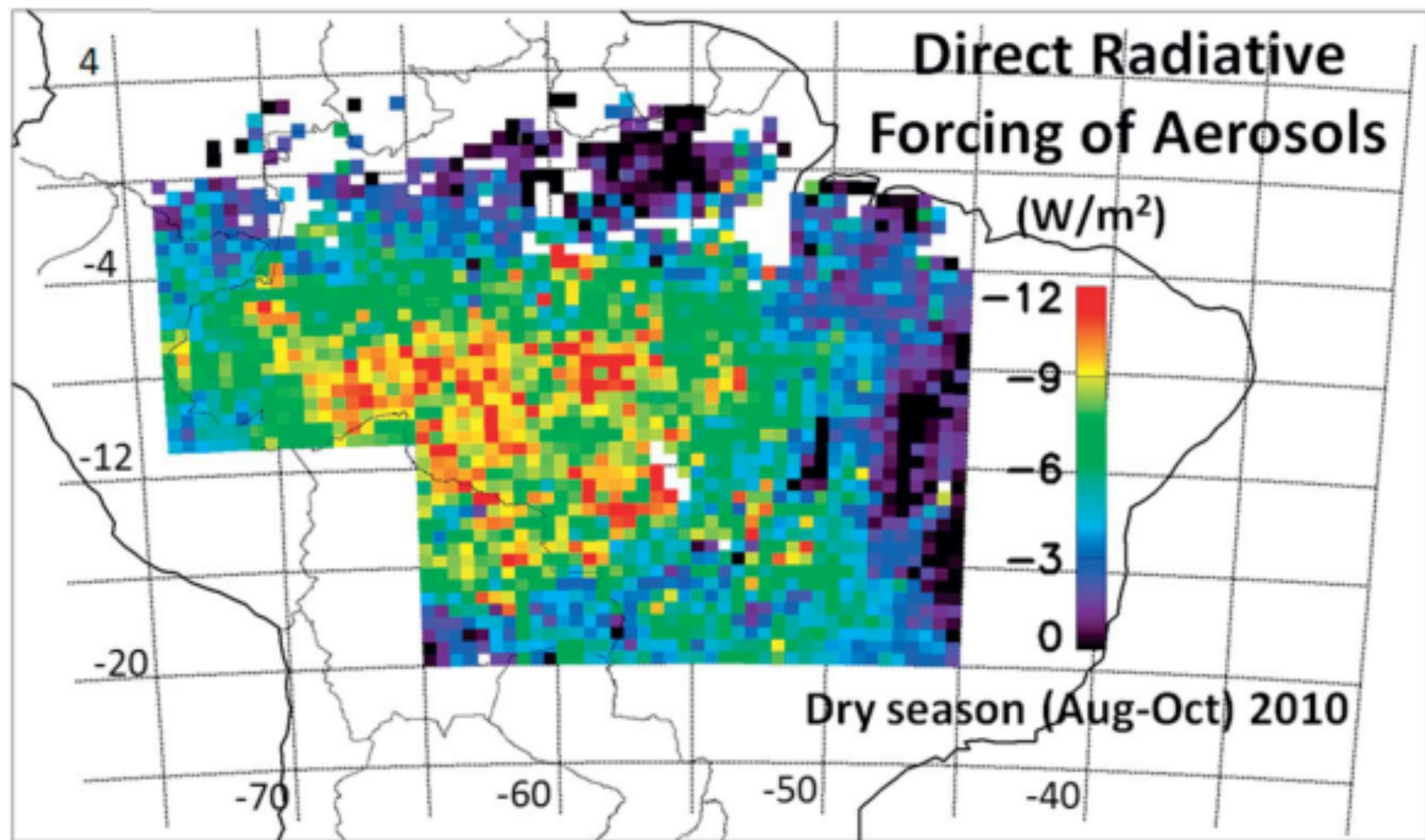
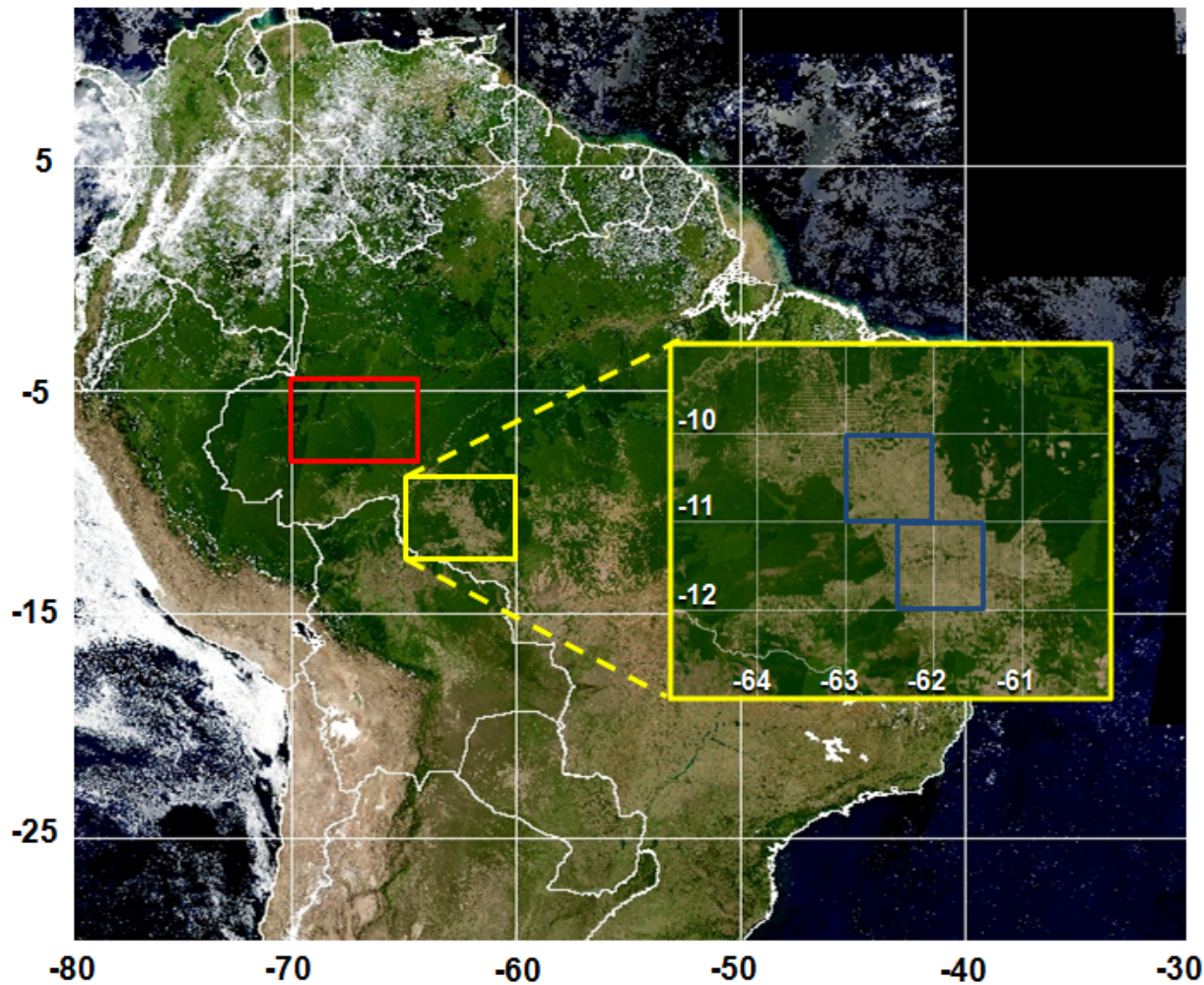


Fig. 18 Average spatial distribution of the direct radiative forcing (DRF) of biomass burning aerosols in Amazonia during the dry season (August to October) of 2010. Forcing derived from calculations using a combination of MODIS and CERES sensors data. During this three-month period, the daily-average radiative forcing of aerosols for the whole area was on average $-5.3 \pm 0.1 \text{ W m}^{-2}$.



**Mean Diurnal
Radiative Forcing
due to change in
surface albedo:
 $-8.0 \pm 0.9 \text{ W/m}^2$**

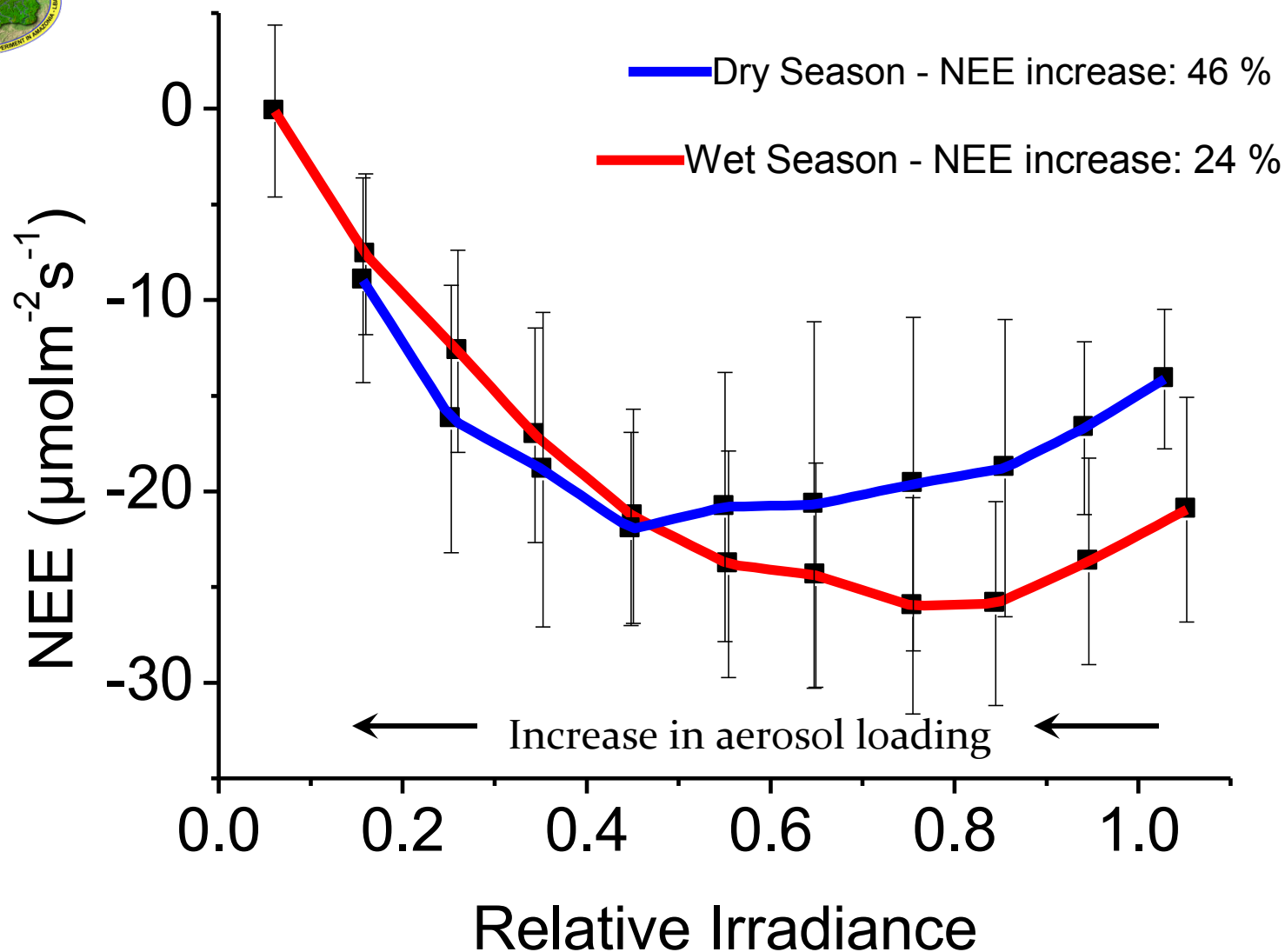
**Mean Diurnal Aerosol
Forcing Efficiency:**
Forest: $-22.5 \pm 1.4 \text{ W/m}^2$
Cerrado: $-16.6 \pm 1.7 \text{ W/m}^2$

Land-use change radiative forcing.
Forested areas are selected in red and
deforested areas are selected in blue.

Strong aerosol effect on forest photosynthesis diffuse radiation have a large effect on CO₂ fluxes

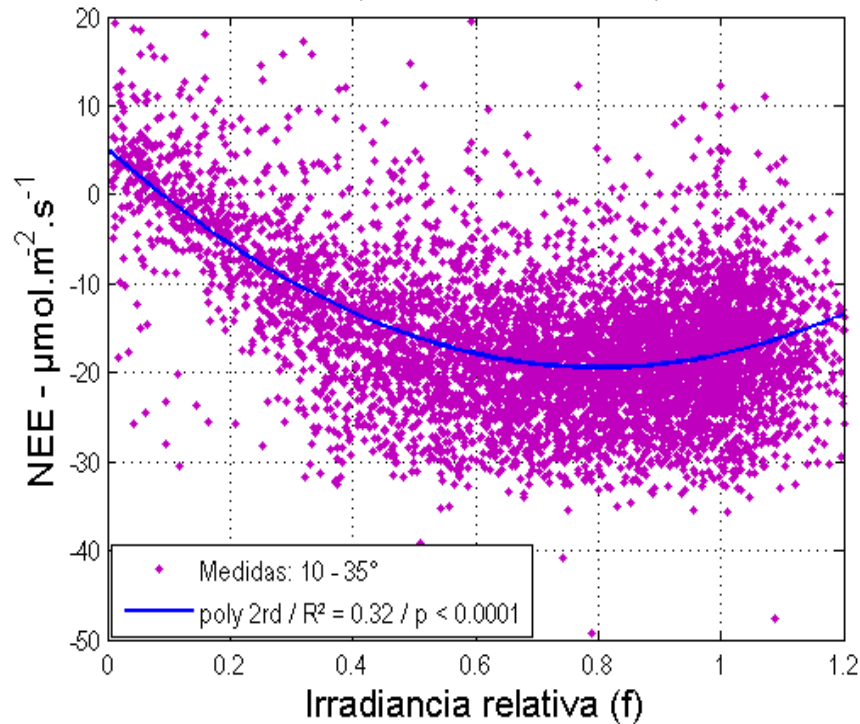


Amazonia Rondonia Forest site 2000-2001



Aerosols effects on NEE – Manaus and

K34 (Jul-Nov /1999-2009)



(Glauber Cirino, INPA, 2013)

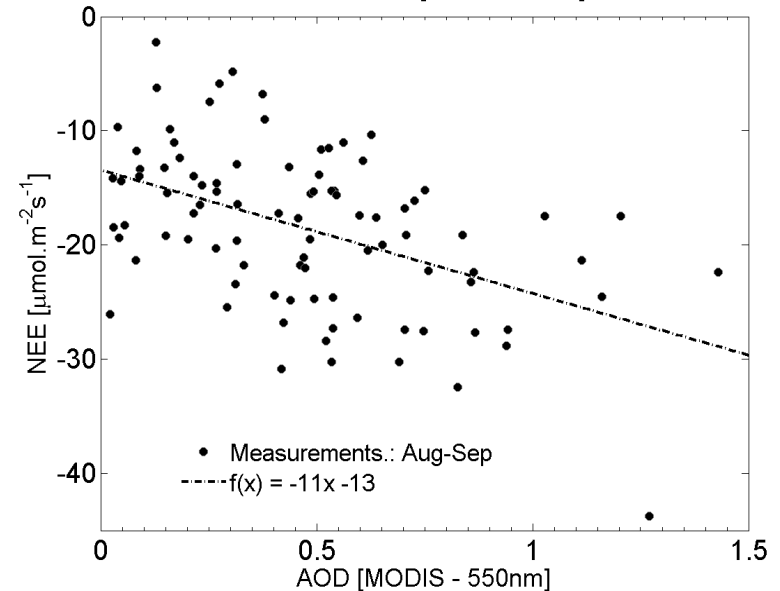
f max: ~ 0,80

NEE (max): ~ -20 $\mu\text{mol}/\text{m}^2\text{s}$
AOT: ~ 0,5

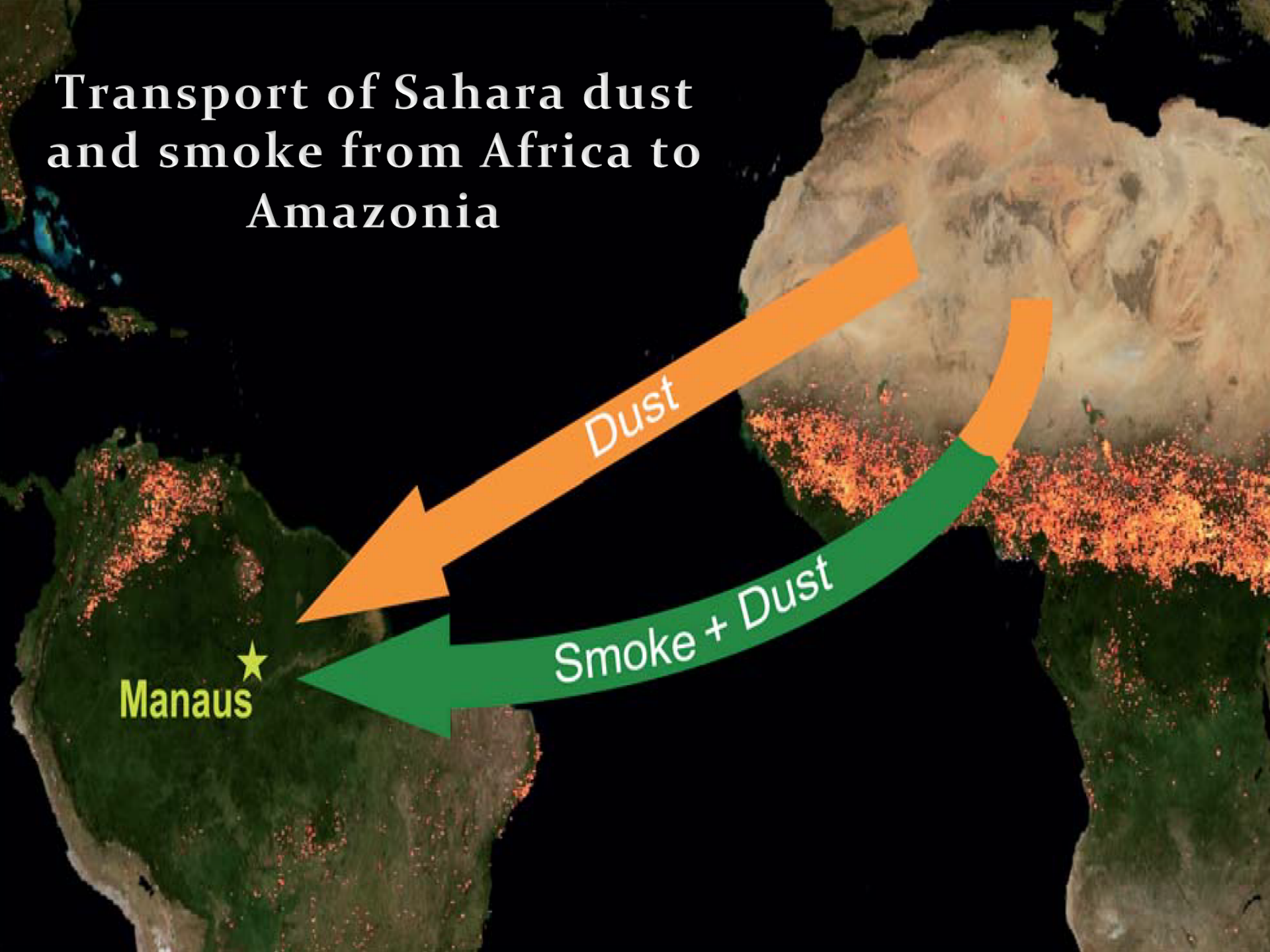
Manaus K34 tower



NEE vs. AOD [Ji-Parana-RO]



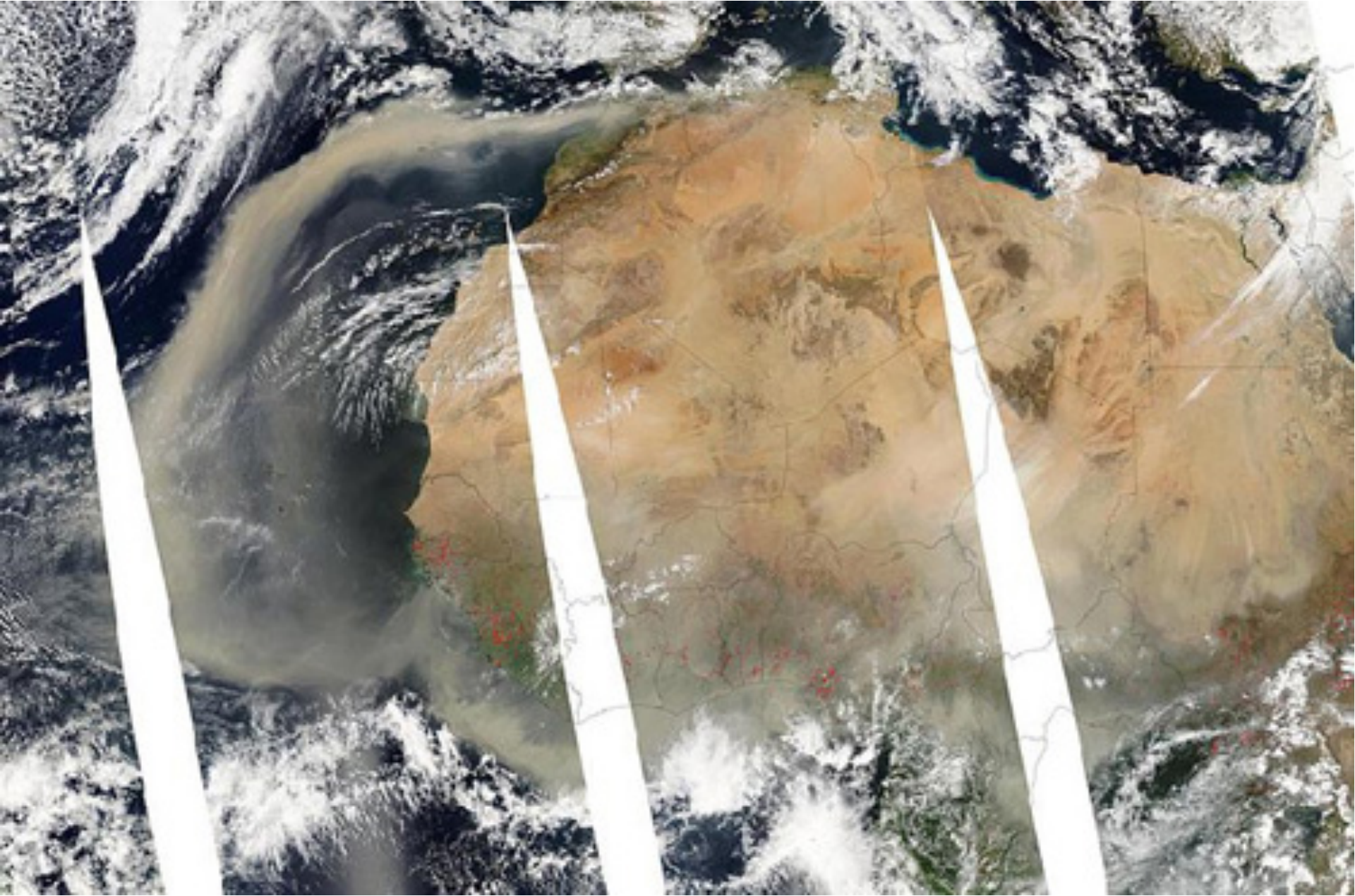
Transport of Sahara dust and smoke from Africa to Amazonia

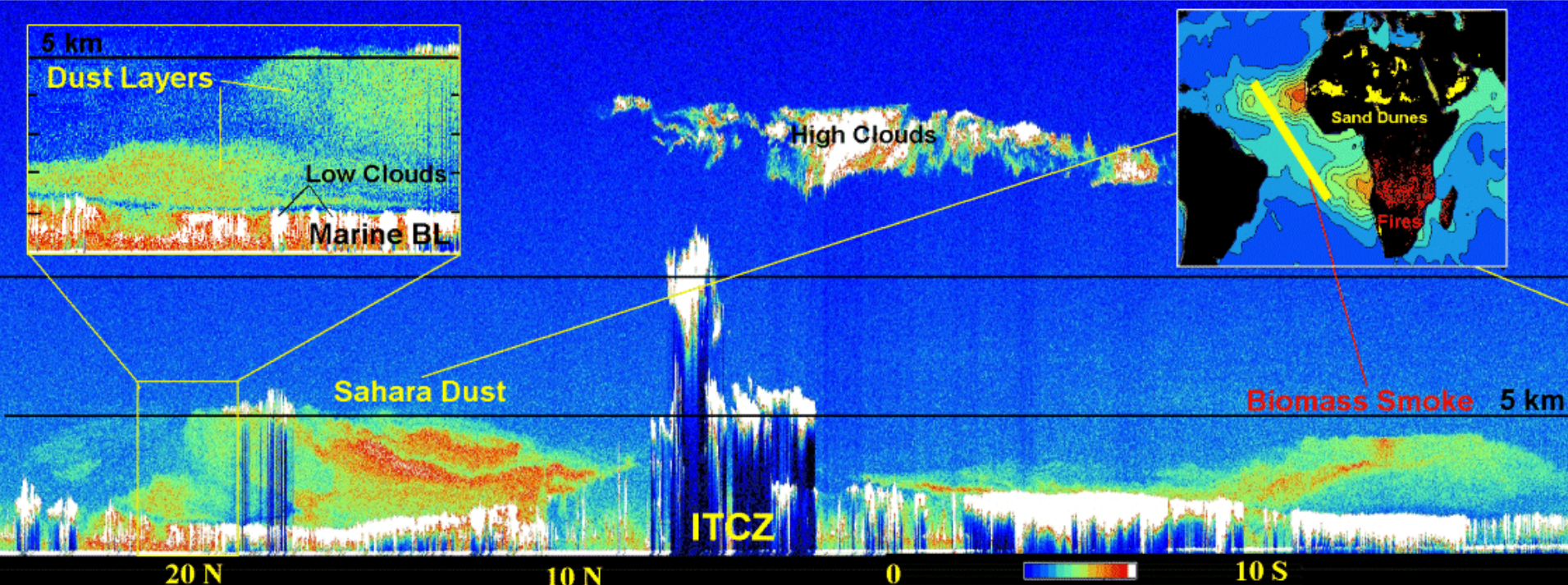
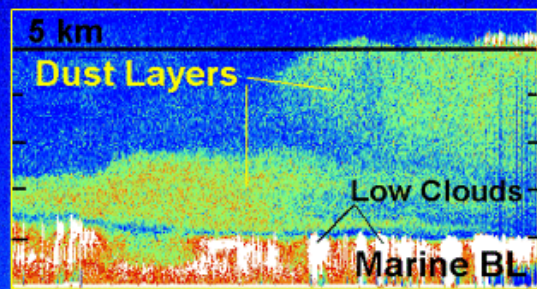
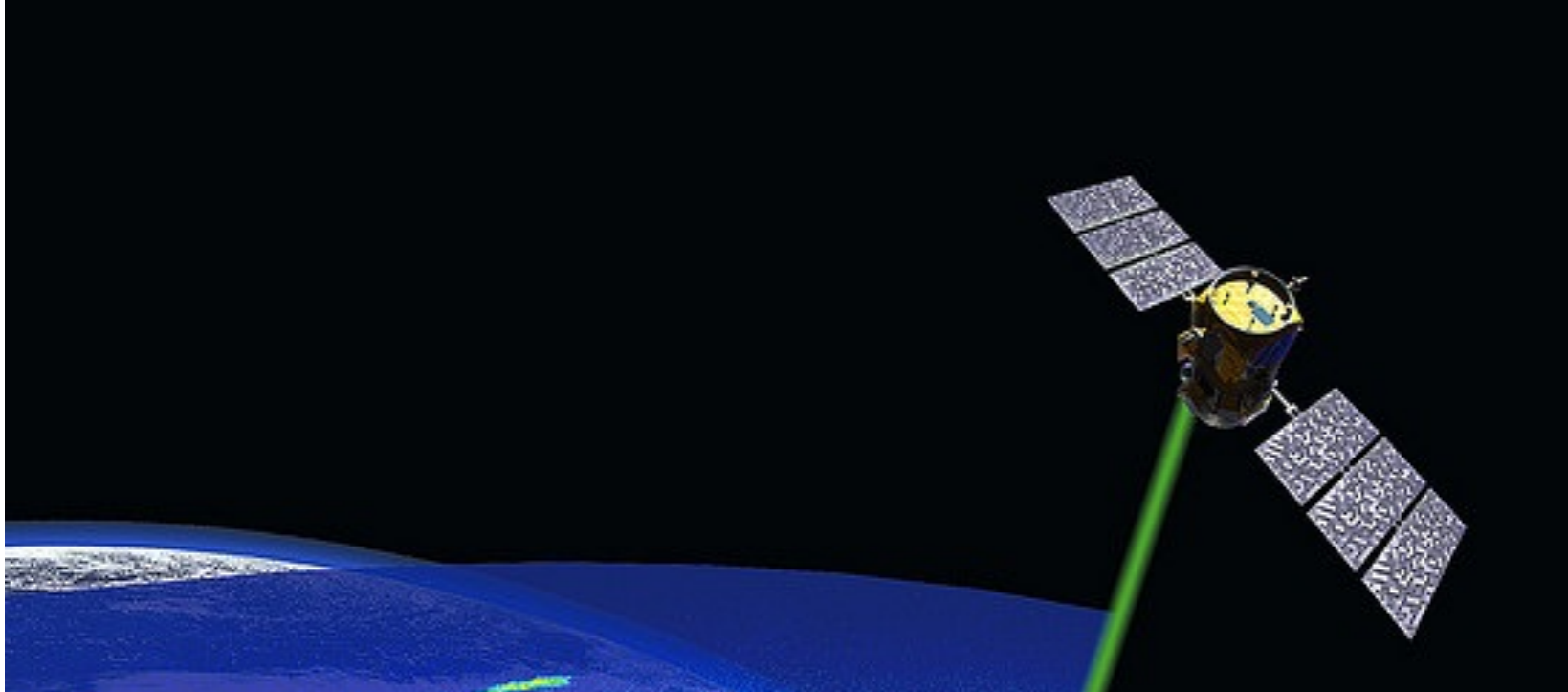


Dust

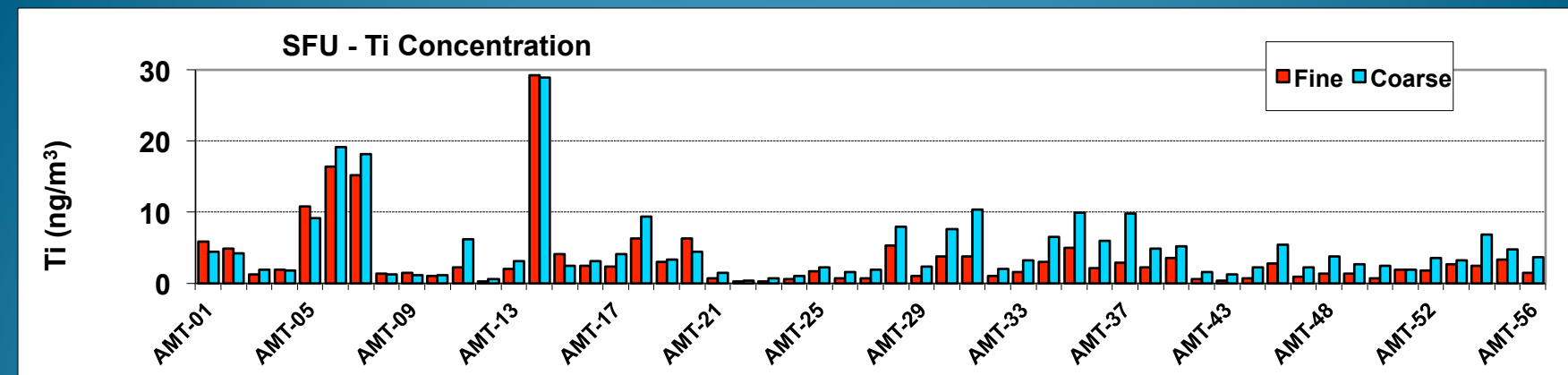
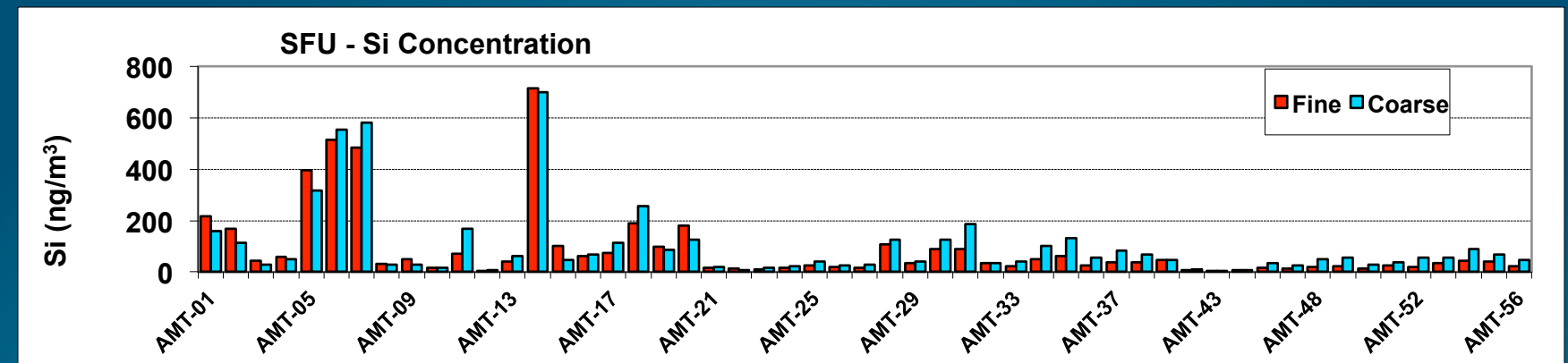
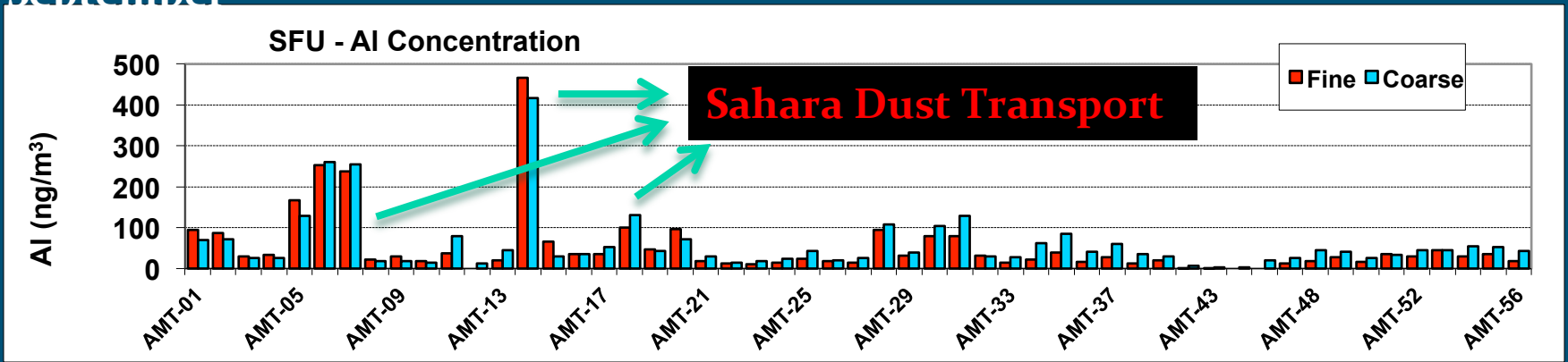
Smoke + Dust

Manaus



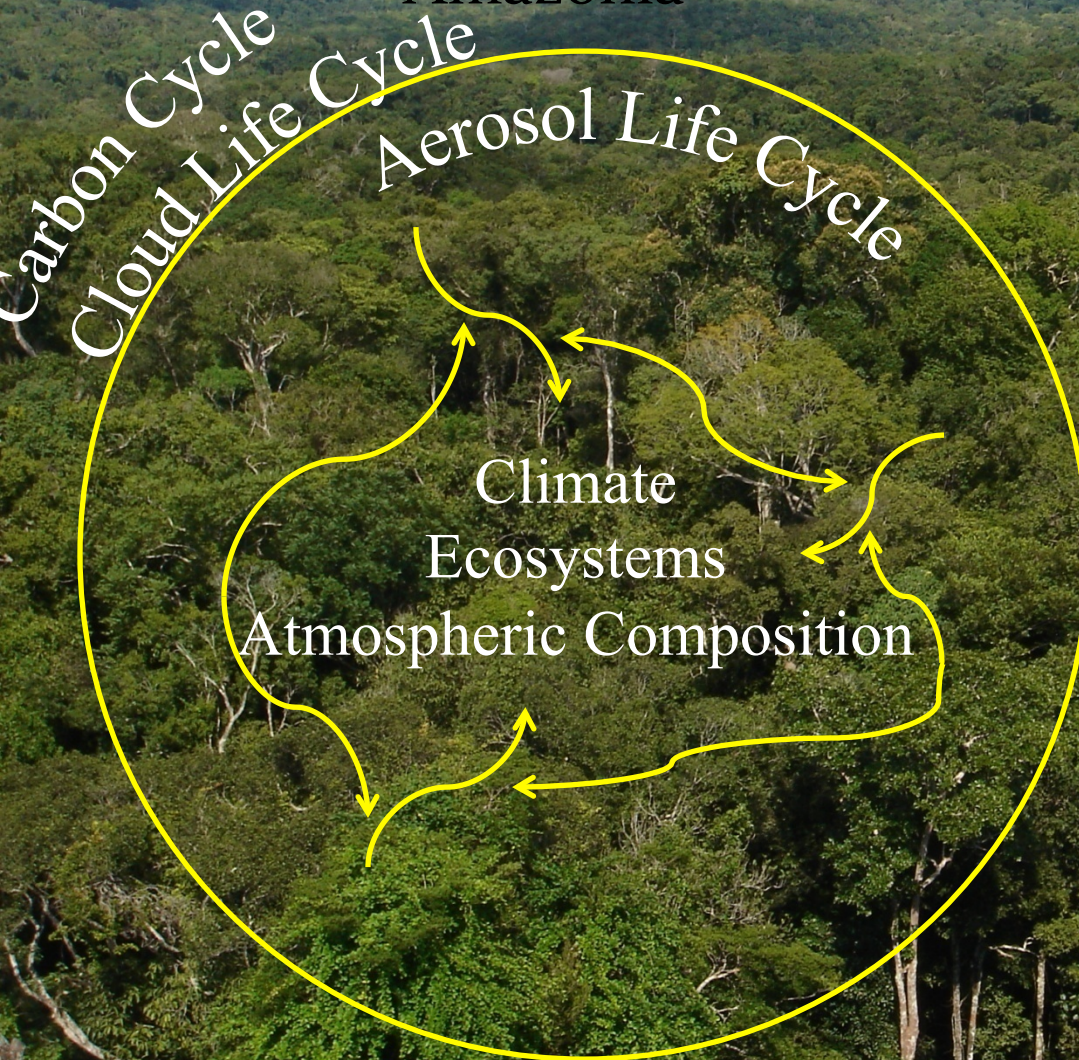


Al, Si and Ti elemental Concentration for fine and coarse mode aerosols Feb. to September



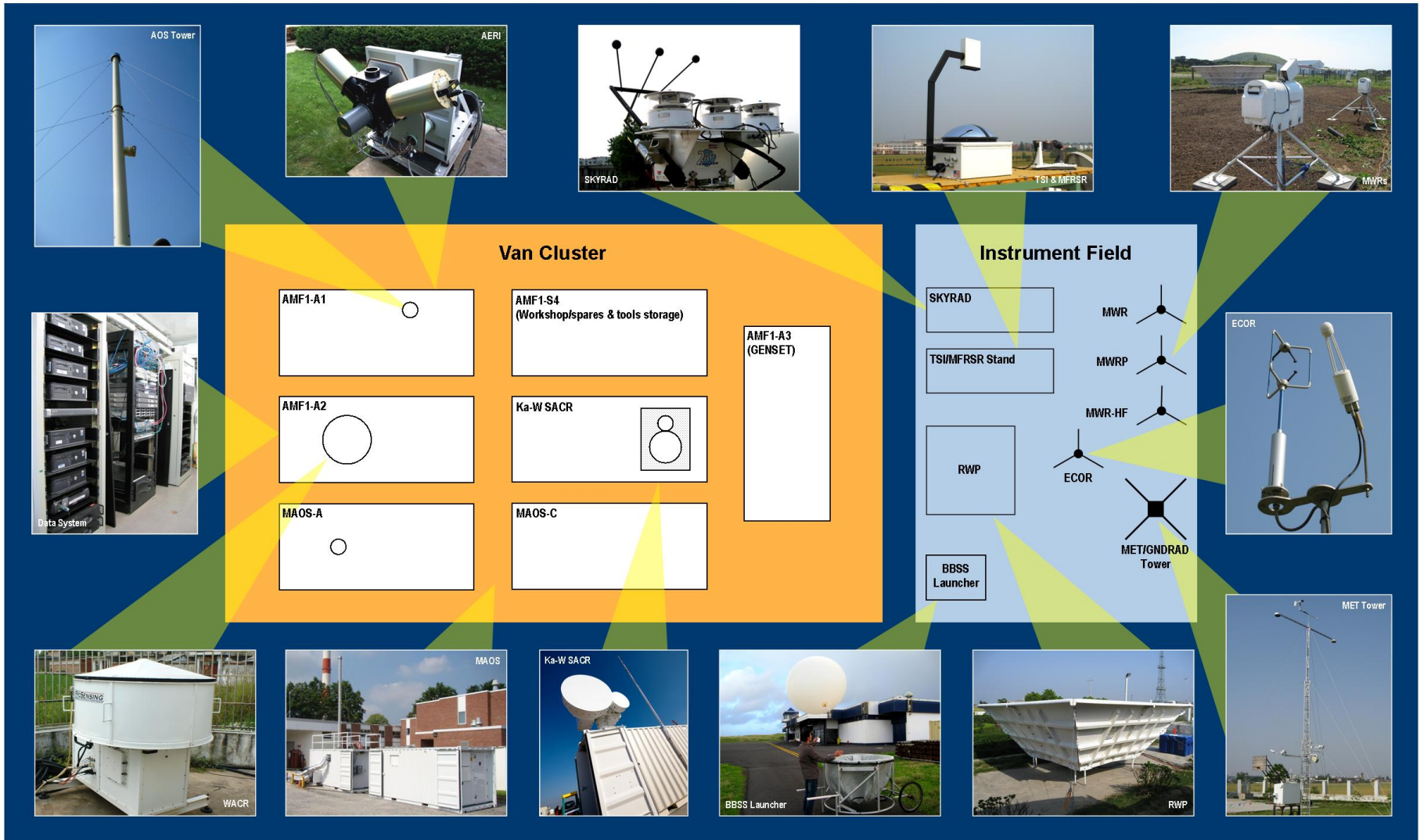
GoAmazon 2014-2015

Ecosystems, Atmosphere
Composition, and Climate in
Amazonia

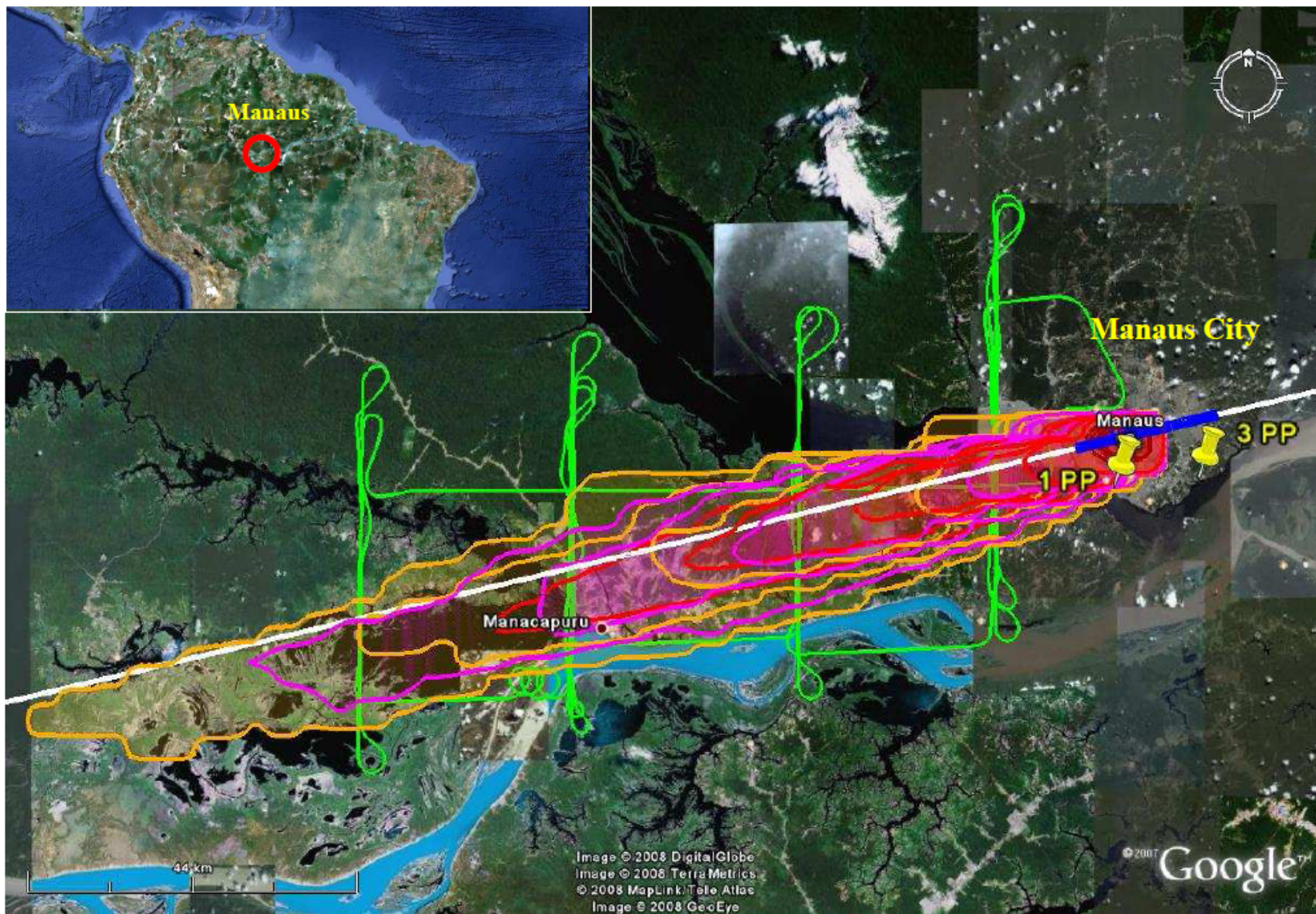


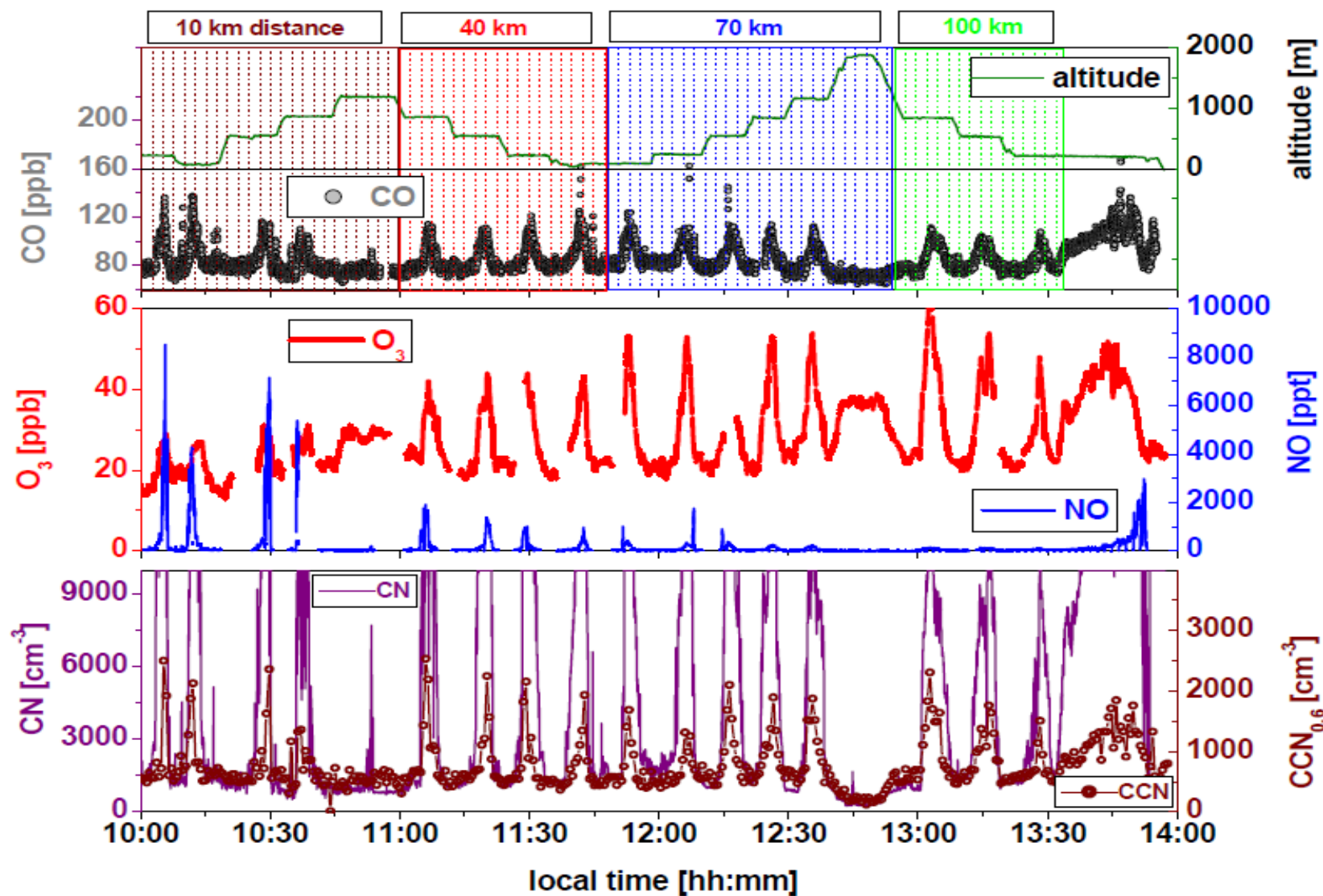
“ARM Mobile Facility in Amazônia” (AMFA2014/5)

ARM Mobile Facility One - Typical Deployment



Experimento GoAmazon 2014

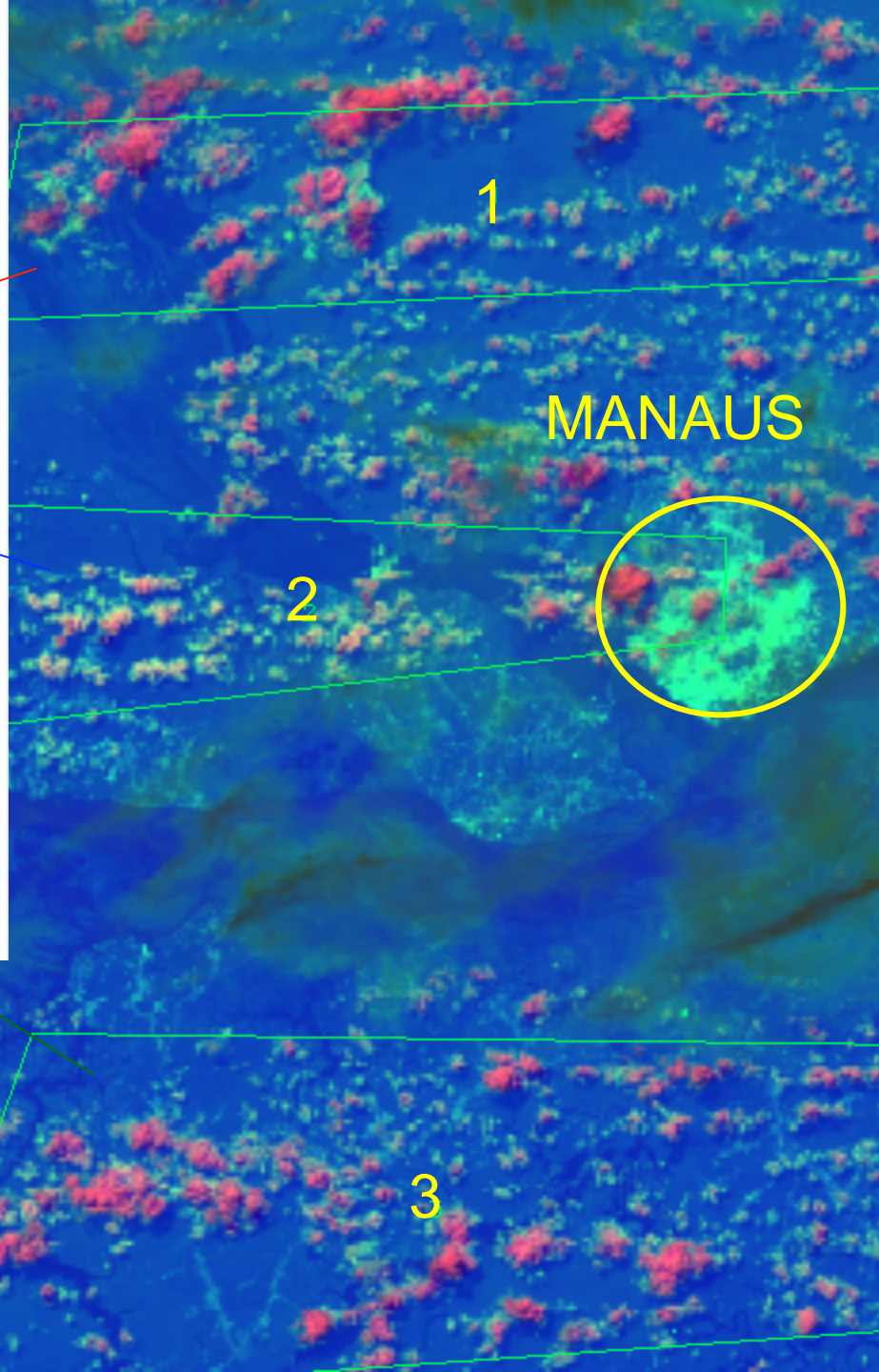
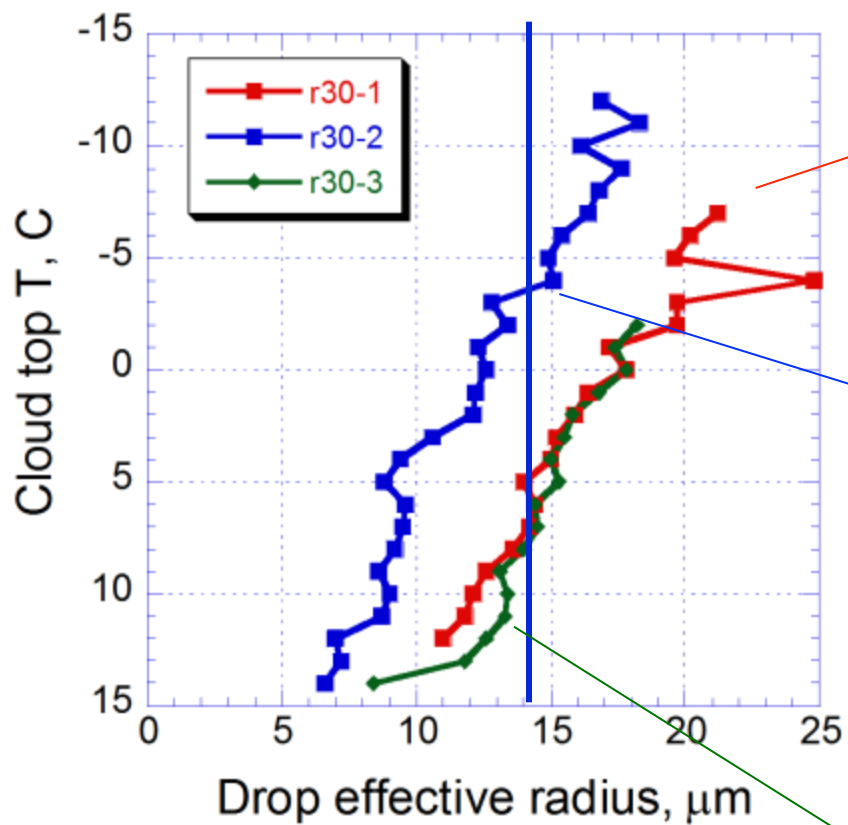




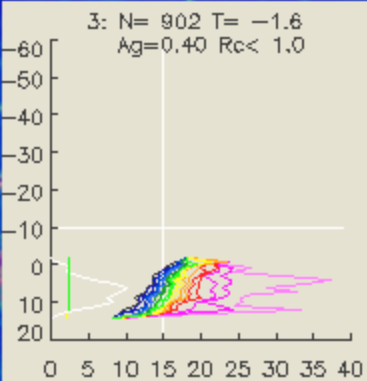
Reference: Kuhn, U.; Ganzeveld, L.; Thielmann, A.; Dindorf, T.; Welling, M.; Sciare, J.; Roberts, G.; Meixner, F. X.; Kesselmeier, J.; Lelieveld, J.; Ciccioli, P.; Kolle, O.; Lloyd, J.; Trentmann, J.; Artaxo, P.; Andreae, M. O., "Impact of Manaus City on the Amazon Green Ocean atmosphere: Ozone production, precursor sensitivity, and aerosol load," *Atmos. Chem. Phys.* **2010**, *10*, 9251-9282.



MANAUS

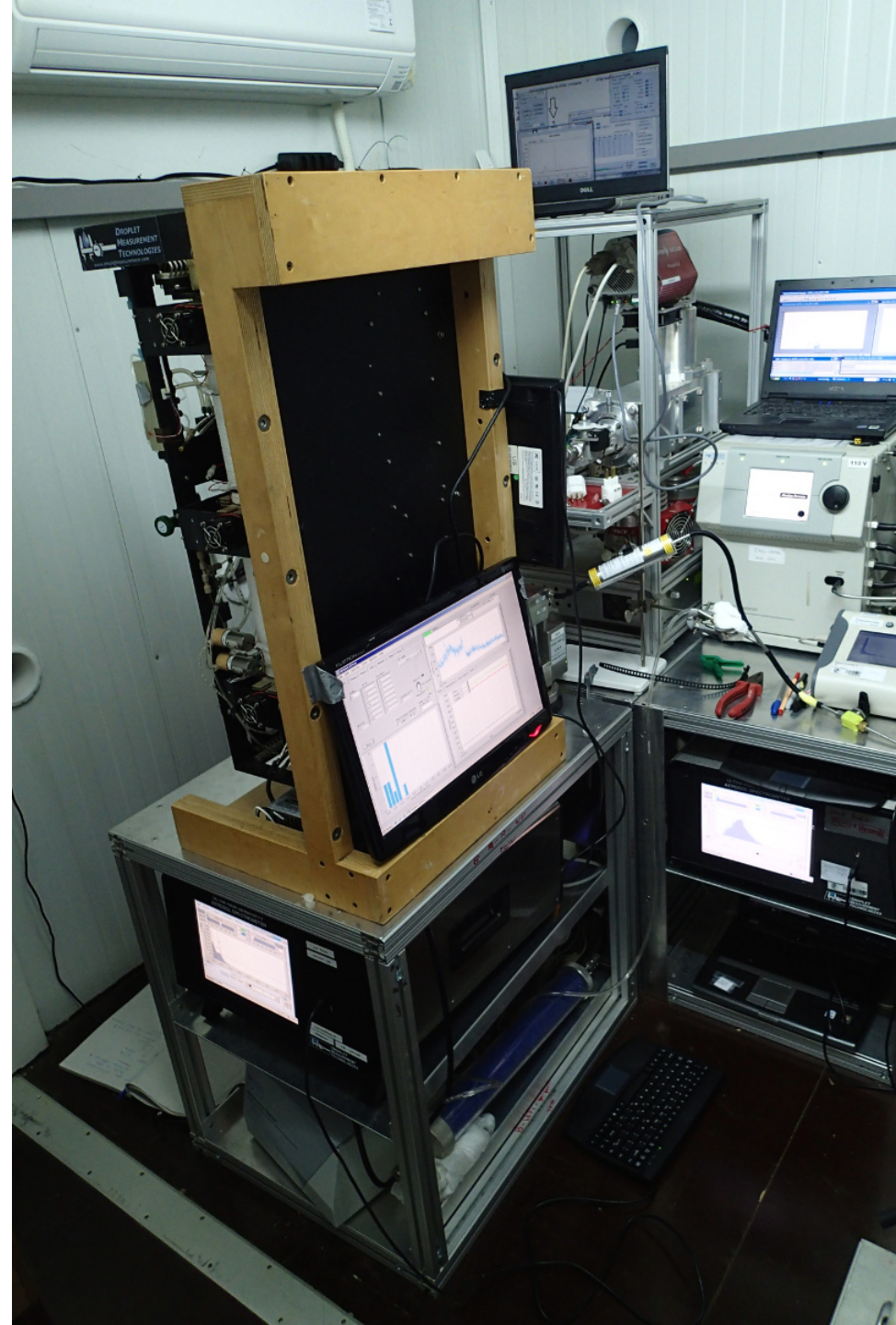


0 5 10 15 20 25 30 35 40



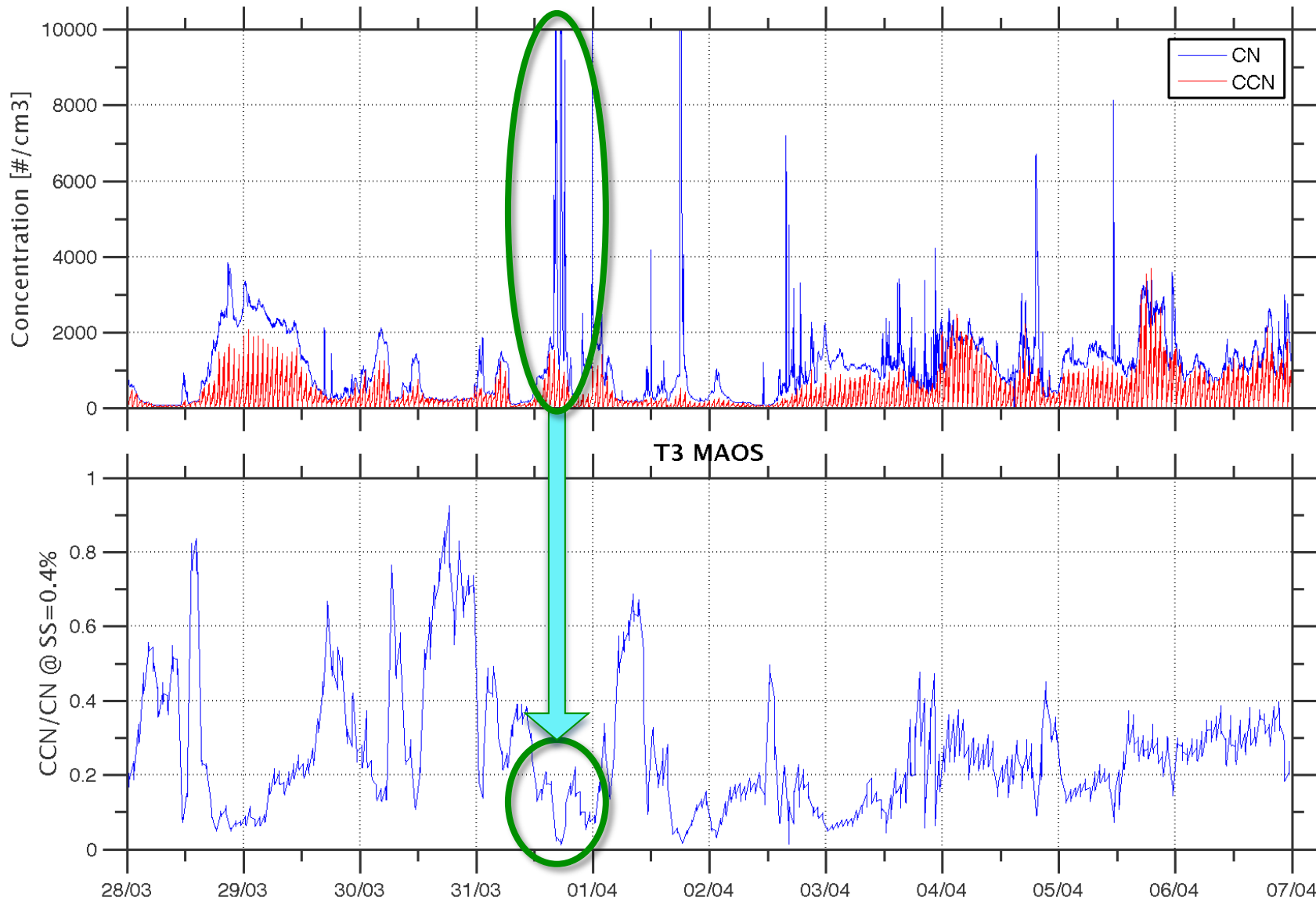
Mira's measurements

- 17-21 March
 - Initial setup
 - Calibration
 - Intercomparison w/ J. Wang at T3
- 22-26 March
 - Moved to ATTO
 - CCNC calibration
 - DMA/UHSAS calibration with PSL's
 - Started data acquis...

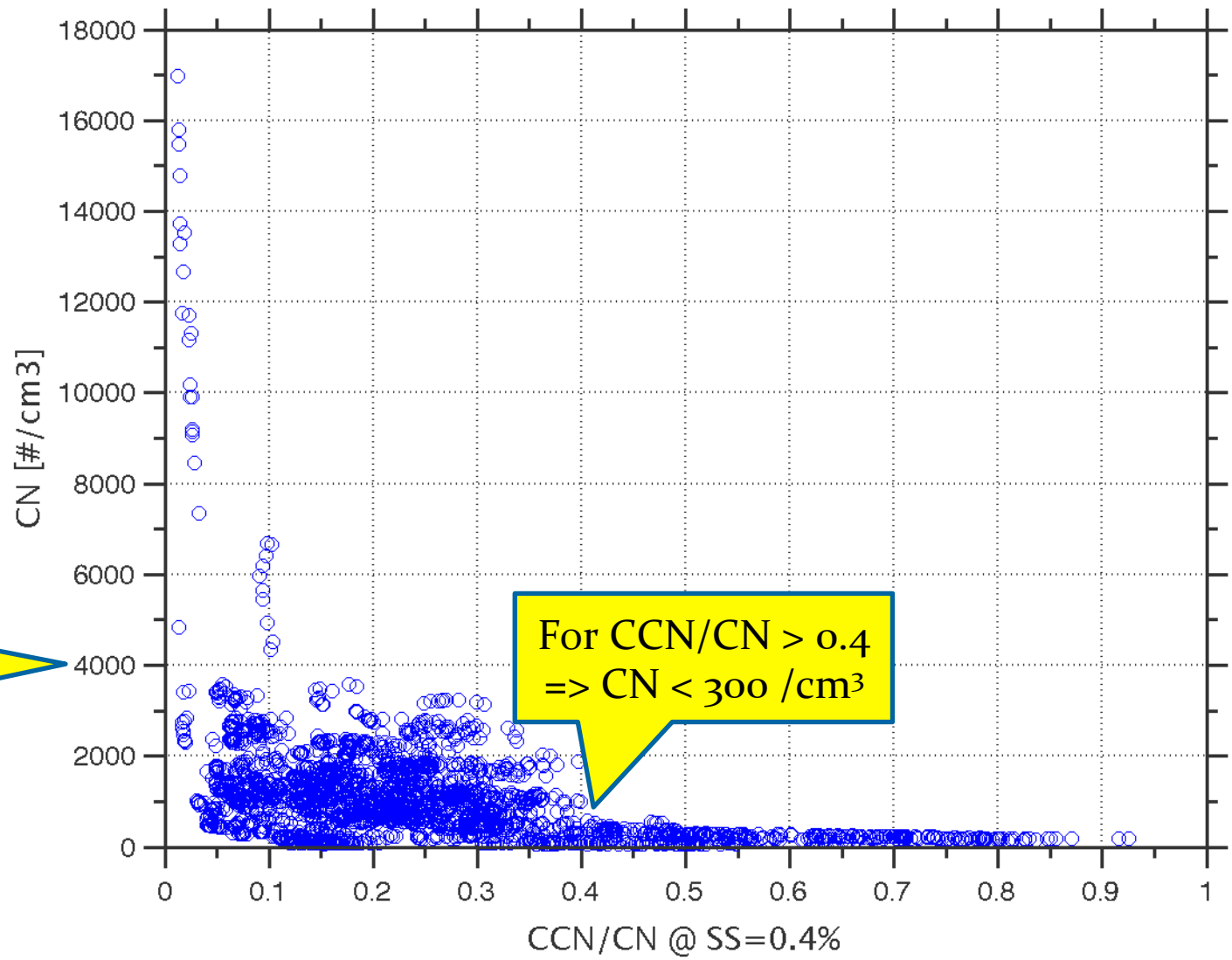


CON/CN MA

ATTO

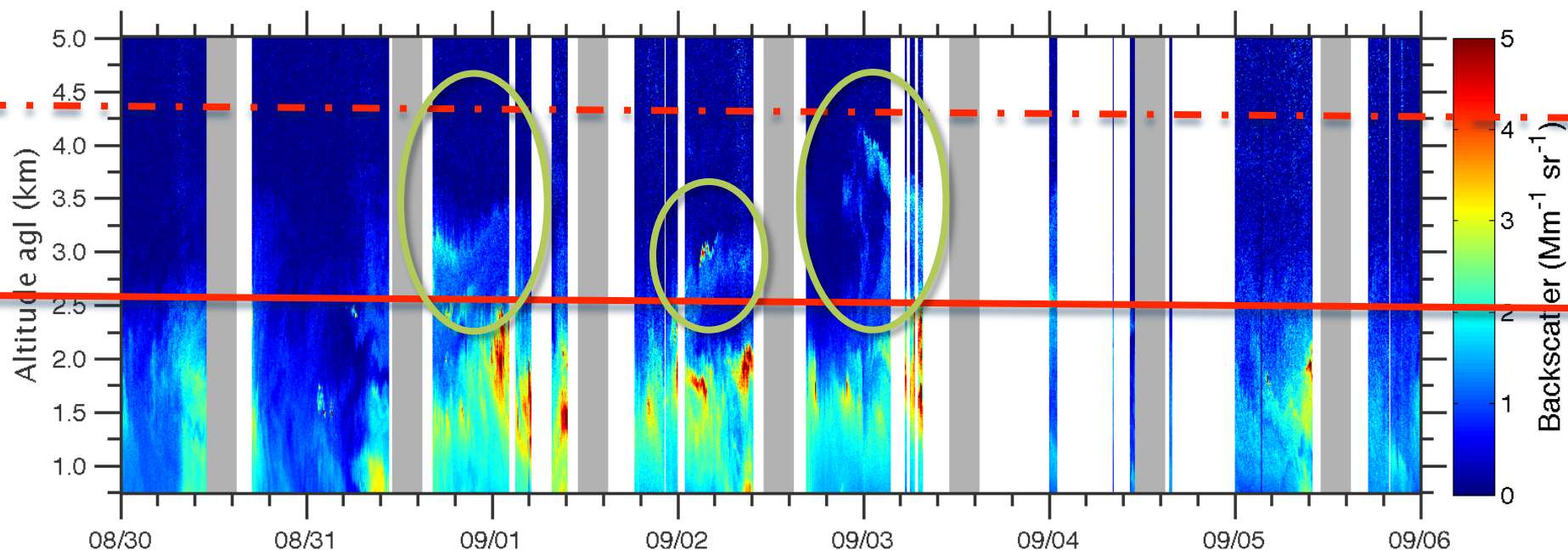


CCN/CN Measurements at ATTO

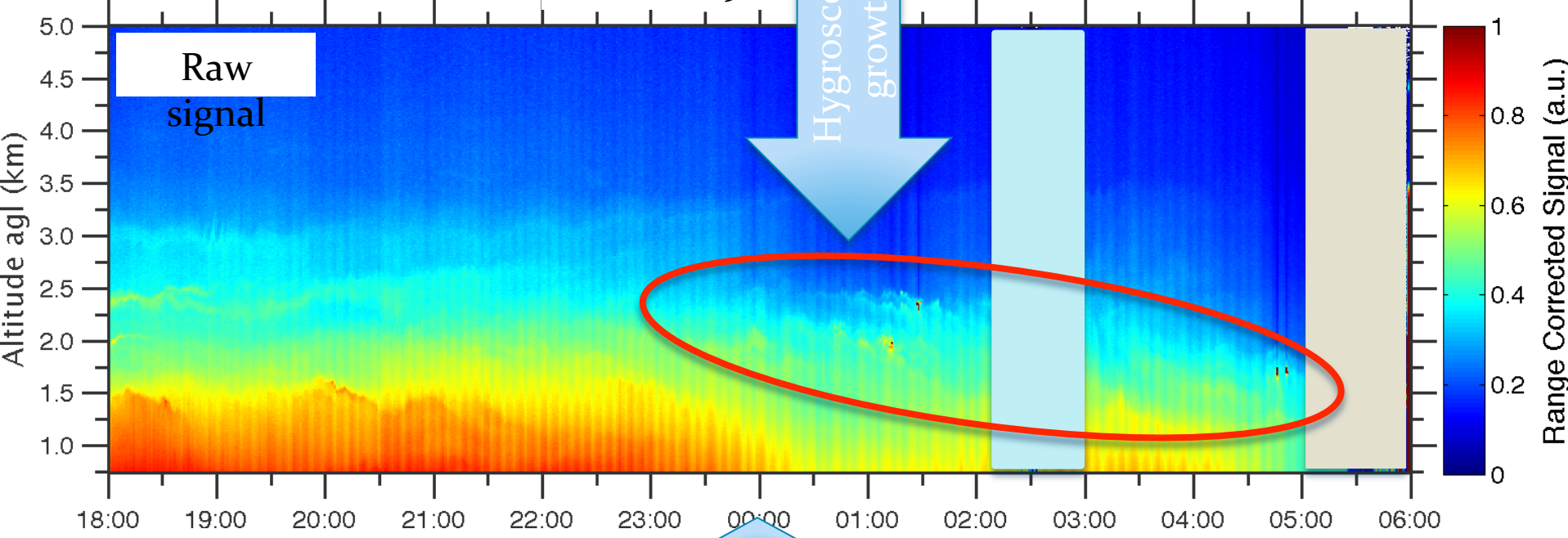
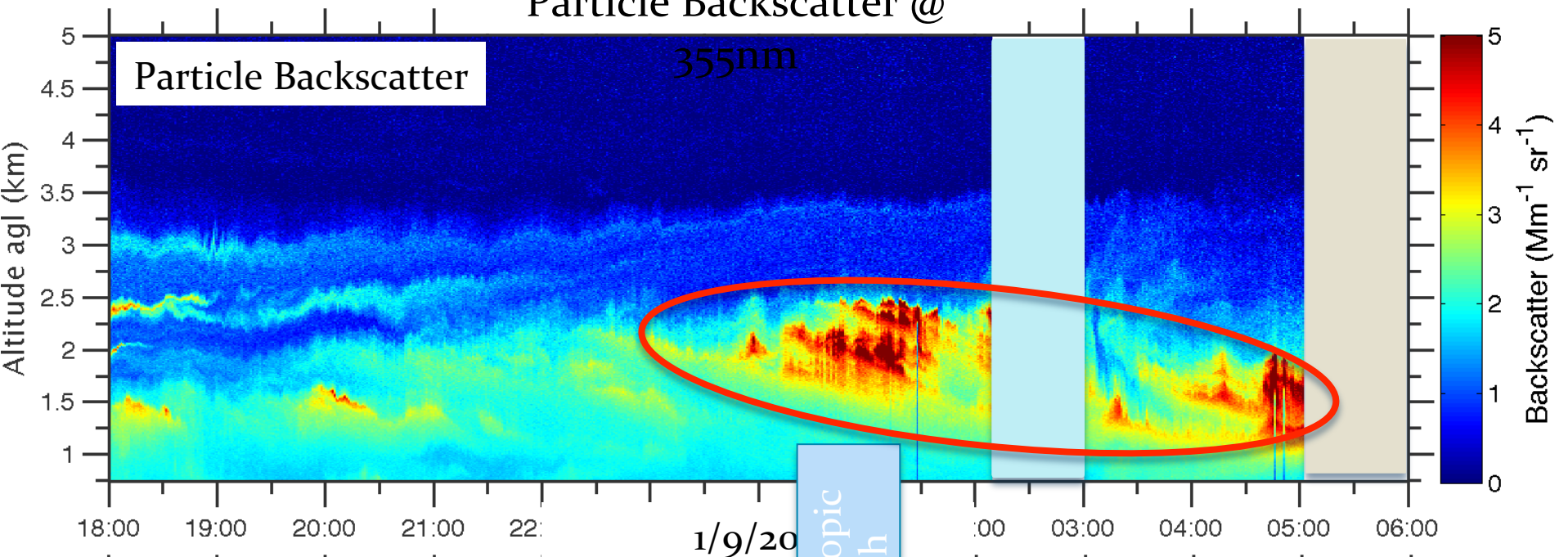


Example of Lidar Measurements

- Particle backscatter coefficient @ 355nm

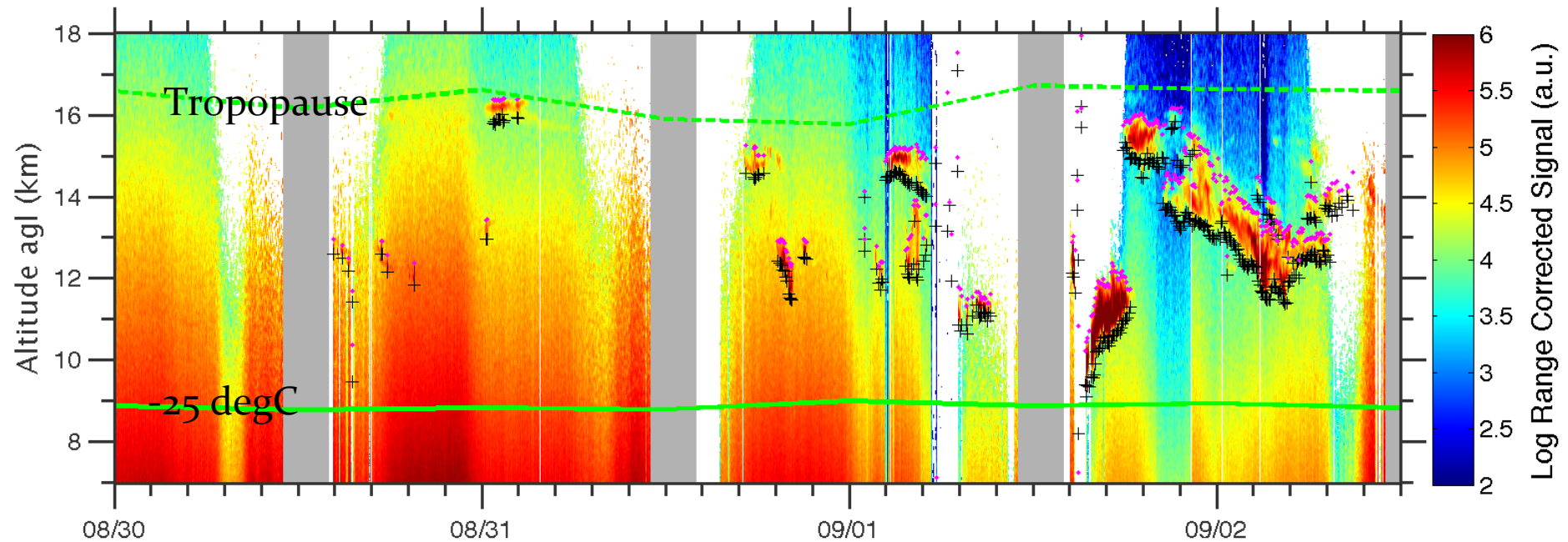


Particle Backscatter @



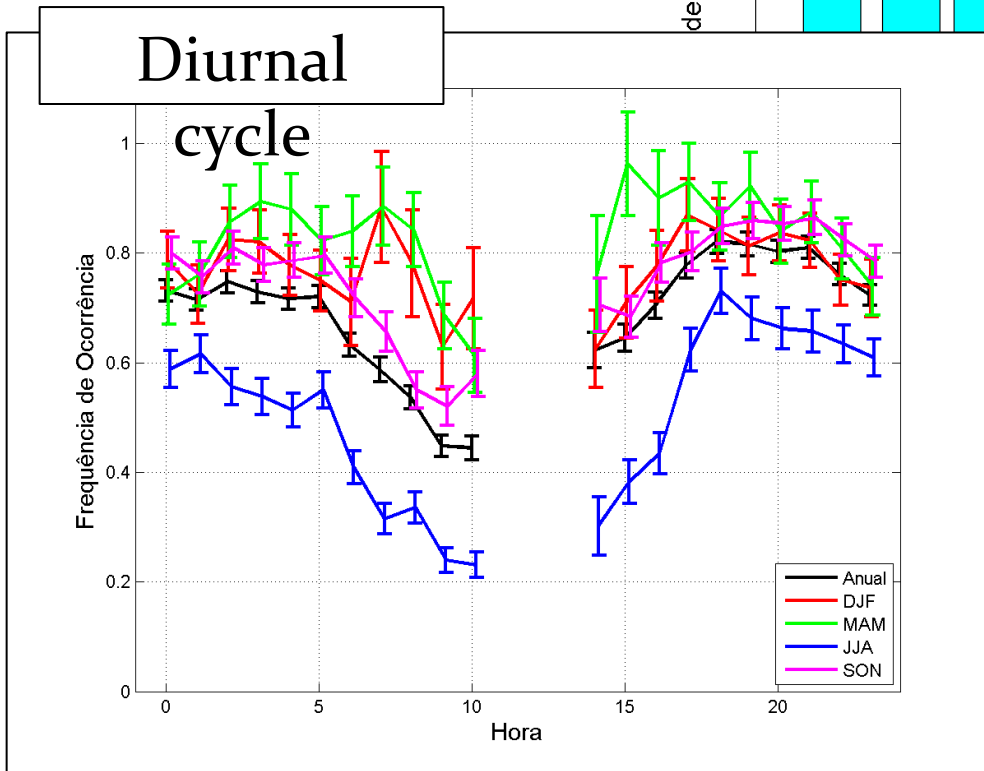
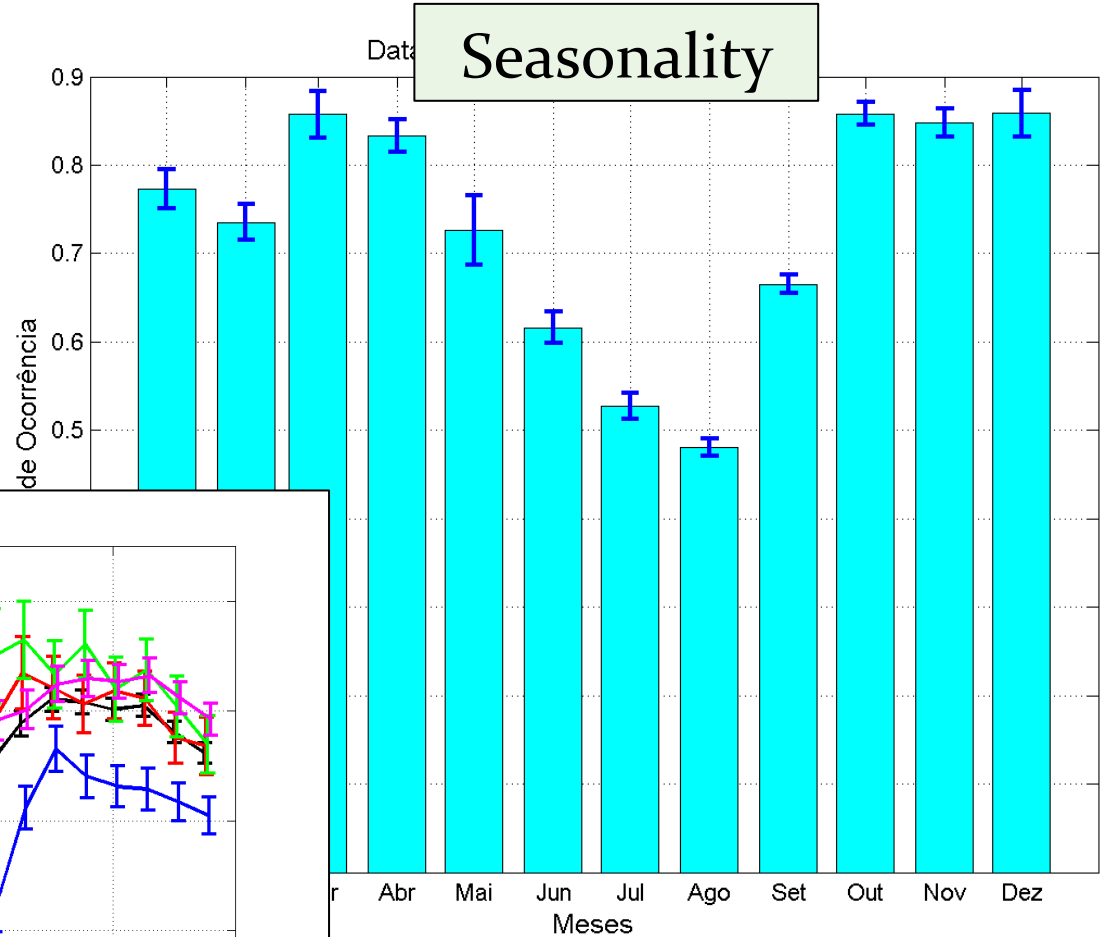
Cirrus Clouds

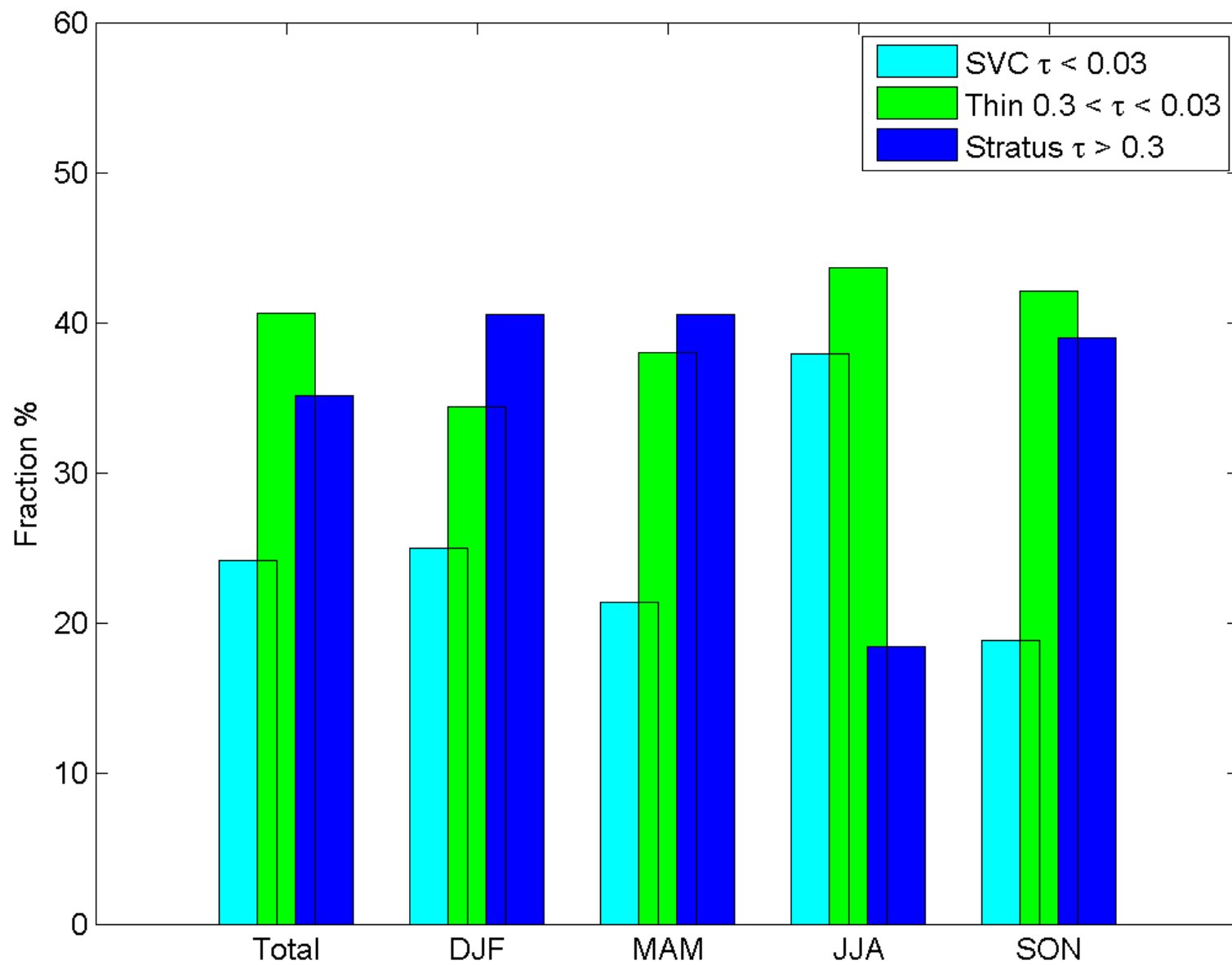
- Cirrus found from 8 to 19.6km
 - Base 12.5 ± 2.4 km
 - Top 14.2 ± 2.2 km



• Cirrus cloud cover at Manaus

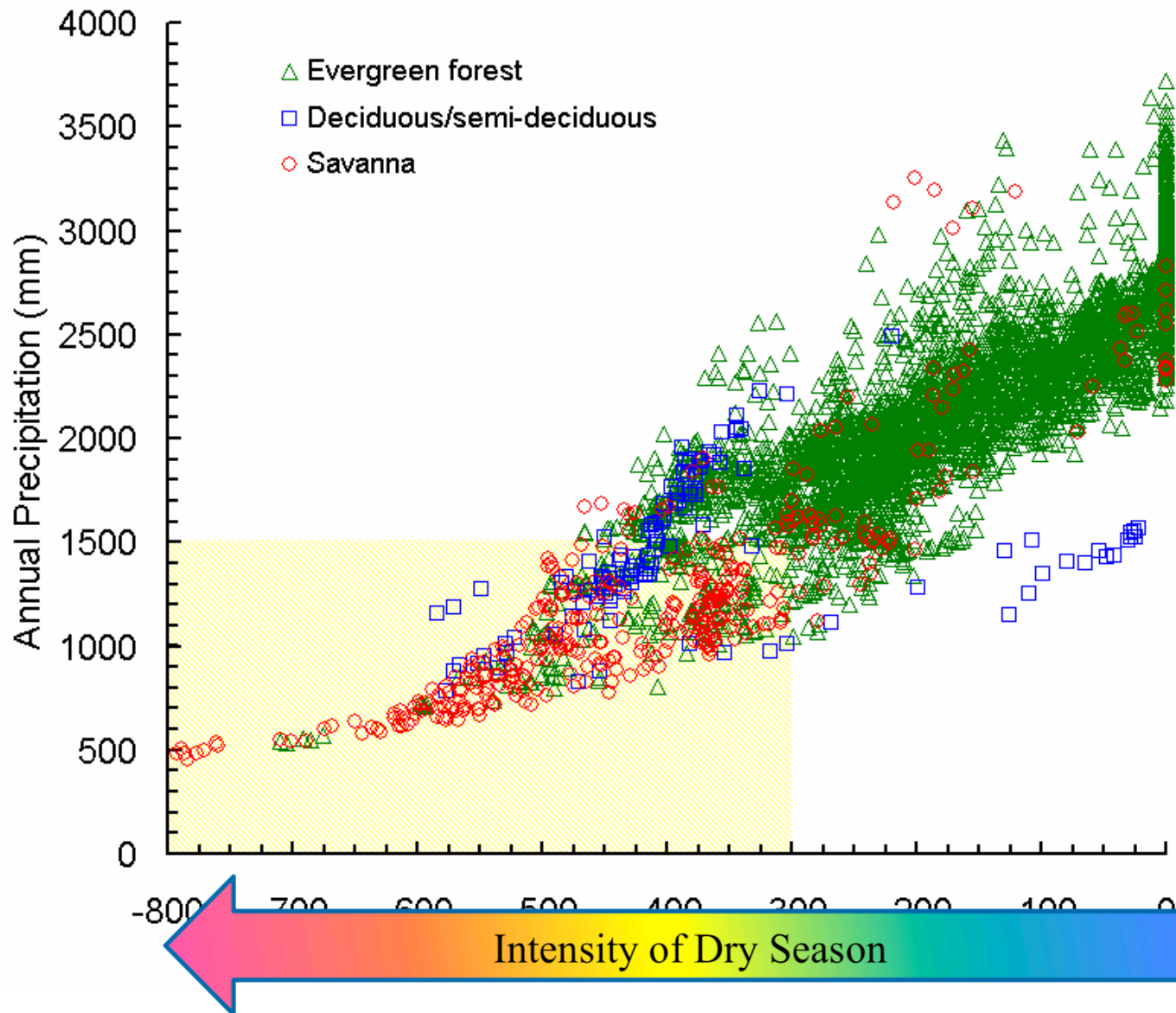
- 83% MAM
- 52% JJA





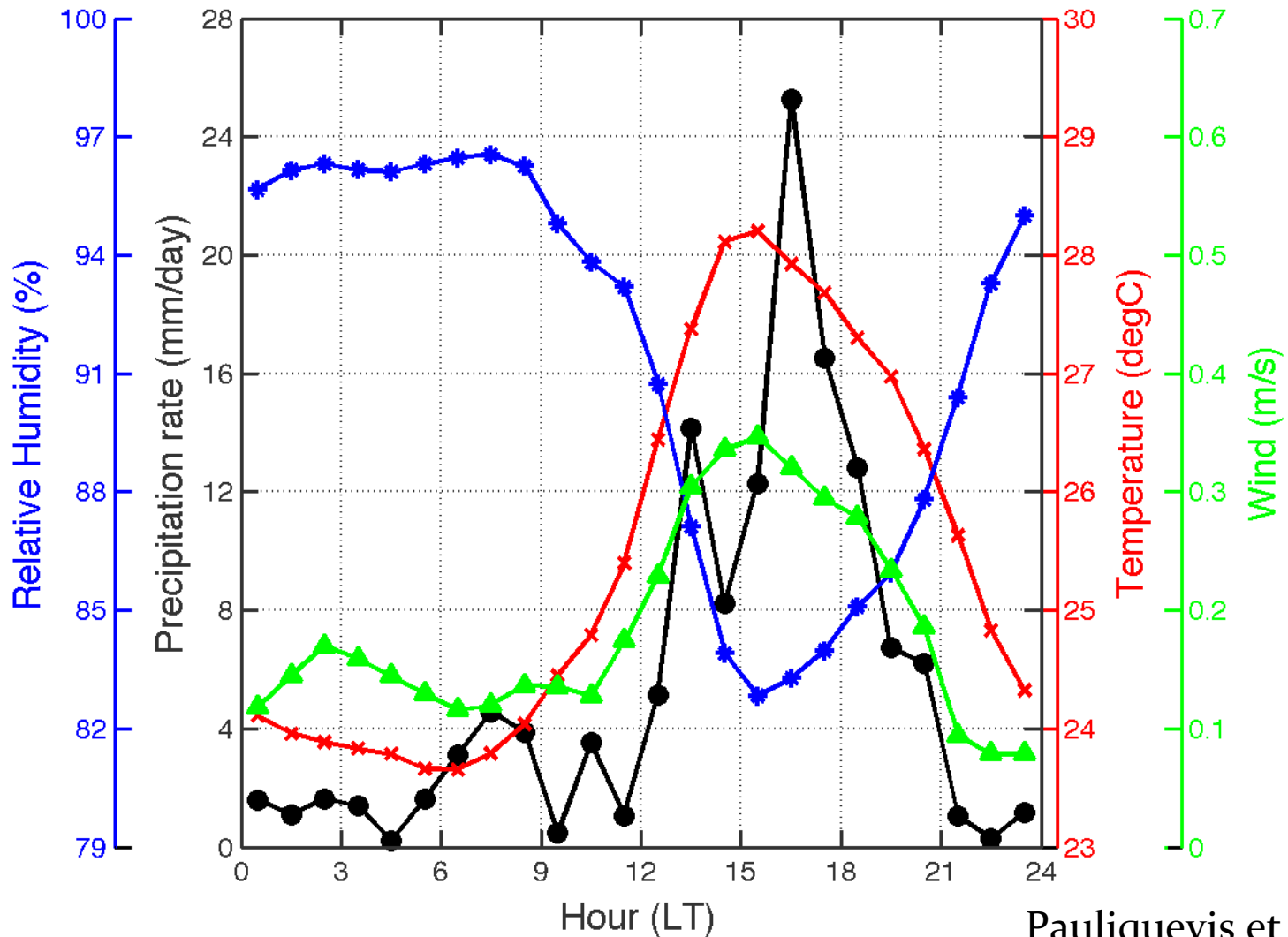
Frequency 78 % 83 % 52 % 77 %

A Rainfall Biogeography of Amazonia

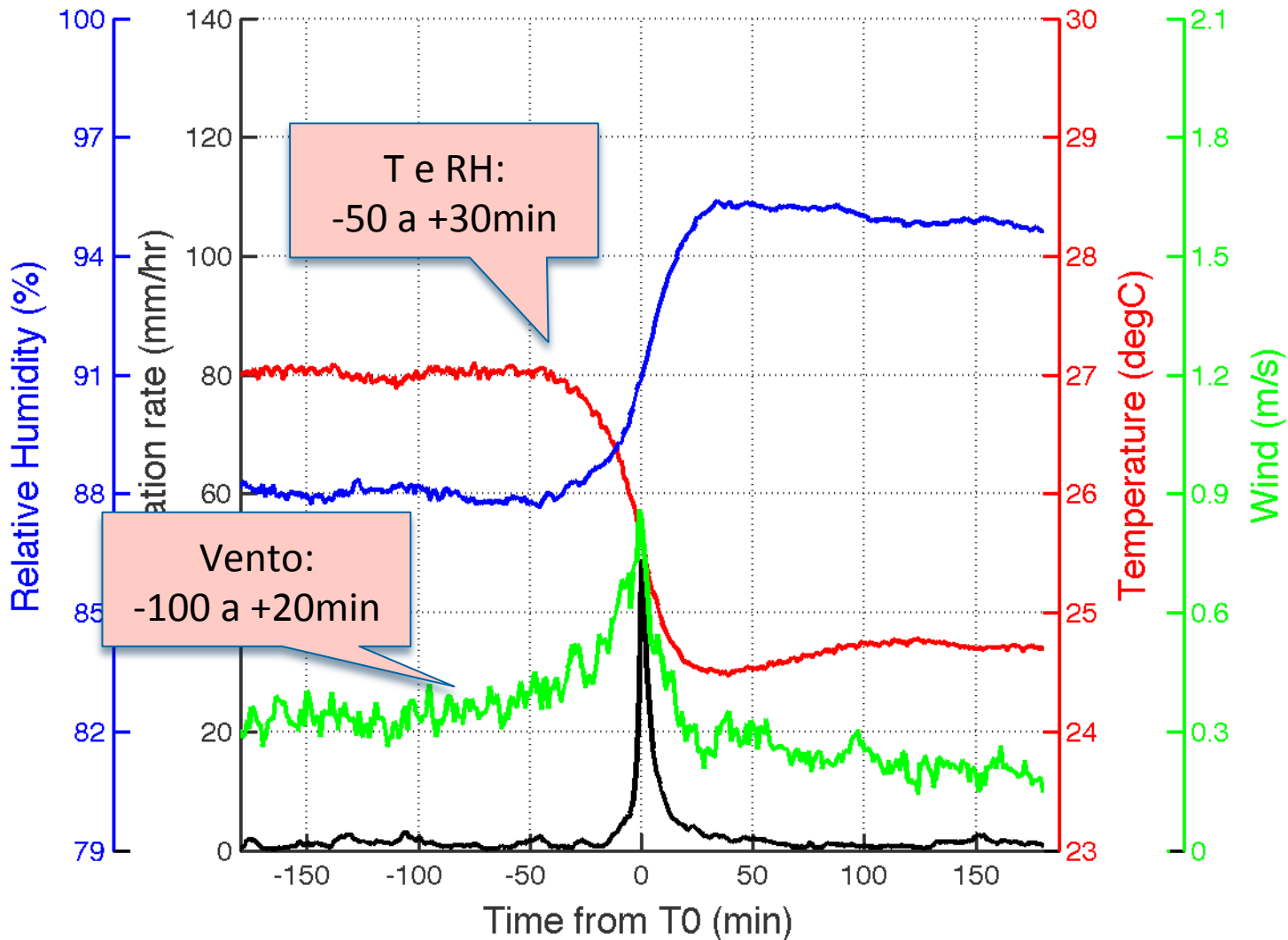


Source: Malhi *et al.*, **Exploring the likelihood and mechanism of a climate-change induced dieback of the Amazon rainforest**, *Proceedings of the National Academy of Sciences*, 2010

Diurnal Cycle @ Embrapa



@ Embrapa, T0= time of max precip



GNSS Dense Network

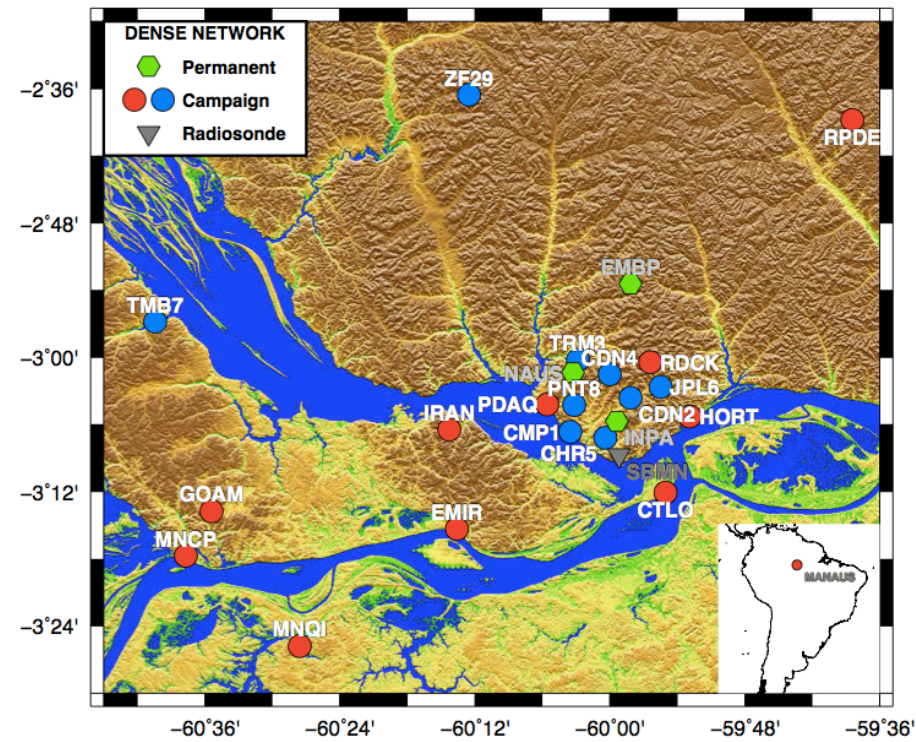
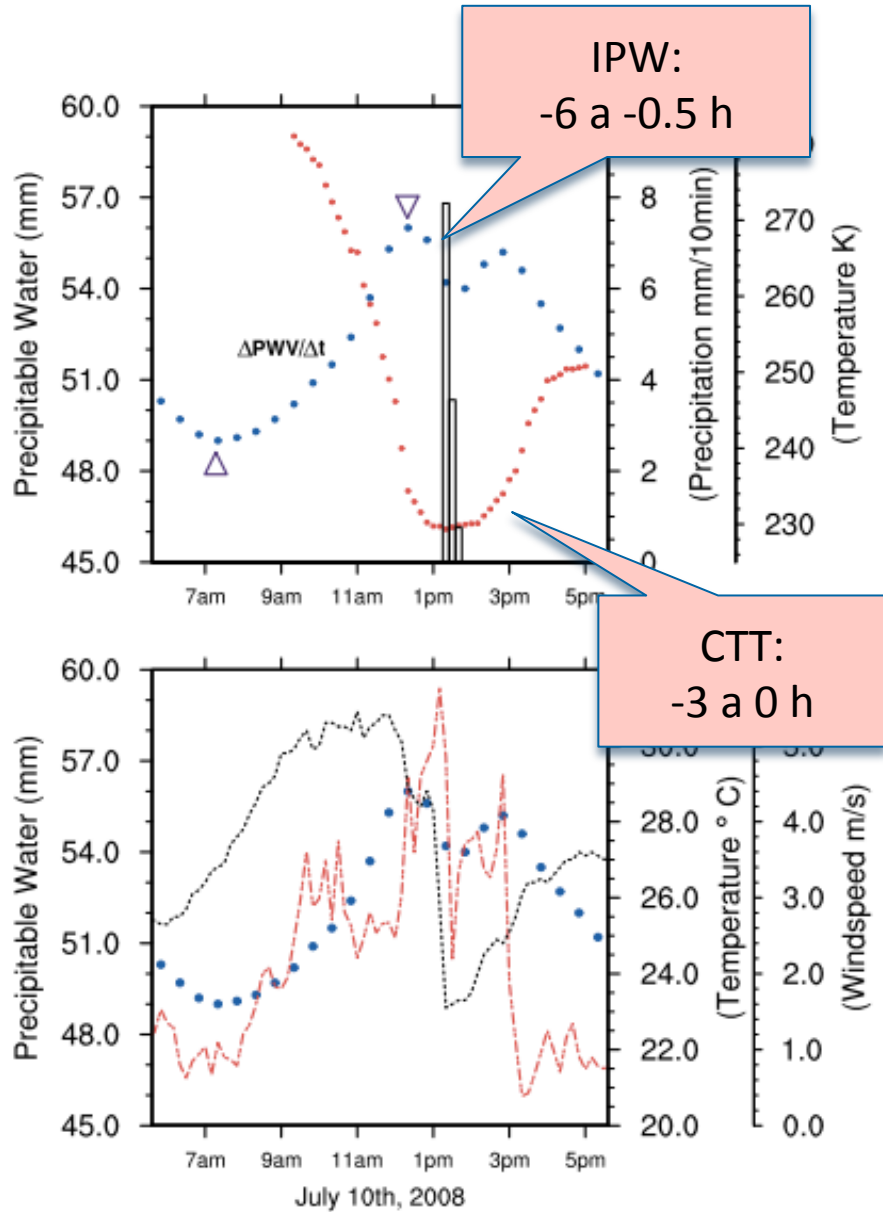


Figure 3. A typical afternoon deep convective event over INPA GNSS/meteorological station. The upper plot contains PWV (blue dots) versus average cloud top temperature (red) and precipitation rate (bars). The 'ramp-up' time calculated for the average $\Delta PWV/\Delta t$ (between triangles) represents the timescale of column convergence (see Equation (2) and text for discussion). The bottom graph plots wind speed (red), temperature (black) and PWV (blue) for the deep convective event.

- Adams et al, Atmos. Sci. Let. 2011
- Adams et al, BAMS 2014 (accepted)



**Amazonia is critical
for water vapor
transport over South
America**

**What processes controls these
fluxes?**

- Zemp et al, ACP 2014

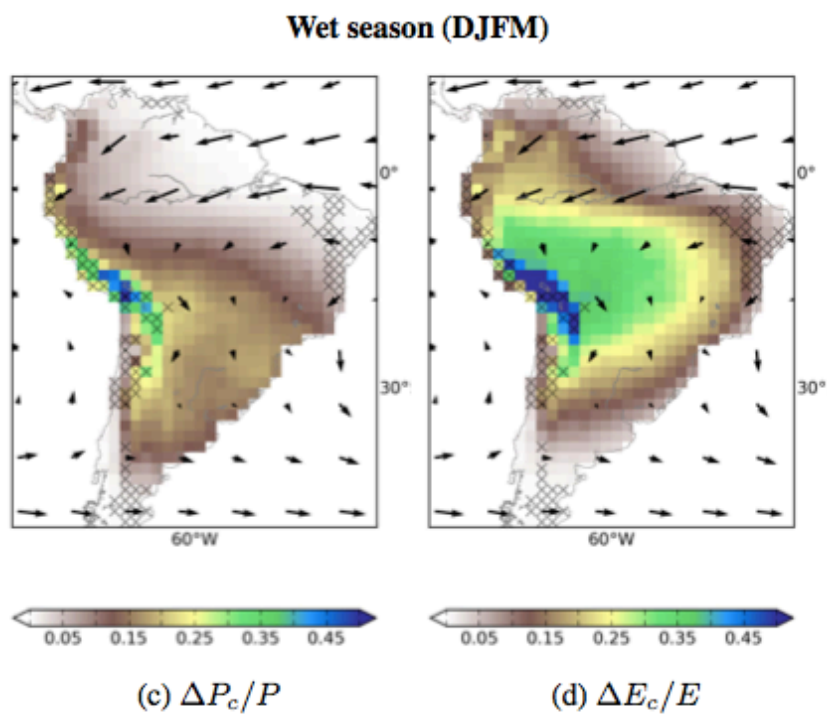
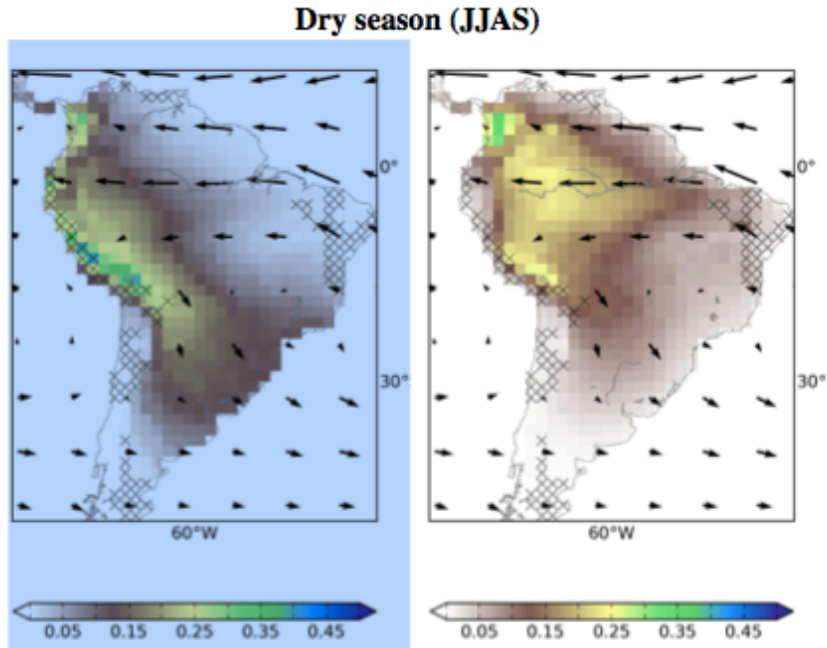
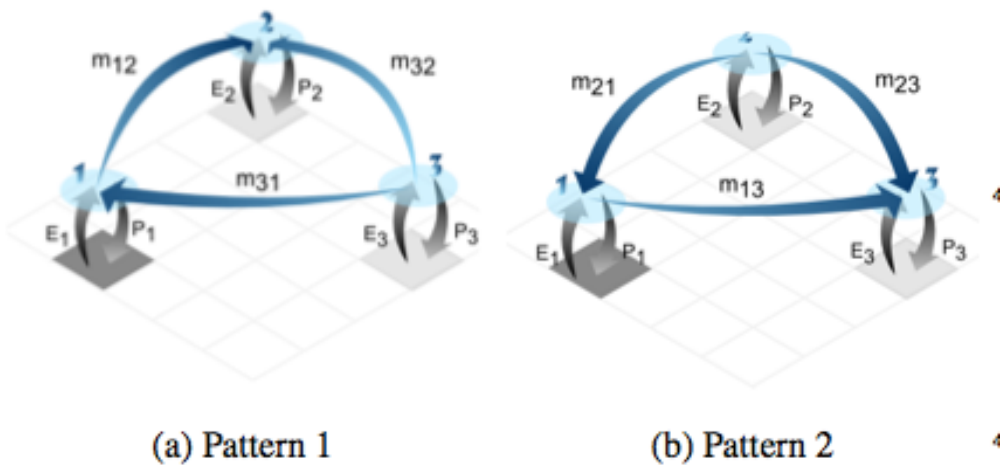
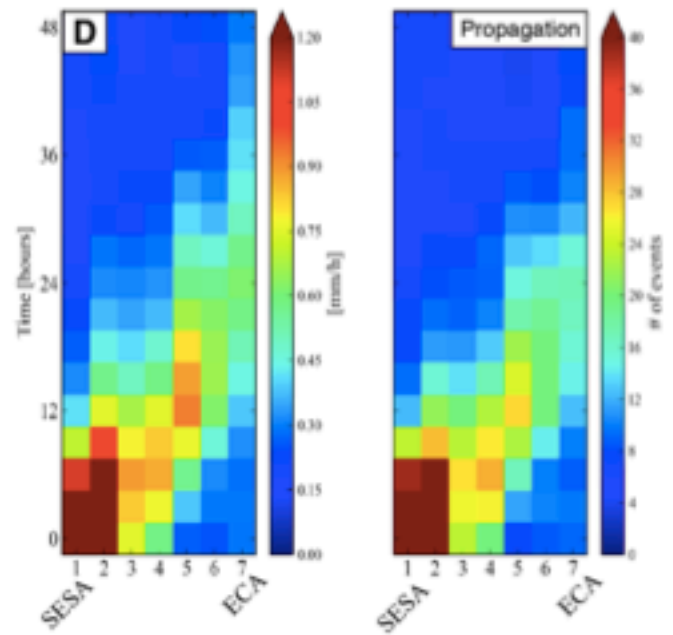
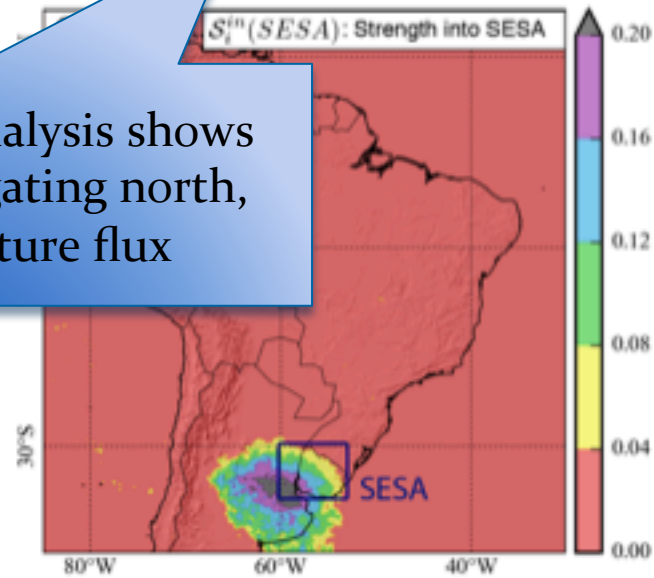
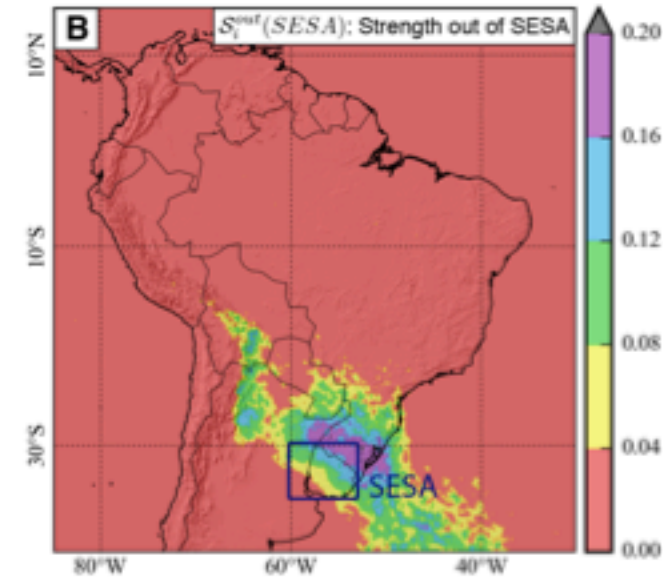
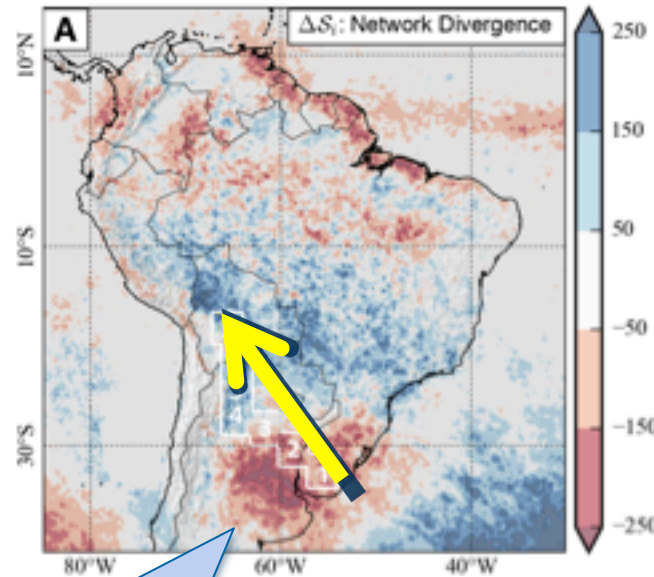


Fig. 5: Fraction of total precipitation originating from cascading moisture recycling ($\Delta P_c/P$) (a,c) and fraction of total evapotranspiration that is involved in cascading moisture recycling ($\Delta E_c/E$) (b,d). While high values of $\Delta P_c/P$ indicate regions which are dependent of cascading moisture recycling for local rainfall, high values of $\Delta E_c/E$ indicate regions which contribute to cascading moisture recycling.

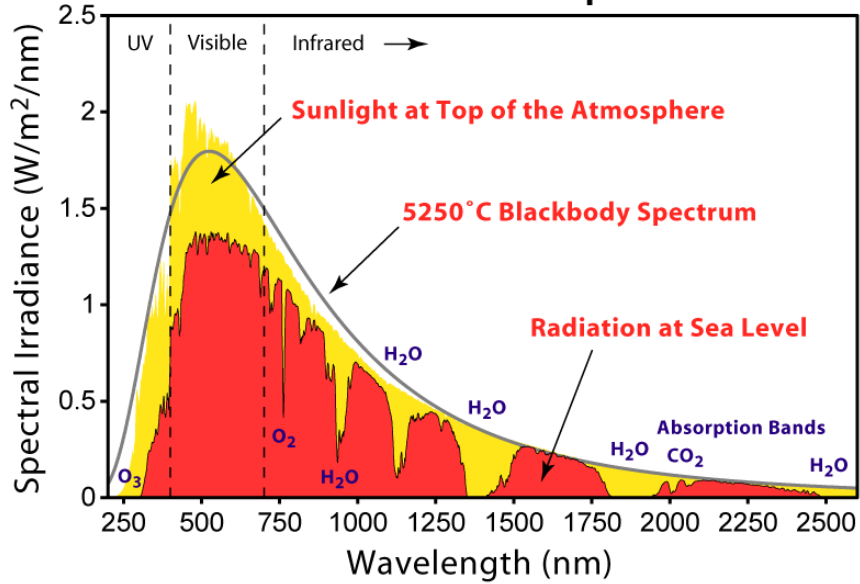
- Boers et al, Nature Comm. 2014



Complex network analysis shows precipitation propagating north, contrary to moisture flux



Solar Radiation Spectrum



hbarbosa@if.usp.br

www.fap.if.usp.br/~hbarbosa